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Bistable systems with stochastic noise: virtues and limits of effective one-dimensional Langevin equations

V. Lucarini^{1,2}, D. Faranda¹, and M. Willeit³

¹Klimacampus, Institute of Meteorology, Hamburg Universität, Hamburg, Germany

²Department of Mathematics and Statistics, University of Reading, Reading, UK

³Potsdam Institute for Climate Impact Research, Potsdam, Germany

Correspondence to: V. Lucarini (valerio.lucarini@zmaw.de)

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Abstract. The understanding of the statistical properties and of the dynamics of multistable systems is gaining more and more importance in a vast variety of scientific fields. This is especially relevant for the investigation of the tipping points of complex systems. Sometimes, in order to understand the time series of given observables exhibiting bimodal distributions, simple one-dimensional Langevin models are fitted to reproduce the observed statistical properties, and used to investigate the projected dynamics of the observable. This is of great relevance for studying potential catastrophic changes in the properties of the underlying system or resonant behaviours like those related to stochastic resonance-like mechanisms. In this paper, we propose a framework for encasing this kind of studies, using simple box models of the oceanic circulation and choosing as observable the strength of the thermohaline circulation. We study the statistical properties of the transitions between the two modes of operation of the thermohaline circulation under symmetric boundary forcings and test their agreement with simplified one-dimensional phenomenological theories. We extend our analysis to include stochastic resonance-like amplification processes. We conclude that fitted one-dimensional Langevin models, when closely scrutinised, may result to be more ad-hoc than they seem, lacking robustness and/or well-posedness. They should be treated with care, more as an empiric descriptive tool than as methodology with predictive power.

1 Introduction

An interesting property of many physical systems with several degrees of freedom is the presence of multiple equilibria (or, more in general, of a disconnected attractor) for a given choice of the parameters. In such a case, the system

does not obey ergodicity and its asymptotic state depends on what is the basin of attraction the initial condition belongs to. Among the many interesting properties of multi-stable systems, we may mention their possibility of featuring hysteretic behaviour: starting from an initial equilibrium $x = x_{\text{in}}$ realized for a given value of a parameter $P = P_{\text{in}}$ and increasing adiabatically the value of P so that the system is always at equilibrium following $x = x'(P)$, we may eventually encounter bifurcations leading the system to a new branch of equilibria $x = x'(P)$ such that, if we revert the direction of variation of P , we may end up to a different final stable state $x_{\text{fin}} = x'(P_{\text{in}}) \neq x(P_{\text{in}}) = x_{\text{in}}$. More generally, we can say that the history of the system determines which of the stable states is realized for a given choice of the parameters.

Whereas hysteretic behaviour has first been discussed in the context of magnetism, climate dynamics offers some outstanding examples where multistability is of great relevance, such as the classical problem of the snowball/snow-free Earth (Saltzman 2002; Lucarini et al., 2010; Pierrehumbert et al., 2011). In this context, the problem which has probably attracted the greatest deal of interest in the last two decades is that of the stability properties of the thermohaline circulation (THC). Since paleoclimatic evidences suggest that the large scale circulation of the Atlantic Ocean presents at least two, qualitatively different, stable modes of operation (Boyle et al., 1987; Rahmstorf, 2002), theoretical and modellistic efforts have long been directed to understanding the mathematical properties of the circulation and the physical processes responsible for switching from one to the other stable mode and those responsible for ensuring the stability of either equilibrium.

Interestingly, it has been possible to construct very simple models of the THC (Stommel, 1961; Rooth, 1982) able to feature most of the desired properties, and models of higher degrees of complexity have basically confirmed the robustness of such properties of multistability, from simple

two-dimensional convective equations models (Cessi and Young, 1992; Vellinga, 1996; Lucarini et al., 2005, 2007) to simplified climate models (Stocker and Wright, 1991; Rahmstorf, 1995; Stocker and Schmittner, 1997). Whereas climate models of intermediate complexity now consistently represent the THC as a multistable system (Rahmstorf et al., 2005), results are not conclusive when full 3-D climate models are considered (Stouffer and Manabe, 2003; Scott et al., 2008). Nonetheless, recent simulations performed by Hawkins et al. (2011) with a full 3-D climate model have successfully reproduced the kind of bistability properties shown by Rahmstorf et al. (2005). In the case of THC, the strength of the hydrological cycle plays the role of dominant parameter, whose variation can lead the system through bifurcations (Sijp and England, 2006, 2011; Sijp et al., 2011). A detailed account of these analyses can be found in Rahmstorf (1995), Scott et al. (1999), and Titz et al. (2001, 2002). The matter is of great relevance for understanding climate variability and climate change because, if the system is close to a bifurcation point, small changes in the parameters value could have virtually irreversible effects, driving the climate system a qualitatively different steady state. As evidenced in many studies (see, e.g. Kuhlbrodt et al., 2007), a transition from the present state of the THC to a state featuring weaker meridional circulation would have very relevant climatic effects at regional and global scale, as the northward ocean heat transport in the Atlantic would be greatly reduced.

The potential of shut-off of the THC is considered a high impact climate risk – even if its likelihood for the present climate is considered very low (Rahmstorf, 2006) – and the conditions under which such a transition can occur are probably the best example of a “climate tipping point” (Lenton et al., 2008).

Since we are dealing in principle with a 3-D fluid with complex thermodynamical and dynamical properties, a lot of efforts have been directing at finding, using suitable scaling and simplified theoretical setting, an approximate one-dimensional ordinary differential equation of the form $\dot{q} = F(q, P)$, where q is the intensity of the THC and P is a set of parameters of the system. Such an equation would be able to represent at least in a semi-quantitative way the evolution of the THC strength as a function of the strength and some parameters only and, by solving $F(q, P) = 0$, would provide the (in general) multiple equilibria corresponding to a specific choice of the set of parameters P . An excellent account of this methodology can be found in Dijkstra (2005). Note that, very recently, a related surrogate one-dimensional dynamics for a salinity indicator has been proposed to fit the output of a comprehensive climate model (Sijp et al., 2011).

The dynamics of multistable systems becomes rather interesting when stochastic forcing is considered. In the most basic case, such a forcing is represented in the form of additive white noise. Noise introduces on one side small scale variability around each of the stable equilibria of the system,

and, on the other side, allows for jumps (large scale variability) driving the system across the boundaries separating the basin of attraction of the fixed points. See Friedlin and Wentzel (1998) for a detailed mathematical treatment of these problems. Since the landmark Hasselmann’s (1976) contribution, it has become clearer and clearer in the climate science community that stochastic forcing components can be treated as quite reliable surrogates for high frequency processes not captured by the variables included in the climate model under consideration (Fraedrich, 1978; Saltzman, 2002). This has raised the interest in exploring whether transitions between stable modes of operation of the THC far from the actual tipping points could be triggered by noise, representing high-frequency (with respect to the ocean’s time scale) atmospheric forcings, of sufficient amplitude (Cessi, 1994; Monahan, 2002). Along these lines, it has become especially tempting to interpret the dynamics of the THC in presence of noise as resulting from an effective Langevin equation of the form $dq = F(q, p)dt + \varepsilon dW$ – where dW is the increment of a Wiener process – as this opens the way to approaching the problem in terms of one-dimensional Fokker-Planck equation. Ditlevsen (1999) suggested the possibility of considering more general stochastic processes for accommodating the statistical properties of observational data.

The Langevin equation approach has also led various researchers to study whether the process of stochastic resonance (Gammaitoni et al., 1998) – basically noise-enhanced response amplification to periodic forcing – could explain the strength of the climate response (in terms of actual THC strength) in spite of the relative weakness of the Milankovic forcing. Note that stochastic resonance, which has enjoyed great success in fields ranging from microscopic physics to neurobiology and perception, was first proposed in a climatic context (Benzi et al., 1982; Nicolis, 1982; Benzi, 2010). Velez-Belchi et al. (2001) provided the first example of a simple THC box-model featuring stochastic resonance, and later Ganopolski and Rahmstorf (2002) observed a similar mechanism in action in a much more realistic climate model.

In this work we would like to examine critically the effectiveness and robustness of using one-dimensional Langevin equations to represent the dynamical and statistical properties of the THC strength resulting from models which feature more than one degree of freedom. This is methodologically relevant, in the context of recent efforts directed at understanding whether the transitions between different steady states associated to the tipping points can be highlighted by early indicators (Scheffer et al., 2008). In Sect. 2 we show how to construct, from data, the form of the effective driving force and the intensity of noise, and propose tests for investigating the robustness of the approach. In Sect. 3 we describe the main properties of the simple box models of the oceanic circulation analysed in this paper. In Sect. 4 we test our methodology by studying, in various cases, whether it is possible to represent consistently the statistical properties

of the THC strength resulting from the stochastic forcing of the models introduced in Sect. 3 using an one-dimensional Langevin model. In Sect. 5 we further expand our analysis by testing whether the matching conditions for observing stochastic resonance are obeyed. In Sect. 6 we present our conclusions.

2 Theoretical background

2.1 Modelling a bistable system

Let's consider a one dimensional Langevin equation for the variable x of the form:

$$dx = F(x)dt + \varepsilon dW, \quad (1)$$

where $F(x)$ is a smooth function of x giving the drift term, W is a standard Wiener process and dW is its infinitesimal increment, so that ε parameterises the strength of the stochastic forcing. As well known, the invariant probability density function (pdf) $\pi(x)$ can be written as:

$$\pi(x) = C e^{-2\frac{V(x)}{\varepsilon^2}} \quad (2)$$

where $V(x)$ is the effective potential such that $dV(x)/dx = -F(x)$ and C is the normalisation constant. The local extrema of the potential correspond to the fixed points of the deterministic system obtained when $\varepsilon = 0$, and in particular its local minima (maxima) giving the stable (unstable) equilibria. Quite intuitively, in the stochastic case, the peaks of the invariant probability distribution correspond to the minima of the potential.

In the prototypical situation of a confining double well potential, where we refer to the position of the right and left minimum of $V(x)$ as x_+ , x_- , and to local maximum as x_0 , the two peaks of the $\pi(x)$ are separated by a dip corresponding to the local maximum of the potential, while for large positive and negative values of x the $\pi(x)$ approaches zero as $V(x)$ diverges for $|x| \rightarrow \infty$. The average rate of transition $r(+ \rightarrow -)$ from the basin of attraction of x_+ to that of x_- can be approximated using the Kramer's formula:

$$r(+ \rightarrow -) = \frac{1}{2\pi} \sqrt{V''(x_+)V''(x_0)} e^{-2\frac{V(x_0)-V(x_+)}{\varepsilon^2}} \quad (3)$$

under the condition that the absolute value of the exponent is larger than one, so that $V(x_0) - V(x_+) \geq \varepsilon^2/2$. This corresponds to the physical condition that the noise is moderate with respect to the depth of the potential well. The average rate of transition $r(- \rightarrow +)$ in the opposite direction can be obtained by exchanging the sign plus with the sign minus in the previous expression. The Kramer's formula basically expresses the general fact that at stationary state a detailed balance conditions applies.

Let's now consider the less idealised case where we observe a scalar output signal $y(x)$ generated by a stochastic or chaotic deterministic flow of the system x living, in general, in an N -dimensional phase space, and let's assume that the empirical pdf $\pi(y)$ is bimodal, so that two peaks are found at $y = y_+$, $y = y_-$, separated by a local minimum at $y = y_0$. We wish to test the possibility of constructing a Langevin equation for the scalar y :

$$dy = F_{\text{eff}}(y)dt + \varepsilon_{\text{eff}}dW \quad (4)$$

so that the statistics generated by Eq. (4) closely resembles the one deriving from the full system. Such a representation would bypass the details of the full dynamics of the system and its construction can be approached by imposing constraints based upon the populations of the basin of attractions of the two modes and upon the transition probability between such basins. Basically, this amounts to defining an effective projected dynamics.

The observation time of the variable y must be long enough to allow for a robust estimate of the pdf and for observing many transitions between the two modes. From the empirical pdf of the considered observable $\pi(y)$ we derive the function $U(y) = -\ln(\pi(y))$. In order to achieve compatibility with Eq. (2), we define

$$2V_{\text{eff}}(y)/\varepsilon_{\text{eff}}^2 = U(y) + \text{const.} \quad (5)$$

The function $U(y)$ contains information on both the effective potential of the system and on the effective intensity of the noise. Substituting the expression of $V_{\text{eff}}(y)$ in Eq. (3), we can write, e.g. the average rate $r(+ \rightarrow -)$ as follows:

$$r(+ \rightarrow -) \approx \frac{\varepsilon_{\text{eff}}^2}{4\pi} \sqrt{U''(y_+)U''(y_0)} e^{-(U(y_0)-U(y_+))} \quad (6)$$

By comparing this expression with the observed transition rate, we can finally find the actual value of ε_{eff} , because all the other terms can be computed from the time series of the y observable. Assuming that in the region between the two minima and the local maximum the function $U(y)$ is smooth, we obtain the following expression, which is numerically more robust as the second derivatives disappear:

$$\begin{aligned} r(+ \rightarrow -) &\approx \frac{\sqrt{2}\varepsilon_{\text{eff}}^2}{\pi} \frac{(U(y_0) - U(y_+))}{(y_+ - y_0)^2} e^{-(U(y_0)-U(y_+))} \\ &= \frac{2\sqrt{2}}{\pi} \frac{(V(y_0) - V(y_+))}{(y_+ - y_0)^2} e^{-\left(2\frac{V(y_0)-V(y_+)}{\varepsilon_{\text{eff}}^2}\right)} \end{aligned} \quad (7)$$

where $G(+ \rightarrow -) = \frac{\sqrt{2}}{\pi} \frac{(U(y_0)-U(y_+))}{(y_+-y_0)^2} e^{-(U(y_0)-U(y_+)}}$ is a factor depending only on the observed probability. Note that Eqs. (6)–(7) are valid under the condition that $U(y_0) - U(y_+) \gtrsim 1$ or, equivalently, that $\pi(y_+)/\pi(y_0) \gtrsim e$. By plugging the obtained value of ε_{eff} into Eq. (5) we derive $V_{\text{eff}}(y)$. We can then reconstruct the effective Langevin equation in the form given in Eq. (4). It is crucial that the same value for

ε_{eff} is obtained when using as benchmark the average rate $r(- \rightarrow +)$, because otherwise we have that the reconstructed dynamics does not obey detailed balance, or, equivalently, there is a mismatch between time-scales of the transitions and steady state populations of the two basins. Such an issue is not relevant in the special case when the pdf of the system is symmetric with respect to y_0 .

2.2 Robustness

Let's now consider an N-dimensional Langevin equation of the form:

$$dx_i = F_i(x_1, \dots, x_N)dt + \varepsilon_{ij}dW_j \quad (8)$$

where the dW_j terms indicate increments of independent Wiener processes and the $F_i(x_1, \dots, x_N)$ are the (generally nonlinear) drift terms. We can write the Langevin equation for the observable y constructed as linear combination of the system variables $y = c_i x_i$ as follows:

$$\begin{aligned} dy &= c_i F_i(x_1, \dots, x_N)dr + c_i \varepsilon_{ij}dW_j \\ &= c_i F_i(x_1, \dots, x_N)dt + \bar{\varepsilon}dW \end{aligned} \quad (9)$$

where, exploiting the independence of the Wiener processes, we have $\bar{\varepsilon} = \sqrt{\sum_{j=1}^N (c_i \varepsilon_{ij})^2}$. If the deterministic part of system (9) features two stable equilibria \mathbf{x}_+ and \mathbf{x}_- , when stochastic noise is added, we will see hopping between the basins of attraction of these two points. When looking at the variable $y = c_i x_i$ as output of the system, we will see a bimodal distribution where the two peaks are centred at $y_+ = c_i x_{i,+}$ and $y_- = c_i x_{i,-}$, respectively, with a local minimum in-between situated at y_0 .

The true dynamics of the observable y is indeed given in Eq. (9), but, if we are provided only with the time series of y , the best we can do in to try to derive the ‘‘best’’ approximate equation – the one-dimensional Langevin model – of the form given in Eq. (4), using the heuristic procedure described above. Comparing Eqs. (4) and (9), we understand that the difference in the drift terms $c_i F_i(x_1, \dots, x_N) - F_{\text{eff}}(y)$ describes the deterministic dynamics of the y variable which cannot be parameterised in terms of the y variable alone. We may expect that if y is a slow variable, which retains the long term memory of the system, such a difference is small, and the dynamics of y is truly quasi-one-dimensional. In there is time-scale separation, one may expect that the impact of the faster variables on its evolution can be effectively expressed as noise, thus determining the value of $\bar{\varepsilon}$ (Saltzman, 2002). Therefore, the ratio $\bar{\varepsilon}/\varepsilon_{\text{eff}}$ gives an indirect measure of how strong is this effect, being close to 1 if the dominant contribution for comes from the direct stochastic forcing into the system.

Since constructing an ad-hoc one-dimensional model from a time-series is typically possible by following the lines described above, we need to introduce some criteria to test the robustness of the approach we have undertaken. This is

needed in order to check whether our methodology is solely descriptive of the statistics of the chosen observable system for a given choice of parameters, or, instead, has predictive power, in terms of allowing one to understand how the statistical properties of the observable change when the parameters of the full system are altered. Obviously, as mentioned above, choosing a suitable variable y will be crucial in ensuring the effectiveness of this one-dimensional parameterisation:

- As a first condition of robustness, we may ask that if we multiply the noise intensity of the true system by a factor α , so that $\varepsilon_{ij} \rightarrow \alpha \varepsilon_{ij}$ in Eq. (8), we obtain that correspondingly $\varepsilon_{\text{eff}} \rightarrow \alpha \varepsilon_{\text{eff}}$ and the effective potential $V_{\text{eff}}(y)$ is not altered in Eq. (4).
- Another condition of robustness is that if in Eq. (9) we alter the noise matrix ε_{ij} in such a way that $\bar{\varepsilon}$ is not altered in Eq. (9), we would like that, correspondingly, ε_{eff} and V_{eff} are not altered in Eq. (4).

These conditions basically require that the reconstructed deterministic drift term is independent of the intensity of the noise – so that an underlying deterministic dynamics is well defined – and the reconstructed noise intensity scales linearly with the actual noise applied to the system, so that we can construct a relationship such as $\varepsilon_{\text{eff}}(\bar{\varepsilon}) \approx \gamma \bar{\varepsilon}$

3 Simple box models of the thermohaline circulation

3.1 Full model

We consider the simple deterministic three-box model of the deep circulation of the Atlantic Ocean introduced by Rooth (1982) and thoroughly discussed in Scott et al. (1999) and Lucarini and Stone (2005a, b). The model consists of a northern high-latitude box (box 1), a tropical box (box 2), and a southern high-latitude box (box 3). The volume of the two high-latitude boxes is the same and is $1/V$ times the volume of the tropical box, where V is chosen to be equal to 2. The physical state of the box i is described by its temperature T_i and its salinity S_i ; the box i receives from the atmosphere the net flux of heat H_i and the net flux of freshwater F_i ; the freshwater fluxes globally sum up to 0, so that the average oceanic salinity is a conserved quantity of the system. The box i is subjected to the oceanic advection of heat and salt from the upstream box through the THC, whose strength is q . The dynamics of the system is described by the evolution equation for the temperature and the salinity of each box. After a suitable procedure of non-dimensionalisation (Lucarini and Stone, 2005a), we obtain the following final form for the temperature and salinity tendency equations for the three boxes:

$$\dot{T}_1 = \begin{cases} q(T_2 - T_1) + H_1 \\ |q|(T_3 - T_1) + H_1 \end{cases} \quad (10a)$$

$$\dot{T}_2 = \begin{cases} \frac{q}{V}(T_3 - T_2) + H_2 \\ |\frac{q}{V}|(T_1 - T_2) + H_2 \end{cases} \quad (10b)$$

$$\dot{T}_3 = \begin{cases} q(T_1 - T_3) + H_3, & q > 0 \\ |q|(T_2 - T_3) + H_3, & q \leq 0 \end{cases} \quad (10c)$$

$$\dot{S}_1 = \begin{cases} q(S_2 - S_1) - F_1 \\ |q|(S_3 - S_1) - F_1 \end{cases} \quad (10d)$$

$$\dot{S}_2 = -\frac{1}{V}(\dot{S}_1 + \dot{S}_3) \quad (10e)$$

$$\dot{S}_3 = \begin{cases} q(S_1 - S_3) - F_3, & q > 0 \\ |q|(S_2 - S_3) - F_3, & q \leq 0 \end{cases} \quad (10f)$$

where $q = k(\rho_1 - \rho_3)$ and $\rho_i = \rho_0(1 - \alpha T_i + \beta S_i)$, so that $q = k(\beta(S_1 - S_3) - \alpha(T_1 - T_3))$, α and β are the usual thermal and haline expansion coefficients, ρ_0 is a baseline density, and k is the hydraulic constant controlling the water transport. Such a parameterisation was first introduced by Stommel (1961) for a hemispheric box model, whereas the approximate linear relationship between the density difference between the two high-latitude regions and the THC strength has been confirmed by simplified yet realistic GCM simulations (Rahmstorf, 1995; Scott et al., 2008).

The freshwater fluxes F_i are considered given constants, with $F_2 = -(F_1 + F_3)/V$, so that $S_1 + V S_2 + S_3 = (V + 2)S_0$ at all times, where $S_0 = 35$ psu is a baseline salinity. Instead, the heat flux $H_i = \lambda(\bar{T}_i - T_i)$ is such that the box temperature is relaxed to a fixed target temperature \bar{T}_i with the time constant λ^{-1} . Such a representation mimics the combined effect of radiative heat transfer and of a meridional heat transport. This also implies that the spatial average of ocean temperature $\bar{T} = (T_1 + V T_2 + T_3)/(2 + V)$ obeys the evolution equation $\dot{\bar{T}} = \lambda(\bar{T} - \bar{T})$, where \bar{T} is the spatial average of the target temperature, so that asymptotically (and, in practise, after few units of λ^{-1}) \bar{T} is a conserved quantity. Therefore, we practically have $\dot{\bar{T}} \approx -(\dot{T}_1 + \dot{T}_3)/V$. We usually have that the internal time scale of the system q^{-1} is much larger than λ^{-1} , thus implying that the thermal relaxation is fast.

The water of the high latitude ocean box where downwelling occurs is warmer and more saline than situated on the opposite side of the planet, since it receives advection from the warm and saline equatorial box. Since the haline contribution is stronger, the downwelling box is denser than the upwelling box. The sign of q is positive if downwelling occurs in box 1, and negative if it occurs in box 3. The buoyancy fluxes in the upwelling and downwelling boxes serve different purposes in determining the dynamics of the system. In fact, the strength of the circulation at equilibrium q_{ref}

depends only on the strength of the buoyancy fluxes in the upwelling box $H_{u,\text{eq}}$ and F_u :

$$|q_{\text{ref}}| = \sqrt{k(\alpha H_{u,\text{eq}} + \beta F_u)},$$

where the sign of q_{eq} is positive if $u = 1$ and negative if $u = 3$. Instead, for a given value of F_u the realised pattern of circulation is stable as long as $F_d \lesssim 3F_u$, which implies that $F_d = 3F_u$ a bifurcation leading to an instability of the system is found. Such an instability exchanges the role of the upwelling and downwelling boxes. Therefore, if, e.g. in the initial state $u = 3$ and $d = 1$, and F_3 is kept fixed, the system features bistability for the following range of values of F_1 : $1/3 F_3 \lesssim F_1 \lesssim 3F_3$. These approximate relations become exact in the limit of infinitely fast thermal relaxation.

Following Scott et al. (1999) and Lucarini and Stone (2005a), we select for the constants of the system the values $k = 1.5 \times 10^{-6} \text{ s}^{-1}$, $\alpha = 1.5 \times 10^{-4} \text{ K}^{-1}$, $\beta = 8.0 \times 10^{-4} (\text{psu})^{-1}$, $\lambda = 1.3 \times 10^{-9} \text{ s}^{-1}$, $V = 2$. When symmetric boundary conditions are considered with $\bar{T}_1 = \bar{T}_3 = 0^\circ\text{C}$, $\bar{T}_2 = 30^\circ\text{C}$, $F_1 = F_3 = 9 \times 10^{-11} \text{ psu s}^{-1}$, we obtain at steady state $|q_{\text{ref}}| = 1.47 \times 10^{-11} \text{ s}^{-1}$. The sign of q depends uniquely on the initial conditions of the integration: we have 50% probability of ending up in either the northern or the southern downwelling state if random initial conditions are chosen. Since the internal time scale $|q_{\text{ref}}|^{-1} \approx 215 \text{ y}$ is much larger than the thermal time scale $\lambda^{-1} \approx 25 \text{ y}$, we conclude that the thermal relaxation is a fast process.

The physical value \tilde{q} of the strength of the thermohaline circulation can be found from the normalised value above as $\tilde{q} = q V_{\text{box},1}$, where $V_{\text{box},1} = V_{\text{box},3} = V_{\text{box},2}/V = 1.1 \times 10^{17} \text{ m}^3$ is the volume of either high-latitude box. Instead, the physical value of the net freshwater flux \tilde{F}_i into box i is obtained as $\tilde{F}_i = F_i V_{\text{box},i}/S_0$; its intensive value per unit surface results to be $\tilde{F}_i = F_i V_{\text{box},i}/A_{\text{box},i} S_0 = F_i D_{\text{box}}/S_0$ where $D_{\text{box}} = 5000 \text{ m}$ is the common depth of the three oceanic boxes. Therefore, our base state features reasonable values for the net poleward transport of freshwater flux – about $2.8 \times 10^5 \text{ m}^3 \text{ s}^{-1} = 0.28 \text{ Sv}$, and for the THC strength – about $1.55 \times 10^7 \text{ m}^3 \text{ s}^{-1} = 15.5 \text{ Sv}$.

3.2 Simplified model

A simplified version of the model given in Eq. (10a–f) can be derived by assuming that the thermal restoring constant $\lambda \rightarrow \infty$ so that the time scale of the feedback $\lambda^{-1} \rightarrow 0$. Thus, the temperatures of the three boxes are such that at all times $T_i = \bar{T}_i$, so that we obtain a reduced dynamical system with only 2 degrees of freedom (d.o.f.):

$$\dot{S}_1 = \begin{cases} q(2S_0 - 3/2S_1 - 1/2S_3) - F_1 & q > 0 \\ |q|(S_3 - S_1) - F_1 & q \leq 0 \end{cases} \quad (11a)$$

$$\dot{S}_3 = \begin{cases} q(S_1 - S_3) - F_3 & q > 0 \\ |q|(2S_0 - 3/2S_3 - 1/2S_1) - F_3 & q \leq 0 \end{cases} \quad (11b)$$

where the THC strength can be written as $q = k\beta(S_3 - S_1)$. Note that, since the system (11a–b) has been obtained by performing a singular perturbation to (10a–f), we need to renormalize the value of the hydraulic constant k in order to obtain $q_{\text{ref}} = 1.47 \times 10^{-11} \text{ s}^{-1}$ at steady state when choosing $F_1 = F_3 = 9 \times 10^{-11} \text{ psu s}^{-1}$ as above. The resulting value is $k = 3.0 \times 10^{-7} \text{ s}^{-1}$, with $|q_{\text{ref}}| = \sqrt{k\beta F_U} = \sqrt{k\beta F_1}$. The same physical scalings described above apply here. Such a simplified model retains the most relevant elements of the dynamics of the full model, even if the thermal dynamical feedback (Scott et al., 1999) are missing.

4 Numerical experiments: fitting the dynamics from the population and the transition rates

4.1 Symmetric forcing to the simplified model

We now modify the dynamical system (11a–b) by including additive noise in both the evolution equations for both S_3 and S_1 so that we obtain the following system of stochastic differential equations in the Ito form:

$$dS_3 = \begin{cases} q(2S_0 - 3/2S_1 - 1/2S_3)dt - F_1dt + \varepsilon_1dW_1, & q > 0 \\ |q|(S_3 - S_1)dt - F_1dt + \varepsilon_1dW_1 & q \leq 0 \end{cases} \quad (12a)$$

$$dS_1 = \begin{cases} q(S_1 - S_3)dt - F_3dt + \varepsilon_3dW_3, & q > 0 \\ |q|(2S_0 - 3/2S_3 - 1/2S_1)dt - F_3dt + \varepsilon_3dW_3 & q \leq 0 \end{cases} \quad (12b)$$

where $dW_{1,3}$ are the increments of two independent Wiener processes. Note that Ditlevsen (1999) proposed the possibility of considering more general noise processes to explain the THC dynamics. Hereby, we stick to the more usual white noise case.

We then perform a set of experiments by integrating the stochastic differential Eq. (12a–b) using the numerical scheme proposed in Mannella and Pallechi (1989) for values of $\varepsilon_1 = \varepsilon_3 = \varepsilon$ ranging from $3.6 \times 10^{-10} \text{ psu s}^{-1/2}$ to $6.2 \times 10^{-10} \text{ psu s}^{-1/2}$. This corresponds to a range of noise strength for the physical freshwater flux of $1.12 \times 10^6 \text{ m}^3 \text{ s}^{-1/2}$ to $1.96 \times 10^6 \text{ m}^3 \text{ s}^{-1/2}$. In more concrete terms, we are exploring stochastic perturbations to the freshwater flux whose variability (standard deviation), over the characteristic internal time scale $|q_{\text{ref}}|^{-1} \approx 215 \text{ yr}$, range between 27 % and 47 % of the baseline value $F_1 = F_3$. Results are presented for sets of 100 ensemble members for each value of ε , with each integration lasting 10^6 yr . The chosen time step is 1 yr.

We wish to study the possibility of defining up to a good degree of precision a consistent stochastic dynamics for the THC strength q involving only q itself and noise. Following

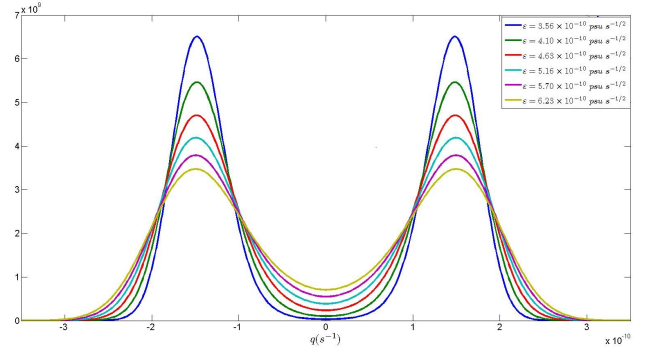


Fig. 1. Empirical probability distribution function for the THC strength in the 2 d.o.f. model for selected values of symmetrically applied noise.

the procedure outlined in Sect. 2, for each value of $\varepsilon_1 = \varepsilon_3 = \varepsilon$, where we have on purpose kept the system's parameters invariant with respect to exchanging the box 1 and the box 3, we attempt the derivation of the deterministic drift term and the stochastic noise defining the effective Langevin equation for the THC strength:

$$dq = F_{\text{eff}}(q, \varepsilon)dt + \varepsilon_{\text{eff}}(\varepsilon)dW \quad (13)$$

where our notation accommodates for a noise-dependent effective drift term, which corresponds to an efficient potential $V_{\text{eff}}(q, \varepsilon)$, such that $F_{\text{eff}}(q, \varepsilon) = -dV_{\text{eff}}(q, \varepsilon)/dq$. The pdfs $\pi(q)$ feature a very strong dependence on the intensity of the noise, with, as expected, higher noise intensity associated to flatter distributions (Fig. 1). We then derive the normalised potential $U(q, \varepsilon) = -\ln(\pi(q, \varepsilon))$ (Fig. 2). By matching the observed hopping rate $r(+ \rightarrow -)$ (Fig. 3a) with the right hand side of the formula given in Eq. (7) – in Fig. 3b we present the values of the factor $r(+ \rightarrow -)$ – we derive for each value of ε the corresponding value of ε_{eff} . Note that for each value of the noise we use only the observed difference between the value of $U(q, \varepsilon)$ evaluated in $q = 0$ and in $q = |q_{\text{ref}}|$ and the value of $|q_{\text{ref}}|$, and the upper bound of ε has been chosen so that $U(q_0, \varepsilon) - U(q_+, \varepsilon) = U(q_0, \varepsilon) - U(q_-, \varepsilon) \gtrsim 1.5$. As shown in Fig. 3c, we obtain that up to a high degree of precision $\varepsilon_{\text{eff}} \approx \gamma\sqrt{2}k\beta\varepsilon = \gamma\varepsilon$, where $\gamma \approx 1$ for all values of ε . Moreover, also in agreement with our expectations given at the end of Sect. 2, we have that $V_{\text{eff}}(q, \varepsilon) \approx V_{\text{eff}}(q)$, so that the effect of adding noise does not impact the deterministic drift term, or, in other terms, a deterministic dynamics is well defined. Figure 4 shows that for all values of noise the obtained effective potentials collapse into a single universal function, apart from an additive constant of no physical significance.

Our experimental procedure has shown quite convincingly that we can reduce the time evolution of the THC strength to a one-dimensional Langevin equation. We wish now to investigate how to derive analytically the drift and the noise term in Eq. (13) and an expression for the hopping rate

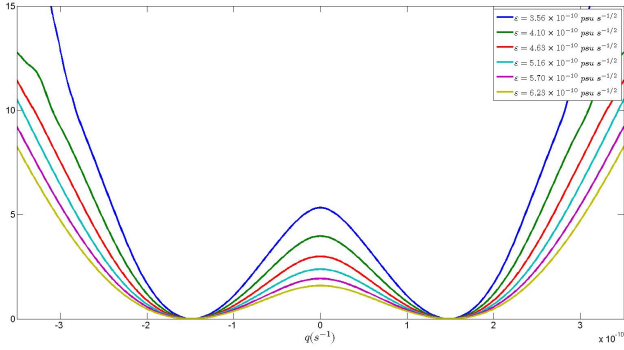


Fig. 2. One-dimensional adimensional potential $U(q)$ obtained as minus the logarithm of the pdf given in Fig. 1.

$r(+ \rightarrow -)$. Rewriting the system (10a–b) with respect to the new variables $q = k\beta(S_1 - S_3) = k(\rho_1 - \rho_3)$ and $Q = k(\rho_1 + \rho_3)$, we obtain the following coupled evolution equations:

$$dq = q(2k\tilde{\rho} - 3/2q^2/|q| - Q)dt + \varepsilon_q dW_q \quad (14a)$$

$$dQ = q(2k\tilde{\rho} + 1/2q^2/|q| - Q)dt - k\beta(F_1 + F_3)dt + \varepsilon_Q dW_Q \quad (14b)$$

where $\tilde{\rho} = \rho_0(1 - \alpha(\tilde{T}_1 + 2\tilde{T}_2 + \tilde{T}_3)/4 + \beta S_0)$ is the average density of the system, we have that, following Eq. (7), $\varepsilon_q = \varepsilon_Q = k\beta\sqrt{\varepsilon_1^2 + \varepsilon_3^2}$, and dW_q, dW_Q are increments of Wiener processes. Note that the drift terms in both Eq. (11) is odd with respect to the $q \rightarrow -q$ transformation and the SDE for the THC strength is in the form of Eq. (9), with $\varepsilon_q - \bar{\varepsilon}$. We assume that the system spends most of its time near the two deterministic equilibria with $q = \mp|q_{\text{ref}}|$, and that over the timescales of our interest the deterministic drift term of the variable Q vanishes. Assuming that the random forcing on Q has little impact on q , we derive the following approximate Langevin equation for the evolution of the THC strength:

$$dq \approx -2q(q^2/|q| - |q_{\text{ref}}|)dt + \varepsilon_q dW_q, \quad (15)$$

where, the drift term is odd with respect to parity and is independent of the noise strength. The corresponding effective potential $V_{\text{eff}}(q)$ is independent of ε and can be written as $V_{\text{eff}}(q) = -q^2|q_{\text{ref}}| + 2/3q^2|q| + \text{const}$. Note that this is a not a quartic symmetric potential but has the same parity properties and is twice differentiable everywhere. As shown in Fig. 4, this functional form closely approximates the experimental findings previously described, with discrepancies where the probability density is exponentially vanishing and small deviations also for $q \approx 0$ (where the density is also low). Moreover, if $\varepsilon_1 = \varepsilon_3$, we obtain that $\varepsilon_q = \bar{\varepsilon} = \sqrt{2}k\beta\varepsilon_1 \approx \varepsilon_{\text{eff}}$, so that the agreement between our experimental and theoretical findings for both the deterministic and stochastic part of the dynamics is quite satisfactory. This suggests that in the experimental setting of Eq. (12a–b)

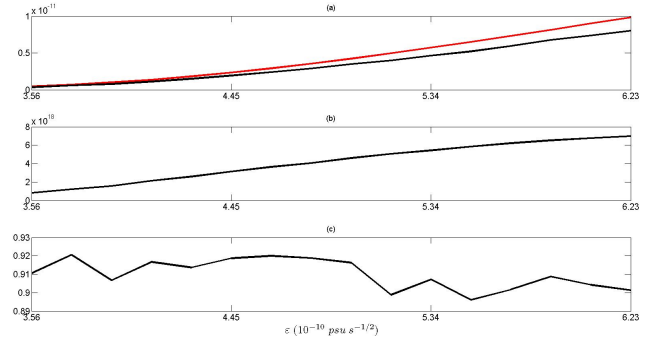


Fig. 3. Goodness of the one-dimensional approach for the reduced 2 d.o.f. model. **(a)** Average rate of hopping between the northern and southern sinking equilibria $r(+ \rightarrow -) = r(- \rightarrow +)$ (black line). The theoretical value is shown with the red line. Data are in units of s^{-1} . **(b)** Geometrical factor $G(+ \rightarrow -) = G(- \rightarrow +)$ of the hopping rate computed from Eq. (7). Adimensional quantity. **(c)** Value of the parameter γ giving the ratio between ε_{eff} (obtained as **a** divided by **b**) and the theoretically derived expression $\sqrt{2}k\beta\varepsilon$. Adimensional quantity.

it is possible to project very efficiently the dynamics of the system on the variable q alone, which seems to capture well the slow manifold (Saltzman, 2002) of the system. Using Eq. (15), we obtain the following approximate expression for the average rate of transition $r(+ \rightarrow -)$ between the northern sinking and the southern sinking state,

$$\begin{aligned} r(+ \rightarrow -) &\approx \frac{2\sqrt{2}}{3\pi} |q_{\text{ref}}| e^{-\frac{|q_{\text{ref}}|^2}{a(k\beta\varepsilon)^2}} \\ &= \frac{2\sqrt{2}}{3\pi} \sqrt{k\beta F_1} e^{-\frac{(F_1)^{3/2}}{a\sqrt{k\beta\varepsilon^2}}} = r(- \rightarrow +), \end{aligned} \quad (16)$$

where the last identity is due to the symmetry of the potential. This formula provides rates in excellent agreement with the outputs of the numerical simulations, as can be seen by comparing the red and the black line in Fig. 3a.

4.2 Asymmetric forcing to the simplified model

The obtained results suggest that the simplified model of the THC with only 2 degrees of freedom allows for a robust treatment of the one-dimensional stochastic dynamics of the THC strength. Nonetheless, in the previous set of experiments we have only verified the first condition for the robustness of the one-dimensional representation (well-posedness for linear scaling on the forcings). In this section, we wish to test how the system behaves when, following Eq. (9), we change the noise matrix ε_{ij} in such a way that the term noise strength $\bar{\varepsilon}$ for the considered observable is not altered. Therefore, we perform a new set of experiments, where the stochastic forcing is exerted only in one of the two boxes, e.g. box 1, so that in Eq. (12a–b) we set $\varepsilon_3 = 0$ and $\varepsilon_1 = \sqrt{2}\varepsilon$ for each value of ε considered in the previous set of experiments. This choice

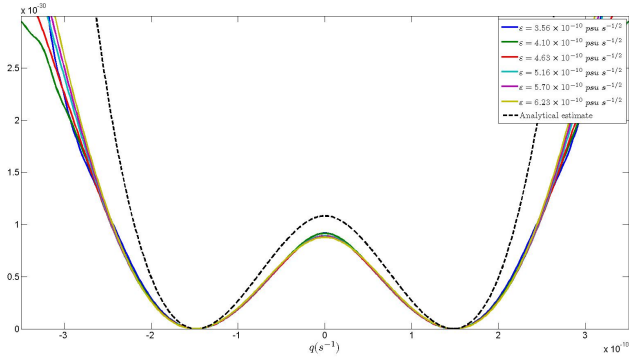


Fig. 4. One-dimensional efficient potential $V_{\text{eff}}(q)$ controlling the evolution of the THC strength in the 2 d.o.f. model with symmetric noise. The results of some numerical experiments are shown together with the analytical estimate.

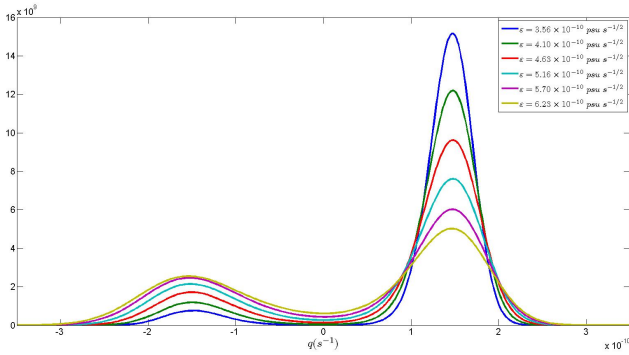


Fig. 5. Empirical probability distribution function for the THC strength in the 2 d.o.f. model with noise applied only in box 1.

guarantees that we have exact correspondence for the values of $\varepsilon_q = \bar{\varepsilon} = k\beta\sqrt{\varepsilon_1^2 + \varepsilon_3^2} = \sqrt{2}k\beta\varepsilon$, so that Eq. (9) looks exactly the same as in the previous set of experiments. Unfortunately, as shown in Fig. 5, following the procedure described in Sect. 2 we obtain for all values of ε an asymmetric probability distribution function, with the northern sinking equilibrium being the most probable state. The prominence of the $q > 0$ conditions become stronger as we consider weaker intensities for the noise. Therefore, the proposed stochastic modelling is not as robust as one could have guessed.

At a second thought the presence of asymmetry in this case becomes clearer. In this case the two sinking states undergo different forcing, because when $q > 0$ the stochastic forcing is exerted only in the box where downwelling occurs, whereas when $q \leq 0$ the stochastic forcing impacts only the box where upwelling occurs. Since, as explained at the beginning of Sect. 3 and discussed in Lucarini and Stone (2005a), the impact of changing freshwater fluxes is different in terms of destabilising the system depending on whether the forcing is applied in the box where downwelling or upwelling occurs, the two states are not equivalent

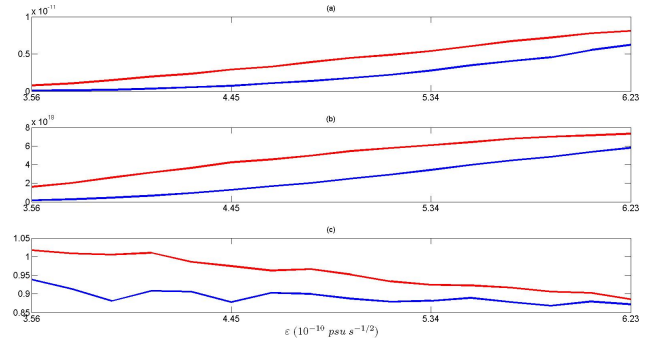


Fig. 6. Goodness of the one-dimensional approach for the non-symmetric case. **(a)** Average hopping rates $r(+ \rightarrow -)$ (blue line) and $r(- \rightarrow +)$ (red line). Data are in units of s^{-1} . **(b)** Geometrical factors $G(+ \rightarrow -)$ (blue line) and $G(- \rightarrow +)$ (red line) computed from Eq. (7). Adimensional quantities. **(c)** Value of the parameter $r\gamma$ giving the ratio between ε_{eff} (obtained as **a** divided by **b**) and the theoretically derived expression $\sqrt{2}k\beta\varepsilon$. Adimensional quantities. The values obtained using the $+ \rightarrow -$ (blue) and $- \rightarrow +$ (red) processes are not compatible.

anymore. In the previous set of experiments, as opposed to that, both boxes were equally (in a statistical sense) stochastically forced, and so that the northern and southern sinking states had equivalent forcings at all times. The more formal mathematical reason why the statistical properties of q are different in the two sets of experiments even if Eq. (9) is apparently the same can be traced to the differences in the correlation between the stochastic forcings to q and Q – compare Eq. (14a–b). In the case of symmetric stochastic forcing in the freshwater fluxes into the two boxes with $\varepsilon_1 = \varepsilon_3$, it is easy to see that the increments to the Wiener processes dW_q and dW_Q are not correlated, whereas when $\varepsilon_3 = 0$ the two quantities dW_q and dW_Q are identical so that their correlation is unitary.

We wish now to test whether, in such an asymmetric setting of forcings, the pdfs of the THC strength scale with the intensity of the noise in such a way to allow the possibility of defining consistently an effective potential $V_{\text{eff}}(q, \varepsilon)$ driving the deterministic part of the one-dimensional stochastic evolution of the THC strength. Given the asymmetry of the pdfs, such an effective potential would, unavoidably, be different from the one derived in the previous set of experiments, so that in no way we can be satisfied in terms of robustness of the one-dimensional Langevin approach. But, if we are able to define such an effective potential, we can deduce that each choice of the correlation matrix for the noise in the full system determines a specific projected effective deterministic dynamics, which is a weaker, but maybe still useful, result.

We follow the procedure described in Sect. 2, and for all the chosen values of ε we have that $U(q_0, \varepsilon) - U(q_-, \varepsilon) \gtrsim 1.5$ and $U(q_0, \varepsilon) - U(q_+, \varepsilon) \gtrsim 1.5$. In Fig. 6a we present the hopping rates $r(+ \rightarrow -)$ (blue) and $r(- \rightarrow +)$ (red). We see that both values increase monotonically with ε , as stronger

noise favours transitions. Note that, quite unexpectedly, for $\varepsilon > 7 \times 10^{-10} \text{ psu s}^{-1/2}$ (beyond the range where our full analysis is performed, not shown), $r(+ \rightarrow -)$ becomes bigger. By simple population algebra, this implies that the fraction of states with $q < 0$ is larger than 1/2, even if the most probable state is given, in all cases, by $q = q_{\text{ref}} > 0$. In Fig. 6b we plot the factors $G(+ \rightarrow -)$ and $G(- \rightarrow +)$. By matching the rate of transition with the corresponding G factor, one obtains for each value of ε the effective noise intensity ε_{eff} such that the probability distribution of the states and the transitions statistics are compatible. As mentioned before, in order to have a consistent picture, we need to obtain the same value of ε_{eff} by either using the $+ \rightarrow -$ or the $- \rightarrow +$ path. In Fig. 6c (compare with Fig. 3c) we show that ε_{eff} does not scale linearly with ε (or, equivalently, γ is not a constant) as found in the symmetric case, and, much more seriously, that there is no consistency between its value as obtained using the statistics of the $+ \rightarrow -$ and $- \rightarrow +$ transitions. Given the mismatch between the time-scale of the transitions and the populations in the two basins, we cannot reconstruct a well-defined effective potential $V_{\text{eff}}(q, \varepsilon)$, so that a consistent one-dimensional Langevin representation of the dynamics and statistics of the THC strength as proposed in Eq. (13) is not possible here. This can be the case if noise can effectively activate non-trivial dynamical processes allowing for a transition between the neighbourhoods of the two steady states with $q = \pm|q_{\text{eff}}|$, where by non-trivial we mean that they cannot be represented even approximately as a function of q only.

4.3 Symmetric forcing to the full system

We now revert to the full system described at the beginning of Sect. 3 and consider the case of symmetric boundary conditions. By adding stochastic perturbations to the freshwater fluxes in a similar fashion as in Eq. (11a–b), so that $F_j \rightarrow F_j + \varepsilon_j dW_j/dt$ with $j = 1, 3$ in Eq. (10a–f) we obtain the following Langevin equations for the variables $q = k(\rho_1 - \rho_3)$ and $Q = k(\rho_1 + \rho_3)$:

$$dq = q(2\tilde{q} - 3/2q^2/|q| - Q)dt + k\alpha\lambda(T_1 - T_3)dt + \varepsilon_q dW_q, \quad (17a)$$

$$dQ = q(2k\tilde{\rho} + 1/2q^2/|q| - Q)dt - \beta(F_1 + F_3)dt + k\alpha\lambda(\tilde{T}_1 - T_1 + \tilde{T}_3 - T_3)dt + \varepsilon_Q dW_Q, \quad (17b)$$

where the same notation as in Eq. (14a–b) has been used. Note that, as opposed to Eq. (11), the deterministic drift term is not odd with respect to the $q \rightarrow -q$ transformation, since in this case explicit temperature dependent terms are present, so that a negative parity is realised only when the signs of both $T_1 - T_3$ and $S_1 - S_3$ are changed. Note that, following the same derivation as in the case of the system with 2 d.o.f. and assuming that over time-scales of interest the drift

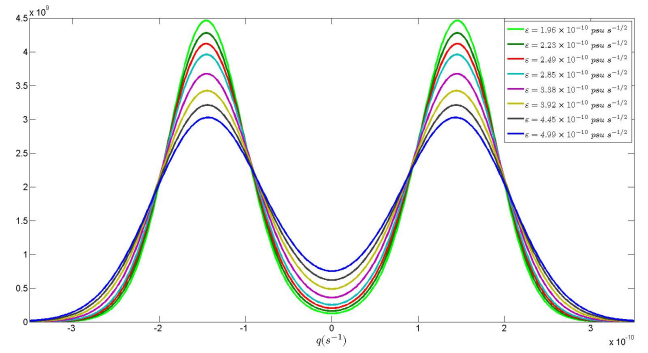


Fig. 7. Empirical probability distribution function for the THC strength in the full model for some selected values of symmetrically applied noise.

of Q is negligible, we end up writing the same approximate autonomous Eq. (15) for the THC strength under the hypothesis that the system spends most time near the two deterministic fixed points $q = \pm|q_{\text{eq}}|$. This suggests that also in this case we might empirically derive a well-defined effective potential $V_{\text{eff}}(q)$ analogous to the one obtained for the 2 d.o.f. model if considering stochastic forcing acting on both boxes 1 and 3.

Therefore, we follow the analysis of the previous subsection, and concentrate to stochastic perturbations to the freshwater flux having the same strength $\varepsilon_1 = \varepsilon_3 = \varepsilon$ in both hemispheres. We first observe that, as anticipated, for a given value of ε the distribution of the THC strength is flatter than in the case of the 2 d.o.f. model (Fig. 7). In order to obtain a pdf analogous to what obtained in the 2 d.o.f. case, in the full model we need to consider a stochastic forcing smaller by about 25 %, which suggests that the full model is, in some sense, less stable. We will discuss this below. As before, we select values of ε such that $U(q_0, \varepsilon) - U(q_-, \varepsilon) = U(q_0, \varepsilon) - U(q_-, \varepsilon) \gtrsim 1.5$ in order to be able to use Kramers' formula as a constraint to check the consistency of our data with the one-dimensional Langevin model. In Fig. 8a we report the hopping rates as a function of the intensity of the noise. The behaviour is qualitatively analogous to what shown in Fig. 3a for the 2 d.o.f. model, but, in agreement with the discussion above, the hopping rate between the $q > 0$ and the $q < 0$ states are consistently higher for the full model when the same stochastic forcing is considered. In Fig. 8b we present the factor $G(+ \rightarrow -) = G(- \rightarrow +)$ introduced in Eq. (7), which depends uniquely on the ratio between the probability density at the two maxima and at the local minimum for $q = 0$. Figure 8c, similarly to Fig. 3c, shows the proportionality constant γ between the value of ε_{eff} and the value of $\tilde{\varepsilon} = \sqrt{2}k\beta\varepsilon$. The parameter γ should be close to unity in the case the projection of the dynamics on q is “trivial” (as in the case of the 2 d.o.f. model with symmetric stochastic forcing), and, more importantly, γ should be approximately independent of ε . In the case analysed here both

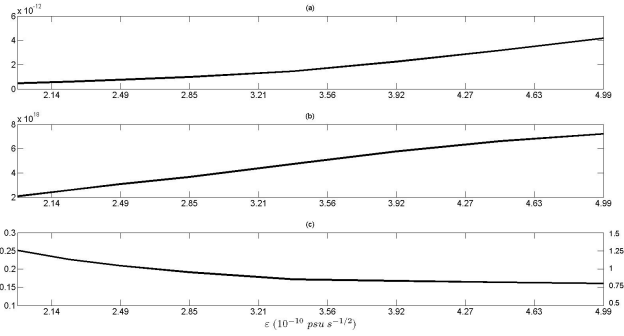


Fig. 8. Goodness of the one-dimensional approach for the full model. **(a)** Average rate of hopping between the northern and southern sinking equilibria $r(+ \rightarrow -) = r(- \rightarrow +)$ (black line). The theoretical value is shown with the red line. Data are in units of s^{-1} . **(b)** Geometrical factor $G(+ \rightarrow -) = G(- \rightarrow +)$ of the hopping rate computed from Eq. (7). Adimensional quantity. **(c)** Value of the parameter γ giving the ratio between ε_{eff} (obtained as **a** divided by **b**) and the theoretically derived expression $\sqrt{2}k\beta\varepsilon$. The value of γ obtained if considering the renormalized value (referred to the 2 d.o.f. model) for k can be read on the right hand side scale. Adimensional quantities.

conditions are not satisfied. The first issue points to the fact that we need to renormalize the constants when developing a lower dimensional projected dynamics (which is exactly what we did when constructing the 2 d.o.f. model from the full model). In fact, if we consider as effective k the one considered for the 2 d.o.f. model, the values of γ are relatively close to 1 (check the scale on the right hand side of Fig. 8c). More critical is the presence of a nonlinear relation between ε_{eff} and ε , which is caused by the fact that in the full model faster modes are excited by noise and impact in a nontrivial way the effective surrogate noise acting on the q variable taken as independent.

When reconstructing from $U(q, \varepsilon)$ the actual effective potential $V_{\text{eff}}(q, \varepsilon)$ using Eq. (5) we obtain that the effective potential is a function of q only, so that $V_{\text{eff}}(q, \varepsilon) = V_{\text{eff}}(q)$ (see Fig. 9) so that it is possible to disentangle completely the drift term from the stochastic forcing. Moreover, such potential is very similar to what obtained in the reduced model with 2 d.o.f., as can be seen by comparing Figs. 4 and 9. The height of the potential barrier between the two minima is slightly lower in the case of the full model analysed here, in agreement with the argument of the destabilising feedback due to the thermal restoring process presented by Lucarini and Stone (2005a). This can be explained as follows: since perturbations in the value of $q - |q_{\text{eq}}|$ are positively correlated to perturbations $T_1 - T_3$ thanks to advection, the contribution $k\alpha\lambda(T_1 - T_3)$ in Eq. (17a) weakens the force driving the system towards the nearby deterministic fixed point, thus enhancing its instability. Overall, the good agreement between Figs. 4 and 9 implies that the deterministic dynamics of the THC strength is robustly consistent between the full and reduced model.

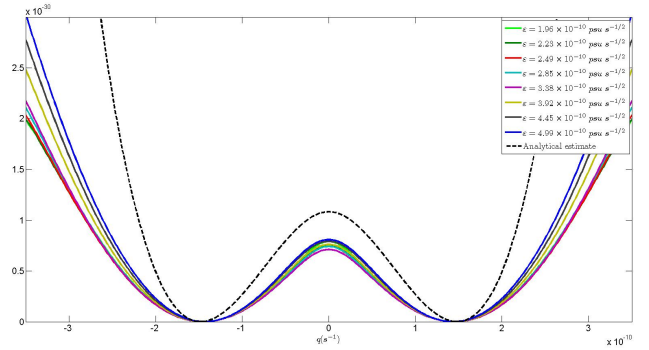


Fig. 9. One-dimensional efficient potential $V_{\text{eff}}(q)$ controlling the evolution of the THC strength in the 5 d.o.f. model with symmetric noise. The results of some numerical experiments are shown together with the analytical estimate. The results are rather similar to what shown in Fig. 4.

5 Numerical experiments: stochastic resonance

Stochastic resonance (Benzi et al., 1982; Nicolis 1982; Gammaitoni et al., 1998) is an exceedingly interesting process whereby noise amplifies the response of the system at the same frequency of a periodic forcing. Typically, it is realised when we consider a Langevin equation of the form:

$$dx = F(x)dt + A \sin(\omega t + \phi)dt + \varepsilon dW \quad (18)$$

where the drift term $F(x) = -dV(x)/dx$ derives from a (symmetric, but necessarily so) potential $V(x)$ with a double well-structure like those considered in this study. We can associate the time dependent drift $G(x, t) = F(x) + A \sin(\omega t + \phi)$ to a time dependent potential $W(x, t) = -\int dx F(x) - A \sin(\omega t + \phi)x = V(x) - Ax \sin(\omega t + \phi)$. The periodic forcing modulates the bistable system, so that one stable state corresponding to one of the two minima results to be less stable than the other every half a period of the forcing, so that every period the populations of the neighbourhoods of these points are, alternatively, exponentially increased and suppressed by a factor $\exp(\pm 2A(|x_- - x_0|)/\varepsilon^2)$. As accurately discussed in Gammaitoni et al. (1998), when we tune the noise intensity ε so that the inverse of the hopping rate given in Eq. (7) for the unperturbed system is approximately equal to half of the period π/ω , the system is in a condition where there is maximum probability for leaving the less stable state into the more stable one, before, randomly, the system switches back. In this case, the system attains a high degree of synchronisation with the input periodic signal, so that the output is basically a square wave with constant phase difference with the sinusoidal forcing.

Interestingly, even if the stochastic resonance is apparently a highly nonlinear process, it can be accurately described using linear response theory, whereby one studies the amplitude of the output signal at the same frequency of the periodic forcing for various values of the noise. Such amplitude,

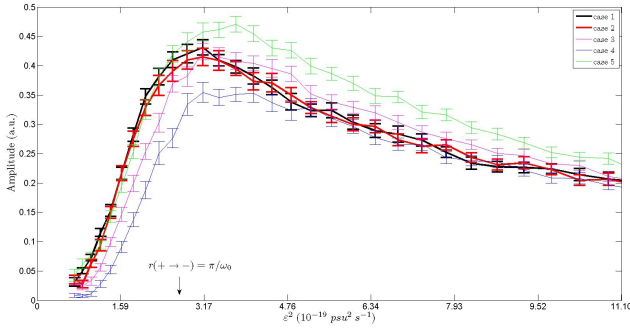


Fig. 10. Response to linear periodic perturbation as a function of the background noise for five experimental settings. Case 1 and 2 are indicative of stochastic resonance. The value of noise realising the approximate matching condition $r(+ \rightarrow -) = r(- \rightarrow +) = \pi/\omega_0$ is indicated.

in the weak field limit, is proportional to the amplitude of the periodic forcing A (Gammaitoni et al., 1998).

Along these lines, we consider the 2 d.o.f. model, and take into account a periodic modulation to the freshwater fluxes, so that $F_1 \rightarrow F_2 + \phi \Delta F \sin(\omega t)$ and $F_3 \rightarrow F_3 + \psi \Delta F \sin(\omega t)$, with symmetric background forcing $F_1 = F_3$. Assuming that the stochastic forcing acts with equal strength $\varepsilon_1 = \varepsilon_3 = \varepsilon$ on both boxes 1 and 3, as analysed in Sect. 4.1, we obtain that Eq. (15), which satisfactorily describes the one-dimensional stochastic dynamics of q , is modified as follows:

$$dq \approx -2q(q^2/|q| - |q_{\text{ref}}|)dt + k\beta(\phi - \psi)\Delta F \sin(\omega t) + \varepsilon_q dW_q, \quad (19)$$

which is exactly in the form of Eq. (15). As we see, for a given value of ΔF , the strength of the periodic forcing to q depends only on the absolute value of the difference $(\phi - \psi)$, and not separately on the values of ϕ and ψ . If the dynamics of q is accurately described by a one-dimensional Langevin equation, we expect to be able to observe the process of stochastic resonance when ω and ε_q are suitably matched. In order to test this, we set the period $\omega = \omega_0 = 2\pi/19\,000$ yr – where the period of 19 000 yr has been chosen because it is long compared to the internal time scale $|q_{\text{ref}}|^{-1} \approx 215$ yr and has paleoclimatic relevance in conjunction to Milankovitch theory (Velez-Bechi et al., 2001; Saltzman, 2002) – choose a moderate value for the amplitude of the sinusoidal modulation $\Delta F = 9 \times 10^{-12} \text{ s}^{-2}$ and study the amplitude of the ω_0 frequency component of the times series of q as a function of ε , and create an ensemble of 100 members for each considered setting. We can state that the phenomenology of stochastic resonance is well reproduced if (a) we find the characteristic peak for the response in the vicinity of a value of ε such that the corresponding hopping rate for the unperturbed system given in Fig. 3a is close to π/ω_0 , and (b) such response depends, for all values of ε , on $(\phi - \psi)$ only. We refer to the scenario where $\phi = -\psi = 1/2$ as case 1, and the scenario where $\phi = 1, \psi = 0$ as case 2. Note that

the case $\phi = 0, \psi = -1$ is identical to case 2 by symmetry. The results are reported in Fig. 10, with the black line corresponding to case 1 and the red line corresponding to case 2. The obtained curves for the amplitude of the response agree very accurately, especially considering the rather small uncertainty, and feature exactly the right shape as presented in Gammaitoni et al. (1998). We observe a relatively broad resonance for values of noise comparable to those inducing in the unperturbed system transitions with average rate similar to the semiperiod of the forcing. Finding quite accurately the signature of stochastic resonance is a further proof that in the special setting of the unperturbed system considered here the dynamics of q is indeed quasi-one dimensional.

We want to contrast this positive outcome with what one obtains by adding periodic perturbations of the same form as above to an “unperturbed” state featuring stochastic forcing acting on box 1 only, described in Sect. 4.2. We choose exactly the same forcing parameters as above and repeat the experiments using the same ensemble size. We refer to the scenario where $\phi = -\psi = 1/2$ as case 3, and the scenario where $\phi = 1, \psi = 0$ as case 4, and to the scenario where $\phi = 0, \psi = -1$ as case 5. Note that here case 4 and case 5 are not equivalent. We obtain (see Fig. 10) that in the three cases, whereas we obtain qualitatively and quantitatively analogous results for the normalised amplitude of the response of the output at the same frequency of the forcing, the curves are distinct with high statistical significance and differ also from what obtained for cases 1 and 2. Such discrepancy would not be possible if the dynamics of q were accurately described with a one-dimensional effective potential plus stochastic noise plus periodic forcing. The fact that the three curves 3, 4, and 5 are not superimposed (and disagree with 1 and 2) further supports the fact that the dynamics of q is not quasi-one dimensional. Nonetheless, an obvious process of resonance (not strictly one-dimensional) is obviously still in place.

6 Conclusions

In this work we have re-examined the classic problem of trying to reconstruct the effective stochastic dynamics of an observable from its time series in the special case of a clearly bimodal empirical probability density function. This issue is especially relevant in climatic and paleoclimatic research, where it is very tempting to try to deduce the large scale, qualitative properties of the climate system, their modulation with time, the potential presence of tipping points through the observations of long time series of proxy data (Livina and Lenton, 2007; Livina et al., 2009). Furthermore, since amplifying mechanisms such as stochastic resonance have been proposed to explain enhanced low frequency variability of the oceanic circulation as a result of slow modulations of some parameters (Ganopolski and Rahmstorf, 2002), such empirically reconstructed statistical/dynamical

properties may be interpreted as starting points to deduce special “sensitivities” of the climate system. Therefore, an important question is to understand how accurate and robust these procedures of reconstruction are.

From this work it is apparent that from the statistical properties of the time series of an observable featuring a symmetric bimodal pdf it seems relatively easy to construct the corresponding one-dimensional Langevin equation by determining the drift term and the intensity of the white noise by basically imposing a self-consistent population dynamics. Assuming that the observable is a function of the phase space variables of a stochastic dynamical systems, it is an obvious temptation to interpret the obtained equation as the description of the projected dynamics for the observable, where the impact of the other (in general, many) neglected degrees of freedom of the system contributes to defining the effective deterministic dynamics and to creating a surrogate white noise term. Nonetheless, if the pdf of the observable is not symmetric, the possibility of constructing a meaningful surrogate stochastic dynamics relies on the fact that one should be able to describe consistently the hopping process and between the two attraction basins and their steady state populations.

In this work we have considered two very simple box models of the oceanic circulation (Rooth, 1982; Scott et al., 1999; Lucarini and Stone, 2005a, b), comprising two high latitude and a low latitude boxes with time-dependent temperature and salinity as testbeds for these methodologies. These models are able to reproduce the bistability properties of the thermohaline circulation, by featuring two possible asymmetric circulations (one mirror image of the other) in presence of symmetric external heat and freshwater forcings. In both models the circulation strength is parameterised as proportional to the difference between the densities of the two high-latitude boxes. The simpler 2 d.o.f. model is suitably derived from the full, 5 d.o.f. model by imposing a fixed temperature for the boxes.

We first impose stochastic forcing of the same intensity on the freshwater forcings to the two high-latitude boxes and observe that the resulting pdf of the thermohaline circulation strength is bimodal and symmetric. More importantly, for both models the dynamics of q can be accurately described with a Langevin equation with a drift term derived from a one-dimensional effective potential plus stochastic noise. An excellent approximation to the true dynamics (as well as to the hopping rates) can be obtained in an explicit form by imposing that the sum of the densities of the two high-latitude boxes is a slow variable. The main difference between the two is that in the full model the nonlinear feedbacks acting on the variable we are neglecting alter in a nontrivial, nonlinear way the effective surrogate noise acting on the q variable. In other terms, in the case of the full model a careful tuning of the noise allows for taking care very accurately – in a statistical sense – of the effect of all the variables we are neglecting. Our results are obtained for a specific value

of the background freshwater forcing $F_1 = F_3$, but the following scaling allow to extensively generalise our findings: $q \sim (F_1)^{1/2}$, $V_{\text{eff}}(q) \sim (F_1)^{3/2}$, $\varepsilon_{\text{eff}} \sim (F_1)^{3/4}$. Things change dramatically when considering the case of stochastic forcing acting only on one of the two high-latitude boxes. We tune our experiments in such a way that, apparently, the Langevin equation for the q variable is not altered with respect to the previous case. Not only we obtain a non-symmetric pdf, but, moreover, it is not possible to reconstruct an approximate but consistent stochastic dynamics for the q variable alone. Therefore, there is no ground for achieving a satisfactory projection of the dynamics, and a one-dimensional Langevin equation cannot be derived.

Finally, we test the possibility of observing the mechanism of stochastic resonance in the simplified 2 d.o.f. model by superimposing a slow periodic modulation on the freshwater fluxes in the two high-latitude boxes to the acting stochastic forcings. Whereas in the scenario where the noise acts with equal strength in both boxes we obtain numerically outputs in close agreement with the theory stochastic resonance for one dimensional systems, thus supporting the idea that a projected dynamics is indeed a good approximation when attempting a description of the properties of q , the opposite holds in the scenario where the noise acts only on one box. This further supports the fact that in this case the dynamics of q is not trivially quasi-one dimensional, and transitions occur through processes that cannot be written precisely as a function of q only.

Our results support the idea that deducing the approximate stochastic dynamics for an observable of a multidimensional dynamical system from its time series is definitely a non-trivial operation. The reconstructed drift term and the noise forcing depend, in general, in a non-trivial way on the intensity and correlation properties of the white noise of the true system. This implies that a “true” projected dynamics cannot be defined. Therefore, in practical applications, it seems tentative to assume that from the pdf of a bimodal observable obtained with a given level of noise it is possible to understand how the bistability property of the full system will change when the level of the input noise is altered. In particular, it seems difficult to be confident on obtaining information on how the rate of transition between the two equilibria will change and on the characteristics of tipping points. This also suggests that recent claims of the possibility of detecting robustly early warning signals for critical transitions at tipping points from time series as proposed in Scheffer (2008) must be carefully checked.

A better understanding of the properties of multistable models can be reached only by going beyond a simplified description of the statistical properties of the observables we are mostly interested into. In order to address these points, we will attempt the kind of critical analysis proposed in this paper on more complex models but still idealised models of the thermohaline circulation, such as that considered in Lucarini et al. (2005, 2007). We will test to what extent a

simple one-dimensional Langevin model for the evolution of the THC strength can be defined and in which range of the parameters determining the boundary conditions, and how the drift and noise terms of the Langevin model – if it can be defined – are related to the internal parameters of the full system. At a more theoretical level, we will try to propose approximate ways for dealing with the transitions between high occupations regimes in multidimensional gradient flows, attempting to derive a markovian description of transitions between discrete states.

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References

- Benzi, R., Parisi G., Sutera A., and Vulpiani A.: Stochastic resonance in climatic change, *Tellus* 34, 10–16, 1982.
- Benzi, R.: Stochastic resonance: from climate to biology, *Nonlin. Processes Geophys.*, 17, 431–441, doi:10.5194/npg-17-431-2010, 2010.
- Boyle, E. and Kegwin, L.: North Atlantic thermohaline circulation during the past 20,000 years linked to high-latitude surface temperature, *Nature*, 330, 35–40, 1987.
- Cessi, P.: A simple box-model of Stochastically forced Thermohaline flow, *J. Phys. Ocean.*, 24, 1911–1920, 1994.
- Cessi, P. and Young, W. R.: Multiple equilibria in two-dimensional thermohaline circulation, *J. Fluid Mech.*, 241, 291–309, 1992.
- Dijkstra, H.: *Nonlinear Physical Oceanography*, Springer, Berlin, 2005.
- Ditlevsen, P.: Observation of alpha-stable noise and a bistable climate potential in an ice-core record, *Geophys. Res. Lett.*, 26, 1441–1444, 1999.
- Freidlin, M. I. and Wentzell, A. D.: *Random perturbations of dynamical systems*, Springer, Berlin, 1998.
- Fraedrich, K.: Structural and stochastic analysis of a zero-dimensional climate system, *Q. J. R. Meteorol. Soc.*, 104, 461–474, 1978.
- Gammaitoni, L., Hänggi, P., Jung, P., and Marchesoni, F.: Stochastic Resonance, *Rev. Mod. Phys.*, 70, 223–287, 1998.
- Ganopolski, A. and Rahmstorf, S.: Abrupt Glacial Climate Change due to Stochastic Resonance, *Phys. Rev. Lett.*, 88, 038501, doi:10.1103/PhysRevLett.88.038501, 2002.
- Hasselmann, K.: Stochastic climate models, Part 1: Theory, *Tellus*, 28, 473–485, 1976.
- Hawkins, E., Smith, R. S., Allison, L. C., Gregory, J. M., Woollings, T. J., Pohlmann, H., and de Cuevas, B.: Bistability of the Atlantic overturning circulation in a global climate model and links to ocean freshwater transport, *Geophys. Res. Lett.*, 38, L10605, doi:10.1029/2011GL047208, 2011.
- Kuhlbrodt, T., Griesel, A., Montoya, M., Levermann, A., Hofmann, M., and Rahmstorf, S.: On the driving processes of the Atlantic meridional overturning circulation, *Rev. Geophys.*, 45, RG2001, doi:10.1029/2004RG000166, 2007.
- Lenton, T. M., Held, H., Kriegler, E., Hall, J. W., Lucht, W., Rahmstorf, S., and Schellnhuber, H. J.: Tipping elements in the Earth’s climate system, *Proc. Natl. Acad. Sci. USA*, 105, 1786–1793, 2008.
- Livina, V. N. and Lenton, T. M.: A modified method for detecting incipient bifurcations in a dynamical system, *Geophys. Res. Lett.*, 34, L03712, doi:10.1029/2006GL028672, 2007.
- Livina, V. N., Kwasniok, F., and Lenton, T. M.: Potential analysis reveals changing number of climate states during the last 60 kyr, *Clim. Past*, 6, 77–82, doi:10.5194/cp-6-77-2010, 2010.
- Lucarini, V. and Stone, P. H.: Thermohaline circulation stability: a box model study, Part I: uncoupled model, *J. Climate*, 18, 501–513, 2005a.
- Lucarini, V. and Stone, P. H.: Thermohaline circulation stability: a box model study, Part II: coupled atmosphere-ocean model, *J. Climate*, 18, 514–529, 2005b.
- Lucarini, V., Calmanti, S., and Artale, V.: Destabilization of the thermohaline circulation by transient perturbations to the hydrological cycle, *Clim. Dynam.*, 24, 253–262, 2005.
- Lucarini, V., Calmanti, S., and Artale, V.: Experimental mathematics: Dependence of the stability properties of a two-dimensional model of the Atlantic ocean circulation on the boundary conditions, *Russ. J. Math. Phys.*, 14, 224–231, 2007.
- Lucarini, V., Fraedrich, K., and Lunkeit, F.: Thermodynamic Analysis of snowball Earth hysteresis experiment: Efficiency, Entropy Production, and irreversibility, *Q. J. Royal Met. Soc.*, 136, 2–11, 2010.
- Mannella, R. and Palleschi, V.: Fast and precise algorithm for computer simulation of stochastic differential equations, *Phys. Rev. A*, 40, 3381–3386, 1989.
- Monaham A. H.: Stabilization of Climate Regimes by Noise in a Simple Model of the Thermohaline Circulation, *J. Phys. Ocean.*, 32, 2072–2085, 2002.
- Nicolis, C.: Stochastic aspects of climatic transitions-response to a periodic forcing, *Tellus*, 34, 1–9, 1982.
- Pierrehumbert, R. T., Abbot, D., Voight, A., and Koll, D.: Neoproterozoic Climate, *Ann. Rev. Earth. Plan. Sci.* 39, 417–460, 2011.
- Rahmstorf, S.: Bifurcation of the Atlantic thermohaline circulation in response to changes in the hydrological cycle, *Nature*, 378, 145–149, 1995.
- Rahmstorf, S.: Ocean Circulation and Climate during the last 120,000 years, *Nature*, 419, 207–214, 2002.
- Rahmstorf, S.: Timing of abrupt climate change: a precise clock, *Geophys. Res. Lett.*, 30, 1510, doi:10.1029/2003GL017115, 2003.
- Rahmstorf, S., Crucifix, M., Ganopolski, A., Goosse, H., Kamenkovich, I., Knutti, R., Lohmann, G., Marsh, R., Mysak, L. A., Wang, Z., and Weaver, A. J.: Thermohaline circulation hysteresis: A model intercomparison, *Geophys. Res. Lett.*, 32, L23605, doi:10.1029/2005GL023655, 2005.
- Rooth, C.: Hydrology and ocean circulation, *Progr. Oceanogr.*, 11, 131–149, 1982.

- Saltzman, B.: *Dynamical Paleoclimatology: Generalized Theory of Global Climate Change*, Academic Press, New York, 2002.
- Scheffer, M., Bascompte, J., Brock, W. A., Brovkin, V., Carpenter, S. R., Dakos, V., Held, H., van Nes, E. H., Rietkerk, M., and Sugihara, G.: Early warning signals for critical transitions, *Nature*, 461, 53–59, 2009.
- Scott, J. R., Marotzke, J., and Stone, P. H.: Interemispheric Thermohaline circulation in a coupled box-model, *J. Phys. Ocean.*, 29, 351–365, 1999.
- Scott, J. R., Sokolov, A. P., Stone, P. H., and Webster, M. D.: Relative roles of climate sensitivity and forcing in defining the ocean circulation response to climate change, *Clim. Dynam.*, 30, 441–454, 2008.
- Sijp, W. P. and England, M. H.: Sensitivity of the Atlantic thermohaline circulation to basin-scale variations in vertical mixing and its stability to fresh water perturbations, *J. Climate*, 19, 5467–5478, 2006.
- Sijp, W. P. and England, M. H.: Precise calculations of the existence of multiple AMOC equilibria in a coupled climate model Part II: transient behaviour, *J. Climate*, in press, doi:10.1175/2011JCLI4246.1, 2011.
- Sijp, W. P., England, M. H., and Gregory, J. M.: Precise calculations of the existence of multiple AMOC equilibria in coupled climate models Part I: equilibrium states, *J. Climate*, in press, doi:10.1175/2011JCLI4245.1, 2011.
- Stocker, T. F. and Schmittner, A.: Influence of CO₂ emission rates on the stability of the thermohaline circulation, *Nature*, 388, 862–865, 1997.
- Stocker, T. F. and Wright, D. G.: Rapid transitions of the ocean's deep circulation induced by changes in the surface water fluxes, *Nature*, 351, 729–732, 1991.
- Stouffer, R. J. and Manabe, S.: Equilibrium Response of Thermohaline Circulation to large Changes in Atmospheric CO₂ concentration, *Clim. Dynam.* 20, 759–777, 2003.
- Stommel, H.: Thermohaline convection with two stable regimes of flow, *Tellus*, 13, 224–230, 1961.
- Titz, S., Kuhlbrodt, T., and Feudel, U.: Homoclinic bifurcation in an ocean circulation box model, *Int. J. Bif. Chaos*, 12, 869–875, 2001.
- Titz, S., Kuhlbrodt, T., Rahmstorf, S., and Feudel, U.: On freshwater-dependent bifurcations in box models of the inter-hemispheric thermohaline circulation, *Tellus*, 54A, 89–98, 2002.
- Velez-Belchi, P., Alvarez, A., Colet, P., Tintor, J., and Haney, R.L.: Stochastic Resonance in the thermohaline circulation, *Geophys. Res. Lett.*, 28, 2053–2057, 2001.
- Vellinga, M.: Instability of two-dimensional thermohaline circulation, *J. Phys. Oceanogr.*, 26, 305–319, 1996.