

How well are Tropical Cyclones represented in reanalysis data sets?

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1 **How Well are Tropical Cyclones Represented in Reanalysis Data Sets?**

2 Kevin Hodges*

3 *Department of Meteorology, University of Reading, Reading, United Kingdom*

4 Alison Cobb

5 *Department of Meteorology, University of Reading, Reading, United Kingdom*

6 Pier Luigi Vidale

7 *National Centre for Atmospheric Science, Department of Meteorology, University of Reading,*
8 *Reading, United Kingdom*

9 **Corresponding author address: Department of Meteorology, University of Reading, Reading,*

10 *United Kingdom.*

11 **E-mail: k.i.hodges@reading.ac.uk**

ABSTRACT

12 Tropical cyclones (TCs) are identified and tracked in six recent reanalysis
13 data sets and compared with those from the IBTrACS best track archive. Re-
14 sults indicate that nearly every cyclone present in IBTrACS over the period
15 1979-2012 can be found in all six reanalyses using a tracking and matching
16 approach. However, TC intensities are significantly under-represented in the
17 reanalyses compared to the observations. Applying a typical objective TC
18 identification scheme, it is found that the largest uncertainties in TC identi-
19 fication occur for the weaker storms; this is exacerbated by uncertainties in
20 the observations for weak storms and lack of consistency in operational pro-
21 cedures. For example, it is unclear whether certain types of storms, such as
22 tropical depressions, subtropical cyclones and monsoon depressions, are in-
23 cluded in the best track data for all reporting agencies. There are definite im-
24 provements in how well TCs are represented in more recent, higher resolution
25 reanalyses; in particular MERRA2 is comparable with the NCEP-CFSR and
26 JRA55 reanalyses, which perform significantly better than the older MERRA
27 reanalysis.

28 **1. Introduction**

29 Tropical cyclones (TCs) are one of the most damaging weather-related natural hazards on the
30 planet, causing 42 % of the United States catastrophe-insured losses in the period 1992-2011
31 (King 2013). Individual intense events can result in severe losses. For example, Hurricane Katrina
32 resulted in an estimated death toll of 1,833 people and financial losses of over \$125 billion (Adeola
33 and Picou 2014). Weaker storms, such as tropical depressions can also have an impact in terms of
34 loss of life and disruption in vulnerable societies (for the Caribbean 2009). It is therefore important
35 to utilise the available data and new analysis techniques to better understand their properties and
36 behaviour, with the aim of mitigating their societal, economic and environmental impacts.

37 Due to the relatively short observational record of TCs, and problems with sampling within
38 the record, there is considerable uncertainty in the variability of TCs in terms of frequency, over
39 climate time scales of the last 100 years (Landsea 2007; Landsea et al. 2009), resulting in un-
40 certainty in the interannual variability and trend detection. The use of reanalyses to detect TCs
41 provides an opportunity to reduce this uncertainty (Truchelut et al. 2013), by allowing the creation
42 of a larger data sample which, when used in conjunction with the historic observational data, can
43 help to provide more confidence in TC numbers than the observations alone. Reanalyses combine
44 observations with a short forecast from a general circulation model (GCM) to produce gridded
45 data sets with regular output intervals constrained by the observations, and can act as a bridge
46 between the observations of TCs and simulated tempestology. However, there can be problems in
47 using reanalyses related to the changing observing system, in particular the introduction of spuri-
48 ous trends (Bengtsson et al. 2004a), and the fact that different reanalyses use different GCMs with
49 different parameterizations and different data assimilation methods, all of which can contribute
50 to differences between them. The study of Schenkel and Hart (2012) previously considered the

51 representation of TCs in the northern hemisphere in several reanalyses, including several of those
52 used in this study, by manually tracking the best track TCs in the reanalyses, and found consid-
53 erable variation in the properties of TCs between the reanalyses, for location, and a consistently
54 large underestimate of intensity (10m winds and Mean Sea Level Pressure) for all the reanalyses.
55 This uncertainty in the representation of TC properties in reanalyses can introduce uncertainty into
56 their detection in these data, so that detection criteria are often tailored to the particular reanalysis
57 of interest (Murakami 2014).

58 Another motivation for a careful study of the properties of TCs as represented by reanalyses
59 is that they are often used as a means of calibrating TC detection and tracking schemes before
60 applying them to climate models (Bengtsson et al. 2007a). This is done by first applying the
61 detection to the reanalyses or operational analyses and adjusting the detection criteria to give
62 similar numbers of TCs to those found in the observations, provided by the TC warning centers
63 best track data. This may be problematic if there are large differences between how reanalyses
64 represent TCs in terms of their properties, such as structure and intensities, or if there are biases
65 in the best track data. The model dynamical core, parameterizations and resolution all play a
66 critical role in determining the output of extreme events in reanalysis data. These vary widely,
67 with in particular newer generations of reanalyses being produced at higher resolutions and with
68 more modern data assimilation systems. For climate models, the IPCC 5th Assessment (2013)
69 stated that there is medium evidence and high agreement that year-to-year count variability of
70 Atlantic hurricanes can be well simulated by modest resolution (100 km or finer) atmospheric
71 GCMs (AGCMs) forced by observed Sea Surface Temperatures (SSTs). Both Strachan et al.
72 (2013) and Roberts et al. (2015) show that 60 km is adequate for simulating interannual variability,
73 although not intensity. Recent work by Murakami (2014) showed that, when considering five
74 reanalyses (also included in this study), the highest resolution reanalysis was not always the best

75 in terms of simulating the TC climatology and properties, nor did the higher-resolution reanalyses
76 produce significantly more intense storms than those with lower resolutions, suggesting that the
77 simulation of TCs in the reanalyses is highly dependent on model formulation (Schenkel and Hart
78 2012) and/or data assimilation strategy. However, if we can understand the uncertainties of TCs
79 in the reanalyses, they may provide a useful means of extending the observations, for example,
80 by extending their lifecycles to include the extratropical transition (Jones et al. 2003) and beyond,
81 which would be useful for TC related extratropical risk analysis and GCM assessment (Haarsma
82 et al. 2013). The use of reanalysis could also assist in the identification of subtropical and hybrid
83 tropical storms (Roth 2002; Guishard et al. 2009), which are also associated with severe weather,
84 providing a more complete set of tropical storm data for use in GCM assessment than is perhaps
85 currently present in best track data; the inclusion of these types of storms in the best track data sets
86 is highly variable between the operational centres.

87 The main aim of this paper is to quantify the uncertainties in how well TCs are represented in a
88 number of recent reanalyses, and how this affects the objective identification of TCs in reanalyses.
89 This is achieved by exploring:-

- 90 1. how well reanalyses represent the observed TCs in the best track data using direct track
91 matching.
- 92 2. how well does an objective identification scheme identify the best track TCs in the reanalyses
93 and what might be the cause of differences.

94 **2. Data and Methods**

95 Data from six recent reanalyses are used in this study and described below. Also used are best
96 track data produced by the tropical warning centers as post season analyses of the TC tracks. These

97 have been combined into the International Best Track Archive for Climate Stewardship (IBTrACS)
98 data set (Knapp et al. (2010)) and are used in this study for verifying the TCs identified in the
99 reanalyses. The IBTrACS-ALL, which includes data from all agencies, is used in this study. The
100 common period of 1979-2012 is used throughout for all data sets, except for one reanalysis where
101 the period is 1980-2012. Throughout the rest of the paper the following nomenclature is used; the
102 term Tropical Cyclone (TC) is used for warm core storms generally and, where appropriate, the
103 term Tropical Storm (TS) is used for TCs with wind speeds greater than 17 m s^{-1} .

104 *a. Best Track dataset*

105 For full details of the IBTrACS-ALL data set, see Knapp et al (2012). The original wind speed
106 data in knots is converted to wind speed in m s^{-1} . The World Meteorological Organization (WMO)
107 standard for reported tropical cyclone wind speed is maximum 10-minute sustained winds at 10
108 m height over a smooth surface; however, this is rarely observed, therefore some discrepancy be-
109 tween agencies is apparent. Different agencies apply different wind-averaging periods, with the
110 East Pacific, North Atlantic (RSMC Miami), and central Pacific (RSMC Honolulu) using 1-minute
111 averaging periods, North Indian (RSMC New Delhi) using a 3-minute period and the other agen-
112 cies using 10-minute averaging periods (Schreck III et al. 2014). The 10-minute wind speeds
113 are converted to 1-minute wind speeds using a factor of 1.13, which has traditionally been used
114 (Harper et al. 2010), and the data from RSMC Miami and New Delhi are used in their original
115 form. However, there are uncertainties in the accuracy and fidelity of this conversion, with differ-
116 ent conversion factors for at-sea, off-sea, off-land and in-land parts of the storm suggested (Harper
117 et al. 2010). Other uncertainties also exist in the best track data, which have been discussed in
118 several studies; a summary of these uncertainties can be found in the appendix of Hodges and

119 Emerton (2015). They include issues relating to location and intensity uncertainties and opera-
120 tional differences between agencies. This is further discussed in the Discussion section.

121 For the analysis of the identified TCs in different ocean basins the IBTrACS basin boundaries
122 (Knapp et al. 2010) have been used, with TCs assigned to a particular ocean basin, based on where
123 the storm reaches maximum wind speed intensity.

124 *b. Reanalysis datasets*

125 Meteorological centers around the world produce reanalysis data sets as an ongoing enterprise.
126 The reanalyses are essentially based on frozen operational numerical weather prediction (NWP)
127 systems. New reanalyses are often released following significant improvements in the models
128 and data assimilation schemes. The reanalyses differ in terms of the models and data assimi-
129 lation methods used to produce them, therefore differences in their output are to be expected. Six
130 recent global atmospheric reanalysis data sets have been analysed for TCs in this study and are
131 summarized in Table 1. They include the European Centre for Medium-Range Weather Forecasts
132 (ECMWF) Interim reanalysis (ERA-Interim) (Dee et al. 2011); the Japanese 25-year reanalysis (JRA25)
133 (Onogi et al. 2007) and 55-year reanalysis (JRA55) (Kobayashi et al. 2015); the National Aero-
134 nautics and Space Administration (NASA) Modern-era Retrospective Analysis for Research and
135 Applications (MERRA) (Rienecker and coauthors 2011) and the following version 2 (MERRA2)
136 (Bosilovich and coauthors 2015; Molod et al. 2015); and the National Centers for Environmental
137 Prediction (NCEP) Climate Forecast System Reanalysis (NCEP-CFSR) (Suranjana and Coauthors
138 2010). The NCEP-CFSR reanalysis is the only coupled atmosphere-ocean-land surface-sea ice
139 reanalysis. The NCEP-CFSR, MERRA and MERRA2 all use different versions of the 3D vari-
140 ational (3D-Var) data assimilation scheme: the Grid-point Statistical Interpolation (GSI) scheme
141 (Shao et al. 2016). For MERRA and MERRA2 the Incremental Analysis Update (IAU) (Bloom

142 et al. 1996; Rienecker and coauthors 2011) system is also used. The data period used for all the
143 reanalyses is 1979-2012, except for MERRA2, which starts in 1980. A key difference between the
144 JMA reanalyses and the reanalyses produced by the other agencies is the assimilation of tropical
145 wind retrievals (TWR). Wind profile data over and around tropical cyclone centers are retrieved
146 from historical data and processed and assimilated as if they were dropsonde observations (Hat-
147 sushika et al. 2006). With the integration of this additional wind data, the intensity of the storms
148 in the JMA reanalyses is found to be improved (Hatsushika et al. 2006). Another difference be-
149 tween the reanalyses is that the NCEP-CFSR uses a technique to improve the representation of
150 TCs by adjusting the location of the tropical vortex to its observed location before the assimilation
151 of storm circulation observations (Suranjana and Coauthors 2010). The MERRA2 reanalysis also
152 uses this method. All the reanalyses in this study make use of quality control processes and bias
153 correction for the diverse range of observations that are assimilated, for example, the variational
154 bias correction of satellite radiances (Dee and Uppala 2009).

155 *c. Tropical cyclone detection method*

156 The analysis of TCs in this study relies on identifying and tracking them. The first step is to track
157 all tropical disturbances, in both hemispheres, before applying two different identification methods
158 to separate the TCs from other tropical systems. This is different from some other schemes where
159 the identification is performed during the tracking and hence only identifies the TC stage of the
160 lifecycle. Though not crucial to this study, the approach taken here identifies much more of the
161 lifecycle, including the precursor and post extratropical transition stages (Jones et al. 2003).

162 For the first step, where all systems in the domain are tracked, the tracking methodology is based
163 on Hodges (1994, 1995, 1999). The domain extends to 60N in the NH and 60S in the SH. The
164 tracking method uses the 6 hourly relative vorticity at the levels 850, 700, 600hPa, vertically aver-

aged. This data is spectrally filtered using triangular truncation to retain total wavenumbers 6-63. The spectral coefficients are also tapered to further smooth the data using the filter described in Sardeshmukh and Hoskins (1984). The spectral filtering acts to remove the noise associated with the smallest spatial scales in the vorticity, which produces more reliable tracking in data of this type, and to remove the large scale background, which is also found to be beneficial. The tracking proceeds by identifying the off-grid vorticity maxima, by applying a maximisation scheme (Hodges 1995), that exceed a value of $5 \times 10^{-6} s^{-1}$ in each time frame (SH scaled by -1). These are initially linked together using a nearest neighbor approach and then refined by minimizing a cost function for track smoothness, subject to adaptive constraints on displacement distance and track smoothness (Hodges 1999). The use of the vertically averaged vorticity is different from some previous studies using this tracking algorithm, where the single level of 850hPa vorticity reduced to T42 resolution was used (Strachan et al. 2013; Roberts et al. 2015; Bell et al. 2013; Bengtsson et al. 2007b; Manganello et al. 2012). The use of the vertically averaged vorticity is found to improve the temporal coherency when a vorticity maximum shifts between levels (Serra et al. 2010; Fine et al. 2016) and results in more of the system lifecycle being detected. A simple vertical average is found to be sufficient, even though the levels are not evenly spaced, since, once spectrally filtered, there is little difference from using the mass weighted vertical average. Only tracks that last at least 2 days (8 time steps) are retained for further analysis. Whilst observed TCs can have lifetimes shorter than 2 days, this only covers the period when they are determined to be TCs, whereas the tracking scheme used here aims to identify the precursor and post-TC stages resulting in much longer lifetimes (see Figure 1c and d) so that using the 2 day threshold is not detrimental to detecting nearly all the observed TCs in the reanalyses, as shown below in the results section.

188 Previous methods used to detect TCs in reanalysis or GCM data rely on applying particular cri-
189 teria, representative of the properties of TCs, such as some thresholds on intensity, e.g. Mean Sea
190 Level Pressure (MSLP) minima, low level wind intensities or vorticity extrema, and a threshold
191 on the warm core structure either determined directly as a temperature anomaly, or inferred from
192 the presence of decreasing winds or vorticity between the lower and upper troposphere, for exam-
193 ple Bengtsson et al. (1995) and related methods. These are often applied as part of the tracking
194 scheme itself, which is different from the approach used here. A minimum period of one day is
195 typically imposed, for which these criteria are satisfied contiguously, and that they are satisfied
196 only over the ocean by imposing the land-sea mask. The criteria based on intensity and structure
197 can be strongly dependent on the model resolution and how processes important to TC devel-
198 opment, such as convection, microphysics and surface drag, are represented in the model. This
199 has resulted in some studies using resolution dependent identification criteria (Walsh et al. 2007;
200 Manganello et al. 2012) or tuning the identification criteria to maximize the detected TCs, for ex-
201 ample in reanalyses compared with observations (Murakami 2014), and some studies have used
202 basin dependent criteria (Camargo et al. 2005). The study of Horn et al. (2014) has shown that the
203 subjective choice of different identification criteria is the main reason for differences between the
204 numbers of TCs identified by different identification schemes.

205 In this study a dual approach is taken to isolate the TCs from all the tracked systems. Taking
206 the tracks identified in the first stage, where all systems are tracked, the first approach used to
207 isolate the TCs is used to see which of the observed TCs in the IBTrACS data set can be found in
208 the reanalyses, without applying any criteria dependent on intensity or structure. This approach
209 makes use of spatio-temporal matching: a track in the reanalyses matches with a track in IBTrACS
210 if the mean separation distance between them, computed over the time period that they overlap, is
211 less than 4^0 (geodesic), and is the least mean separation distance if more than one track satisfies

212 this criterion, where any amount of temporal overlap is allowed. This will be termed the "di-
213 rect matching" method. A similar approach has previously been used for extra-tropical cyclones
214 (Hodges et al. 2003). The relaxed criterion on the temporal overlap is chosen because, in general,
215 the TCs in IBTrACS have much shorter lifetimes compared to the tracks in the reanalyses pro-
216 duced by the tracking scheme. Several diagnostics are produced from the matched tracks, such as
217 the mean separation distance distribution, lifetime distribution and intensity distribution based on
218 low level winds, 10m and 925hPa, and MSLP.

219 The second approach used to isolate the TCs from all the tracked systems is to objectively
220 identify them using a typical set of identification criteria based on intensity and structure; this will
221 be termed the "objective detection" method. The criteria used are similar to those used previously
222 with this tracking algorithm (Bengtsson et al. 2007a,b; Strachan et al. 2013). This requires adding
223 additional fields to the tracks, namely the T63 vorticity at levels 850, 700-200hPa to provide
224 intensity and warm core criteria. This is done by recursively searching for a vorticity maximum
225 at the different levels using the maximum at the previous level as a starting point for a steepest
226 ascent maximization applied to the B-spline interpolated field. A search radius of 5° (geodesic)
227 is used centered on the location at the previous level. The same approach is used in the Southern
228 Hemisphere (SH) by multiplying fields by -1. Also added are the Mean Sea Level Pressure (MSLP)
229 minimum and maximum winds at 10m and 925hPa as alternative measures of TC intensity. For
230 MSLP a steepest descent method is used with the B-spline interpolation and a search radius of 5°
231 (geodesic) centered on the tracked vorticity center to find the closest pressure minimum, whilst
232 for the winds a direct search for the maximum winds within 6° of the tracked center is used. The
233 criteria for identification are:

- 234 1. the T63 relative vorticity at 850 hPa must attain a threshold of at least $6 \times 10^{-5} s^{-1}$.

- 235 2. the difference in vorticity between 850 and 200 hPa (at T63 resolution) must be greater than
236 $6 \times 10^{-5} s^{-1}$ to provide evidence of a warm core.
- 237 3. the T63 vorticity center must exist at each level between 850 and 200hPa for a coherent
238 vertical structure.
- 239 4. criteria (1) to (3) must be jointly attained for a minimum of 4 consecutive time steps (one
240 day) and only apply over the oceans.
- 241 5. tracks must start within 30S to 30N.

242 The approach used here means that the tracking and identification is performed at a common
243 resolution for all the reanalyses, making the tracking and identification as resolution independent
244 as possible, although the actual model resolution will still have some impact on the identification.

245 The TCs identified by the objective detection method are also matched against the observed
246 tracks in IBTrACS, using the same criteria as in the direct matching method, to determine the hit
247 and miss rates of the identification scheme.

248 The tracking is applied to each full year, January-December, for the Northern Hemisphere (NH)
249 and July to June the following year in the Southern Hemisphere (SH), resulting in 34 years in the
250 NH and 33 in the SH (33 and 32 respectively for MERRA2).

251 **3. Results**

252 In this section the ability of the different reanalyses to simulate different aspects of TC behavior
253 is assessed and compared to the observed TC activity, as represented by the IBTrACS database
254 described in the Best Track dataset subsection.

255 *a. Direct Matching Results*

256 The number of TCs in IBTrACS that match with a storm in the reanalyses for each reanalysis
257 using the direct matching method are summarized in Table 2 for both NH and SH. This shows that
258 $\sim 95\%$ of the TCs in IBTrACS are identified in the reanalyses in the NH and $\sim 92\%$ in the SH.
259 The different reanalyses are remarkably similar in this respect. In general the TCs not found in the
260 analyses tend to be the weakest and/or short lived TCs in IBTrACS in both hemispheres. Some of
261 the missing TCs fail to pass the 2 day lifetime threshold imposed on the reanalysis tracks. There is
262 also some evidence that the number of missing TCs in the reanalyses, according to the matching
263 criteria, are reduced in the later period after 2000: compared to the earlier period, the number of
264 matches increases to $\sim 98\%$ in both NH and SH. This improvement may be associated with the
265 assimilation of improved observations, in particular the availability of surface scatterometer winds
266 from the QuikSCAT satellite data from mid-1999 until the end of 2009 and continuing with similar
267 data from other remote sensing platforms since then.

268 To see how the TCs identified in the reanalyses by the direct matching method compare with
269 those in IBTrACS several sets of statistics are produced.

270 1) LOCATION

271 Figure 1a and b show distributions for the mean separation distance (geodesic distance) between
272 the identical reanalysis tracks and those of IBTrACS, obtained using the direct matching method,
273 in the NH and SH respectively. In the NH (Figure 1a) the majority of TCs identified in the re-
274 analyses have a mean separation from those in IBTrACS of less than 2^0 (220km), with the peak
275 of the distribution for each reanalysis typically at less than 1^0 (110km). The smallest mean sep-
276 aration distances occur for JRA55, with the distribution peak at 0.5^0 (56km) and the largest for
277 MERRA, with the distribution peak at 1^0 and the other reanalysis somewhere in between. The

278 JRA55 separation distances are comparable with those from the much higher resolution (T1279;
279 16km)) operational analyses of ECMWF (Hodges and Emerton 2015) (Appendix), which may be
280 a consequence of the assimilation of the TWR observations in JRA55. This conjecture is strength-
281 ened by the fact that JRA25, which also assimilates TWR data, is comparable in terms of the mean
282 separation distances to the much higher resolution NCEP-CFSR reanalysis. It is also apparent that
283 MERRA2 has improved over MERRA with respect to the separation distances. In general the
284 mean separation results for the NH (Figure 1a) are consistent with those found by Schenkel and
285 Hart (2012) for the identical reanalyses considered. In the SH (Figure 1b) a rather similar picture
286 is seen, with each of the reanalyses occurring in the same order as in the NH of best to worse.
287 Whilst the separation distances appear slightly larger for some reanalyses in the SH, i.e. ERAI and
288 MERRA, the others are comparable with the results in the NH, highlighting the improvement in
289 the SH in the more recent reanalyses compared with older reanalyses.

290 2) LIFETIME

291 Figure 1c and d show the lifetime distributions in the NH and SH respectively. In the NH it is
292 apparent that the TCs identified in the reanalyses have much longer lifetimes than the TCs in the
293 observations. This is a consequence of not imposing any criteria during the tracking to identify
294 TCs. Imposing the TC detection criteria during the tracking would truncate the tracks to the TC
295 stage alone, and would introduce a dependency of the lifetime on the chosen criteria and how well
296 TCs are represented in the reanalyses in terms of intensity and structure. The extended lifecycles
297 include pre-TC stages, e.g. easterly waves and the stage after extratropical transition. Some of
298 the reanalysis TCs can exist for longer than one month, in which time a precursor disturbance
299 can travel across an ocean basin, develop into a TC and recurve to high latitudes undergoing
300 extratropical transition, whereas none of the observed TC tracks last this long. The distributions for

301 the different reanalyses are quite close together, showing that rather similar lifetimes are obtained
302 for all the reanalyses. A similar set of results is obtained in the SH, although the distributions for
303 the reanalyses are a little noisier, due to the smaller number of observed TCs in this hemisphere.

304 3) LATITUDE OF MAXIMUM INTENSITY

305 The latitude at which the maximum intensity is attained in terms of the 10m winds is shown for
306 the NH and SH in Figure 1e and f respectively. In the NH the distributions show that, whilst most
307 TCs in the reanalyses attain their maximum intensity at similar latitudes to those in the observa-
308 tions, there are some TCs that attain their maximum intensity at much higher latitudes. A possible
309 cause for this behavior is that, because of the longer lifecycles that are identified in the reanalyses,
310 some storms only attain their maximum intensity as they recurve to higher latitudes and become
311 larger and better represented at synoptic scales. Whilst this could be addressed by restricting the
312 reanalysis tracks to just the TC stage, this would mean either truncating the tracks where they
313 overlap with the best track data (Hodges and Emerton 2015), or using the detection criteria based
314 on intensity and structure discussed above to define the TC part of the lifecycle. Either of these
315 approaches introduces a degree of subjectivity: the first as it depends on the different operational
316 practices of the operational agencies, and the second because it depends on how well TCs are rep-
317 resented in the different reanalyses. Also, for this part of the study, we want to see what exactly is
318 in the reanalyses in terms of TC lifecycle and restricting the lifecycles defeats this objective. This
319 is also important for future work, such as studies of extratropical transition and risk associated with
320 TCs and their later lifecycle stages in extratropical regions. A similar situation may also occur for
321 the TC stage itself, where the relatively low resolution of the reanalyses means that TCs are not
322 well represented at the small spatial scales of TCs in the tropics, but become better represented
323 as they move to higher latitudes. A similar picture is seen for the SH (Figure 1f). This type of

324 behaviour is often seen for TCs identified in relatively low resolution climate model simulations
325 (Manganello et al. 2012).

326 4) INTENSITY

327 Also examined are the maximum intensity distributions of the TCs for three intensity measures:
328 minimum MSLP and maximum 10m and 92hPa wind speeds, which are shown in Figure 2 for
329 both NH and SH TCs. For both MSLP (Figure 2a, b) and 10m wind speeds (Figure 2c, d) in the
330 NH and SH it is clear that all the reanalyses underestimate the intensity of TCs compared to the
331 observations and that the intensities are model dependent. This is not surprising considering the
332 relatively low spatial resolutions of the reanalyses where the assimilation of observations cannot
333 correct for this. Previous studies with dynamical downscaling of individual historical TCs, such
334 as Katrina, have shown that resolutions $\sim 1-5\text{km}$ with a non-hydrostatic model are necessary to
335 simulate TC inner-core processes correctly in order to enable the right magnitude of wind inten-
336 sities (Davis et al. 2008) to be simulated. However, some studies using hydrostatic models with
337 parameterized convection at resolutions $\sim 10\text{km}$ can certainly produce TCs with depths as large
338 if not larger than observed (Manganello et al. 2012). Coupling to the ocean has also been found
339 to be important in correctly simulating TC intensity (Kilic and Raible 2013), although only the
340 NCEP-CFSR reanalysis applies any such coupling and its impact on the reanalysis and TCs is
341 uncertain.

342 The results for intensity based on the MSLP (Figure 2a, b) show that in general the more recent
343 reanalyses, NCEP-CFSR, JRA55 and MERRA2 have deeper TCs; this is more evident in the SH,
344 although in both hemispheres few TCs reach depths below 940hPa. The more recent reanalyses
345 may be performing slightly better with respect to this intensity measure, possibly due to better
346 use of the available observations and improved models, and not necessarily due to resolution. For

347 10m wind speeds (Figure 2c, d), much larger differences are seen between the different reanal-
348 yses, although, as already mentioned, none of them can simulate the strongest intensities seen
349 in the observations. NCEP-CFSR has the most intense TCs in terms of 10m wind speeds, with
350 some TC almost attaining intensities of 50 ms^{-1} (Category 3 TS) but with no Category 4 or 5
351 (Saffir-Simpson scale) TSs. The weakest maximum 10m wind speed intensities are produced by
352 the MERRA reanalysis with no TCs surpassing 30 ms^{-1} , which barely reaches Category 1 TS.
353 However, the more recent MERRA2 reanalysis shows a significant improvement being compara-
354 ble with the JRA55 reanalysis in having TCs that can almost attain 10m wind speeds of 40 ms^{-1}
355 (Category 1 TS), although less than that seen for the NCEP-CFSR reanalysis. The results for the
356 reanalyses TC 10m wind speeds show similar behavior in both hemispheres. The results for both
357 10m wind and MSLP maximum intensities are generally consistent with those of Schenkel and
358 Hart (2012) for the NH. One problem with using the 10m winds from the reanalyses is that they
359 are not a direct model prognostic field, but are computed as a diagnostic, though not necessarily
360 in the same way for each reanalysis. They are generally computed as an extrapolation from the
361 lowest model level to the surface using profile functions and corrected when over land for terrain
362 roughness to conform to the World Meteorological Organization (WMO) standard for SYNOP
363 observations (for example see, ECMWF (2015)). However, for some reanalyses this is not done
364 for the actual analyses: for example in MERRA, it is performed during the IAU cycle, so does not
365 see the full analysis increment, and is an average over four model time steps (private communica-
366 tion, Michael Bosilovich, NASA). To evaluate the uncertainty further, the maximum wind speeds
367 at the 925hPa pressure level associated with the TCs are also considered (pressure level winds are
368 obtained by interpolation between model levels); the TC 925hPa winds are shown in Figure 2e,
369 f for the NH and SH respectively. The downside to using the 925hPa winds is that there are no
370 available observations with which to compare with, although this is not critical here, where we

371 just want to see if the same differences between the reanalyses, as seen for 10m winds, occur at
372 this level. The results for the wind speed intensity at 925hPa show a rather different perspective
373 from those at 10m, with both NCEP-CFSR and MERRA2 having comparable values in the tail of
374 the distribution with values as high as 60 ms^{-1} . The MERRA reanalysis is now comparable with
375 the other reanalyses of JRA55, JRA25 and ERAI.

376 5) WIND SPEED-PRESSURE RELATIONSHIP

377 The wind speed-pressure relationship is often used by the operational centers to estimate winds
378 from pressure measurements and surface pressure from wind measurements, for which various
379 quadratic empirical relationships have been developed based on cyclostrophic balance (Knaff and
380 Zehr 2007). Hence, the wind-pressure relationship of TCs is often considered in studies of TCs in
381 models and reanalyses (Roberts et al. 2015) to compare with the observed relationship, although
382 it should be noted that the observations may themselves be estimated from one of the empirical
383 relationships, which can differ between agencies (Knaff and Zehr 2007). Figure 3a shows the
384 wind-pressure relationship for the observations and the TCs identified in the different reanalyses
385 using the direct matching method in the NH. The wind-pressure relationship is determined using
386 the 10m wind speeds and MSLP values, by determining the maximum attained 10m wind speed
387 and taking the MSLP value at the same time. The results show that all the reanalysis reflect
388 the underestimate of both the 10m wind speeds and MSLP depths of the TCs, this being most
389 prominent for MERRA. This can be related to the radius of maximum wind (RMW), computed for
390 the reanalyses at the time of maximum 10m wind intensity, and shown for the NH in Figure 3c. The
391 RMW is not available for all the agencies in IBTrACS but we estimate it at the time of maximum
392 wind intensity, based on the simple Rankine model described by Knaff and Zehr (2007), this gives
393 RMW values for the observations predominately below $\sim 100\text{km}$ (1^0) and a peak around $\sim 50\text{km}$

394 (0.5⁰). This is consistent with the findings of Kimball and Mulekar (2004) for North Atlantic TSs
395 who made use of an extended "best track" data set. For all the reanalyses the RMW are seen to be
396 too large (Figure 3c). Assuming gradient wind balance for the TCs, and the fact that RMWs are
397 too large and wind intensities are too low for the reanalyses implies that the pressure difference
398 between the storm centers and the environment is also too low, consistent with the wind speed-
399 pressure relationship in Figure 3a. The fact that the NCEP-CFSR has the strongest wind intensities
400 and one of the smallest RMWs is also consistent with the result in Figure 3a that NCEP-CFSR is
401 closest to the observed wind speed-pressure relationship, whereas MERRA, which has the weakest
402 maximum wind speeds and large RMWs, is the worst of the reanalyses in this respect. MERRA2
403 shows a significant improvement over MERRA in terms of the wind speed-pressure relationship
404 which can be understood in terms of the improved maximum wind speeds and lower RMWs. In
405 fact, MERRA2 has the lowest RMWs, although is not as strong in intensity (10m wind speed) as
406 NCEP-CFSR.

407 The fact that NCEP-CFSR appears to perform the best in terms of the wind speed-pressure rela-
408 tionship may be the result of the vortex relocation scheme used by the NCEP-CFSR assimilation
409 system, which, as pointed out by Schenkel and Hart (2012), will result in improved vortex location,
410 which in turn may lead to improved TC intensities as a result of the TC being in the correct envi-
411 ronment. Allied to this, Schenkel and Hart (2012) also pointed out that observations within the TC
412 vicinity are less likely to be rejected by the assimilation scheme, due to smaller differences with
413 the first-guess field. However, the situation is likely more complex than this, as MERRA2 also
414 uses the vortex relocation method and has the lowest RMWs but is not the most intense in terms
415 of wind speed. JRA55, on the other hand, with a similar resolution to MERRA2, has the smallest
416 location errors (Figure 1a, b), does not use vortex relocation, but does assimilate best track data as
417 synthetic dropsondes (Hatsushika et al. 2006) and has comparable intensities to MERRA2 and a

418 wind speed-pressure relationship, also very similar to MERRA2. Hence, it appears that there are
419 complex trade-offs occurring within the assimilation systems. In the SH the wind-speed pressure
420 relationship (Figure 3b) and RMWs (Figure 3d) appear to be very similar to those in the NH: in
421 particular the wind-speed pressure relationship appears to be closely associated with the ordering
422 of the 10m wind speeds of the reanalyses shown in Figure 2b.

423 *b. Objective Identification*

424 Following the assessment of how well TCs are represented in the chosen reanalyses it is of
425 interest to see how existing objective TC identification schemes perform in order to try and un-
426 derstand the impacts of the differences between reanalyses on objective TC identification. This
427 is important, as objective schemes are the only way to identify TCs in climate model simulations
428 and they are often contrasted with reanalyses as a means of verification at comparable resolutions.
429 As Murakami (2014) has shown, detection schemes have to be tuned to particular reanalyses to
430 optimally detect TC/TS frequencies. This is also what tends to happen in operational settings,
431 where detection schemes are often tuned to a particular operational setup, so that applying them to
432 data from a different operational center can give very different numbers of detected TCs from the
433 in-house method (c.f. Fig. 22 of Kobayashi et al. (2015)). Some schemes also adjust identification
434 criteria by ocean basin (Camargo and Zebiak 2002) to account for model biases. However, these
435 are not appealing approaches in the climate model context, where a fixed set of criteria, applied in
436 a common resolution framework, will provide a better comparison between different model sim-
437 ulations or different climate scenarios (Shaevitz and Coauthors 2014). To assess how one such
438 scheme performs, the objective detection method described in the methodology section, based on
439 the vorticity at multiple levels between 850 and 200hPa, is applied to the vorticity tracks obtained
440 from the tracking of all vorticity centres.

441 1) ANNUAL COUNTS

442 The annual average TC counts are determined for each ocean basin (Figure 4) and are shown in
443 Figure 5. In the NH the annual number is in reasonably good agreement with the observations of
444 IBTrACS apart from MERRA, which has ~ 30 fewer identified TCs, whilst the other reanalyses
445 are slightly over or under in number, a result also previously noted by Murakami (2014) using the
446 same criteria. However, in the SH the identification has resulted in a much higher number than in
447 the observations, which occurs for all the ocean basins. The overestimation is particularly large
448 in the South Pacific (SP) region; the South Atlantic (SA) region also has more identified systems
449 than are in the observations. These differences will be discussed further in the Discussion section.

450 2) MATCHING AGAINST IBTRACS

451 To further analyse the objectively identified TCs, they are matched against the observed TCs of
452 IBTrACS using the direct matching method to identify the common storms between the two and
453 the false positive and negative detections. The results of this track matching are shown in Table 2
454 in terms of the Probability of Detection (POD) and False Alarm Rate (FAR). The POD is defined
455 here as the number of matched storms for each reanalysis divided by the total number of storms in
456 the observations, and the FAR by the number of non-matched storms in each reanalyses divided
457 by the total number of storms in the same reanalysis. Also shown in Table 2, for comparison, are
458 the POD for the direct matching results, before applying the objective criteria, discussed in the
459 Direct Matching Results subsection, which shows an almost uniform detection rate of 0.95 across
460 all the reanalyses in both hemispheres, although this is lower in the SH than the NH. The reason
461 why the POD for the SH is lower for the pre-criteria matching is likely related to differences in
462 the observations that are assimilated in the reanalyses between the two hemispheres, as there is no
463 dependence on structure or intensity for detection for these results.

464 For the POD based on using the objective detection method the values are much lower, with the
465 best detection for JRA55 and the worst for MERRA in both hemispheres, although POD is higher
466 in the SH than the NH, possibly due to differences in sample sizes. The FAR (Table 2) shows
467 values ranging from 0.16 for JRA25 to 0.36 for NCEP-CFSR in the NH. The fact that JRA25 has
468 the lowest FAR may be related to this reanalysis having the lowest resolution, hence, detecting
469 fewer small scale and possibly weaker storms; this could be investigated using GCMs of varying
470 resolution. In the SH, FAR is much higher, as might be expected from the previous discussion, due
471 mostly to the higher number of TCs detected compared with the observations. From these values
472 of POD and FAR it is apparent that, although similar numbers of TCs are detected in the NH using
473 the objective detection method, they need not be identical to the ones in the observations.

474 To explore the POD and FAR values in more detail the storms that are in the observations and that
475 match and do not match with those identified in the re-analyses, using both identification methods,
476 pre-objective direct matching and post-objective matching, are further analyzed relative to their
477 attained category in the observations according to the Saffir-Simpson scale determined from the
478 1min. observed winds. Hence, the IBTrACS storms are partitioned into the categories according to
479 the 1min. winds before matching them against the reanalysis tracks, as previously described. Since
480 different agencies use different wind intensity scales, this approach provides a more consistent
481 classification across the different ocean basins. Since some weak storms in IBTrACS have no wind
482 information, they are excluded from this analysis; Murakami (2014) excluded tropical depressions
483 from their study, although it is unclear how this is achieved for the reanalyses, apart from applying
484 the agency wind thresholds. The results of this analysis by category are shown in Tables 3 and
485 4 for the NH and SH respectively. In the NH, Table 3 shows that for the objectively identified
486 TCs it is the weakest categories that have the poorest level of matches between the reanalyses
487 and IBTrACS, in particular for the tropical depressions, although many tropical depressions in

488 IBTrACS are excluded due to lack of wind information. However, for the TS category (between
489 tropical depression and Category 1) the best performing reanalyses at this level, JRA25 and JRA55,
490 match with 78.5% of IBTrACS storms, while for the worst performing (MERRA) only 41.6% of
491 IBTrACS storms match. For the higher TS wind speed categories the percentage of matches
492 with IBTrACS steadily increases with category on progressively smaller sample sizes, i.e. 92%,
493 98%, 99.5% and 100% for CAT1-CAT5 respectively, for the best performing JRA25 and JRA55
494 and considerably worse for MERRA (63.5, 75, 83, 82.5, 92%) with NCEP-CFSR and MERRA2
495 comparable with JRA25 and JRA55. Re-calculating the POD for just Cat1-Cat5 TS (Table 5) the
496 best performing reanalyses, JRA25 and JRA55 now have values 0.95.

497 In the SH, Table 4 shows that a fairly similar situation occurs as in the NH for the objectively
498 identified TCs, except that it is apparent there are virtually no tropical depressions available to
499 compare with in the observations, either because very few of this category of storms have any
500 wind values, or more likely that they are not generally included in the best track data sets in this
501 hemisphere; this is discussed further in the discussion section. The best degree of matches again
502 occurs for the JRA25 and JRA55, ranging from 84-89% for the weakest TSs (TS Category) to
503 95% for CAT5.

504 The POD, for CAT1-CAT5 objectively identified TS only, shown in Table 5, shows that for this
505 intensity range the values are comparable in both hemispheres and comparable with the results
506 in the study of Murakami (2014) who restricted their study to this intensity range, although they
507 used different skill metrics compared to here and in the study here there is no special tuning of the
508 objective detection parameters for each reanalysis as in Murakami (2014).

509 For the TCs identified using the direct matching method (pre-objective), previously discussed
510 in the Direct Matching Results subsection, the matching by observation category (not shown)

511 indicates consistently high POD values as reported in the Direct Matching Results subsection for
512 all categories and reanalyses.

513 To understand the nature of the TCs, identified by the objective detection method, in the reanal-
514 yses that do not match with the IBTrACS TCs, in particular in the SH, those that do not match are
515 binned according to the latitude of their genesis. For the SH this is shown in Figure 6(a). This
516 shows essentially two groups of storms: those with genesis within 0-20S and those with genesis
517 occurring south of 20S. The genesis for all TCs in IBTrACS is almost entirely within 0-20S (not
518 shown). Examining these two groups on non-matching objectively identified TCs separately, a
519 scan of the tropical storm advisories (discussed later) indicates that some of the identified storms
520 in the first group can be found in the advisories but not IBTrACS; this is discussed further in the
521 Discussion section. Figure 6(b) shows examples of two tracks identified in ERAI that do not match
522 with IBTrACS: the track labeled "Storm 1" occurs in January of 2011 and is a storm that possibly
523 occurs in the RMSC Nadi advisories, named 02F, but is not in IBTrACS, probably because it did
524 not develop further into a true TS. Even so, it seems a substantial storm with 10m winds in ERAI
525 over 20 m s^{-1} whilst near Australia. Figure 6(c) shows the infrared satellite image, which presents
526 an asymmetric structure, unlike a true TS, with this storm more likely to be a hybrid warm core
527 TC. The second storm shown in Figure 6(b) originates south of 20S, where very few IBTrACS
528 storms have their genesis. This particular storm seems to have formed in the vicinity of the South
529 Pacific Convergence Zone (SPCZ) and travels south eastward with relatively weak 10m winds in
530 ERAI $\sim 15 \text{ m s}^{-1}$ through a region of very little habitable land. It has no reference in any tropical
531 storm advisories, yet its structure in the satellite imagery (Figure 6(d) shows some similarities with
532 "Storm 1" (Figure 6(c)) and it may also be a hybrid TC. As shown by Yanase et al. (2014) (Figure
533 1) using the Hart phase space classification of cyclones (Hart 2003), applied to reanalysis data,
534 storms found between 20-40S in the SH summer tend to be hybrid storms. There are also storms

535 in IBTrACS that do not match with an analysis track, but these tend to be the weakest storms
536 below Category 1 as shown in Tables 3 and 4. These issues are further discussed in the Discussion
537 section.

538 **4. Discussion**

539 There are several possibilities for the poorer performance of the objective detection method in
540 the SH compared with the NH in terms of the detection, relative to the observed TCs in IBTrACS.
541 As shown above, the discrepancy in numbers is closely associated with the weakest storms, trop-
542 ical depressions and tropical storms (below Category 1). The first possibility for the differences
543 between the NH and SH objective detection may be due to different biases in the best track data
544 in the SH compared with the NH; the second is due to different biases in the representation of
545 TCs in the reanalyses between the NH and SH; the third is due to the selection criteria used by the
546 objective detection method to identify TCs in the reanalyses being not selective enough, or being
547 mainly tuned to the NH. These will be addressed in turn.

548 In terms of possible biases in the IBTrACS observations, it is possible that the SH is observed
549 differently than in the NH. The SH is sparsely inhabited in particular regions, such as the SP and
550 SA, so that less emphasis may be placed on detection except for the most intense systems likely
551 to make landfall (Kucas et al. 2014). Related to this is the application of different storm detection
552 procedures in the different warning centers that produce the best track data (Velden and Coauthors
553 2006b; Kueh 2012). Storm classification is primarily based on the interpretation of satellite obser-
554 vations using empirical relationships such as the Dvorak scheme (Velden and Coauthors 2006a);
555 there is little aircraft reconnaissance apart for the North Atlantic with some other limited cover-
556 age associated with field campaigns and in specific regions, e.g. Taiwan (DOTSTAR) (Wu and
557 Coauthors 2005). The uncertainties of applying operational detection and classification schemes

558 when storms are relatively weak and show a poor organization (Torn and Snyder 2012) may make
559 deciding between whether a tropical disturbance should be classified as a tropical depression and
560 counted in best track, or is some other tropical storm such as a subtropical or hybrid cyclone, dif-
561 ficult and dependent on subjective forecaster interpretation. Gyakum (2011) states that "there is
562 presently no single set of objective criteria that, if applied operationally, would irrefutably support
563 a forecasters analysis of cyclone type (subtropical, hybrid or tropical)". It is also unclear whether
564 all agencies report weaker storms such as tropical depressions consistently in their best track anal-
565 yses, and hence whether they make their way into IBTrACS. For example, HURDAT, produced by
566 the National Hurricane Center (NHC), and which forms part of IBTrACS, and covers the North At-
567 lantic and North Eastern Pacific includes subtropical cyclones (Landsea and Franklin 2013) where
568 as the Joint Typhoon Warning Center (JTWC), which covers the Western North Pacific, South
569 Pacific and South and North Indian Oceans, do not routinely include subtropical cyclones (Kucas
570 et al. 2014; Gyakum 2011) unless they undergo Tropical Transition (TT) (Bentley et al. 2016;
571 McTaggart-Cowan et al. 2013). Even within a single ocean basin where multiple agencies are op-
572 erational, considerable uncertainties exist between different best track data sets. For example, Ren
573 et al. (2011) and Barcikowska et al. (2012), highlight significant differences between JTWC and
574 Japan Meteorological Agency (JMA) best track data in the Western North Pacific (WNP) in terms
575 of frequency and intensity of TCs, with better agreement for frequencies for Category 2 TS and
576 above; this is exactly where our objective detection scheme performs best in both hemispheres.

577 Therefore, uncertainties in the interpretation of the observations for the weaker tropical storms,
578 and different agency operational procedures, may result in their exclusion from the best track
579 archive. Several reassessments of best track data, in particular in the SH, have resulted in the inclu-
580 sion of some additional storms, but also the removal of some others (Diamond et al. 2012) so that
581 actual numbers are not significantly changed. However, evidence that the SH may be being treated

582 differently for tropical storms in the observations than in the NH, in particular with respect to the
583 weaker sub-tropical and hybrid storms, can be seen by considering the tropical storm advisories.
584 Information on weak tropical disturbances, together with TCs, is available in text based reports
585 from the warning agencies, such as the JTWC "significant tropical cyclone advisories". However,
586 not all this information is included in the best track post season analysis and hence IBTrACS.
587 For example, in the South Pacific, IBTrACS reports 5 storms in the 2011/2012 season (July-June)
588 but scanning the advisories (Regional Specialized Meteorological Center (RSMC) Nadi) results
589 in a much larger number of tropical disturbances, ~ 20 . A more quantitative comparison can be
590 made using the combined advisories from each warning center, for each year, in each hemisphere
591 (July-June in the SH). This information has been collated by Padgett and Young (2016) from 1998
592 onwards for both hemispheres, although some very weak systems are not included. Comparing
593 the numbers in the advisories with those in IBTrACS over the period 1998-2012, which overlaps
594 with our study period; in the NH, IBTrACS has on average 69 storms per year and the advisories
595 72, hence the advisories have $\sim 4\%$ more storms, whereas for the SH, IBTrACS has on average 28
596 storms per year and the advisories 39, hence the advisories have $\sim 40\%$ more storms. Hence in the
597 NH it appears that a much larger proportion of the storms in the advisories make their way into
598 the best track data than in the SH. This can partially explain the difference in numbers between
599 IBTrACS and the TCs identified by the objective detection method in the reanalyses in the SH. It
600 was discussed in the Matching Against IBTrACS subsection that some of the storms identified in
601 the reanalyses appeared to be in the advisories but not IBTrACS.

602 Tropical disturbances and subtropical cyclones occur in all the ocean basins, and it seems that
603 whether or not they contribute to the best track data may vary between the NH and SH and be
604 dependent on the warning center procedures. The SPCZ and South Atlantic Convergence Zone
605 (SACZ) are known to be associated with weak tropical depressions and subtropical cyclones in the

606 SH, as well as more intense tropical cyclones in the South Pacific (Vincent et al. 2011). A similar
607 situation occurs in the North Pacific associated with the Mei-Yu front (Lee et al. 2006). The South
608 Atlantic is not known as a very active TC region, due to relatively cool Sea Surface Temperatures
609 and relatively high vertical wind shear. However, several studies have highlighted this region as
610 susceptible to the formation of subtropical cyclones (Evans and Braun 2012; Gozzo et al. 2014)
611 often in association with the SACZ. This is also seen in simulations produced with high-resolution
612 GCMs where they are often identified as TCs (Roberts et al. 2015). The study of Gozzo et al.
613 (2014), based on reanalysis data, found on average 7 subtropical cyclones per year with genesis
614 between 20-30⁰S, a number that is remarkably similar to the number of systems objectively de-
615 tected in the reanalyses in this study in the SA region. The majority of the sub-tropical cyclones
616 identified by Gozzo et al. (2014) do not seem to have made it into the advisories or best track data,
617 either because they are too weak, even for the advisories, or possibly because in general they are
618 moving away from land and therefore not a threat (Kucas et al. 2014). Another possibility is that
619 SA sub-tropical cyclones are more asymmetric than those found in the North Atlantic (Evans and
620 Braun 2012) and hence do not satisfy the criteria for inclusion in the TC best tracks. A similar
621 situation may also occur in the South Pacific. If these additional uncertainties in the best track data
622 are considered together with the numbers in the advisory data, then the actual numbers of TCs
623 occurring in the SH may not be too far away from the numbers objectively identified here in the
624 reanalyses. The results from the Matching Against IBTrACS subsection suggest that some of the
625 differences between numbers in the SH between the objective identification used in this study and
626 IBTrACS may be related to the identification of hybrid or sub-tropical cyclones by the objective
627 identification scheme.

628 Other regions where subtropical or hybrid storms may need to be considered are the cool seasons
629 in the eastern north Pacific, where they are called Kona storms (Kodama and Businger 1998).

630 Monsoon depressions may also be confused with weak tropical cyclones in the reanalyses as these
631 also have a warm core aloft structure and occur in the north and south Indian ocean, western
632 Pacific and Australian region (Hurley and Boos 2015). They represent an additional uncertainty
633 in the best track archive, as they are occasionally included in the best track data in the western
634 Pacific via the JTWC (Hurley and Boos 2015); however, as with subtropical cyclones, this is not
635 done consistently for all agencies. These may also contribute to uncertainties in the best track data
636 in the north and south Indian ocean and South Pacific.

637 The second possibility for the differences in the numbers of TCs detected by the objective detec-
638 tion method in the reanalyses and IBTrACS in the NH and SH concerns the quality of the reanaly-
639 ses in the two hemispheres, which may affect how TCs are represented and hence contribute to the
640 uncertainties in their detection in the reanalyses. The primary observations assimilated in the SH
641 come from satellite observing platforms, which generally provide relatively coarse vertical reso-
642 lutions, whereas in the NH the surface-based observing system provides a more diverse range of
643 observations, including from sondes and aircraft. The use of direct satellite radiance assimilation,
644 variational bias correction and modern assimilation methods has resulted in much better extraction
645 of the information content in the observations, including for older observations (Rienecker et al.
646 2012). Discriminating between weak TSs, sub-tropical cyclones and other systems in the reanal-
647 yses is a problem in both hemispheres for the objective detection method, but could be more of
648 a problem in the SH if the TCs are not as well simulated and storms, including sub-tropical or
649 hybrid storms, do not have the correct structure. This could be exacerbated if there are more of
650 the weaker type of storms in the SH associated with the convergence zones as discussed above,
651 which, allied to the difficulty in separating these storms from other systems, may be a factor in
652 the differences between the number of storms in IBTrACS and the number detected by the objec-
653 tive detection method in the reanalyses in the SH. The only way to test this is by using observing

654 system experiments, where the NH observing system is degraded to that of the SH and the data as-
655 simulation re-run. These types of experiments have been performed in the past and have shown the
656 relative importance of the different types of observations used in the reanalyses and how changes
657 to the observing system may affect the reanalysis (Bengtsson et al. 2004b; Whitaker et al. 2009).
658 However, it is very time consuming and expensive to re-run modern data assimilation systems,
659 even if we had access to the same systems used to produce the reanalyses used here. Hence this is
660 beyond the scope of this paper. However, studies using the same detection criteria as used here,
661 applied to relatively high resolution climate model simulations for the current climate (Gleixner
662 et al. 2013; Strachan et al. 2013; Roberts et al. 2015), have found similar results to those found
663 here for the reanalyses, in that similar TC numbers to observations are found in the NH, albeit with
664 some model dependent basin by basin biases, and a larger number of TCs than in the observations
665 in the SH. This may indicate that the difference in the number of SH storms from the observations
666 are not necessarily related to differences in the quality of the reanalyses in the two hemispheres,
667 but more with possible biases in the best track data and possibly the detection criteria used in our
668 objective scheme, discussed next.

669 The larger bias in the number of TCs identified by the objective detection method in the SH
670 compared with the NH relative to observations may also be related to the detection criteria used
671 here, and whether they are selective enough for the data used, so that more tropical depressions,
672 subtropical cyclones and hybrid cyclones are identified as TCs, possibly related to the quality of
673 the reanalyses as discussed above. TC detection schemes, applied to model or reanalysis data,
674 are certainly sensitive to the detection criteria and tracking methodology employed (Horn et al.
675 2014), especially for weaker storms, as shown in this paper, and are most often tuned for the
676 NH. An alternative approach would be to apply more selective criteria to remove subtropical and
677 hybrid cyclones from the detection, based on previous studies focussed on studying subtropical

678 cyclones, for example the Hart phase space parameters (Guishard et al. 2009; Evans and Braun
679 2012; Yanase et al. 2014). Another idea in the literature suggests using TC development pathways
680 (McTaggart-Cowan et al. 2013), whereby tropical cyclogenesis is categorized according to dy-
681 namical metrics, although this would necessarily introduce added complexity and possibly more
682 parameters to choose subjectively. It would also remove these types of storms in the NH, so that,
683 whilst the numbers detected in the SH may compare better with the observations, the numbers may
684 compare less favourably in the NH. However, it might allow a better focus on the different storm
685 types.

686 It is likely that all three of the issues discussed above can lead to TC detection biases in the
687 reanalyses relative to the best track data.

688 No TC tracking and/or identification scheme will be perfect and, although TC identification
689 schemes can be re-tuned against the observations separately for the NH and SH or for individual
690 ocean basins if necessary (Camargo and Zebiak 2002) to take account of possible deficiencies in
691 the detection and the observational biases, this does not seem like a good idea if TC detection is
692 to be applied to model simulations where methodological consistency is important.

693 **5. Summary and conclusions**

694 The study of TCs in six recent reanalyses has shown that all the reanalyses are capable of rep-
695 resenting nearly all the TCs present in the best track archive of IBTrACS, with a detection rate of
696 $\sim 98\%$ in the period since 2000 and slightly lower before this. However, how well the TCs are
697 represented in the reanalyses, in terms of their properties, is less encouraging, with wind intensities
698 significantly lower than in the observations and pressures too high in value. Although significant
699 amounts of observations are assimilated by the data assimilation systems used in the reanalyses,
700 in particular from satellites, this is unable to correct these deficiencies in the TC properties, due

701 to the still too low model resolution and dependence on parameterized processes used in the re-
702 analyses. Additional methods of assimilating observations in the vicinity of the TCs and vortex
703 relocation can help improve this situation, but not to the extent where intensities get anywhere near
704 those observed at current reanalysis resolutions. However, it is apparent that there has been some
705 improvements in the representation of TCs in the more recent reanalyses of NCEP-CFSR, JRA55
706 and MERRA2; in particular MERRA2 shows a significant improvement over the older MERRA
707 reanalysis in terms of wind and MSLP intensities. Separation distances between TCs identified
708 in the reanalyses and the observations have also improved with the more recent reanalyses. The
709 improvements in the intensities and location are most likely due to the increases in model horizon-
710 tal resolutions and the use of improved data assimilation and bias correction systems, which are
711 capable of extracting more information content from the older observations, as well as resulting in
712 less observation rejection and the introduction of new and better calibrated observing systems in
713 recent years. This progress is likely to continue as new reanalyses are produced with ever higher
714 resolutions, such as the new ECMWF ERA5 reanalysis. Further improvements in data assimilation
715 are expected as well as the introduction of new and more accurate observing systems, although the
716 downside to this may be the introduction of spurious trends in TC properties.

717 The other aspect explored in this study is how well objective TC detection schemes are capable
718 of detecting the same TCs that are in the observations using a widely used identification scheme.
719 This is important in order to have confidence in these schemes when applied to climate model sim-
720 ulations and for comparisons made between models or experiment scenarios. This part of the study
721 highlighted the problem of detecting TCs at the low intensity end of the TC intensity range: in par-
722 ticular, tropical depressions and up to category 1 (Saffir-Simpson), with gradual improvements in
723 the detection rate with increasing TS category. This raises several issues: are the current detection
724 schemes used at operational centers and for climate studies of TCs, which all have a rather similar

725 methodology of user chosen thresholds on intensity and/or structure, selective enough; are TCs
726 represented well enough in the reanalyses; are there problems with observational biases for weak
727 storms? The answer to these questions is probably that all three play a role in differences found
728 between the objective identification of TCs in reanalyses and the observed best track data. It is
729 clear the intensities, and probably structure, are not well enough simulated in the reanalyses, which
730 will cause problems when trying to discriminate between weak TSs and other tropical systems. In
731 terms of more selective criteria, other approaches could certainly be introduced, such as the phase
732 space approach, but this will also depend on how well TCs are represented in the reanalyses and
733 the introduction of subjective thresholds on the phase space parameters (Yanase et al. 2014). How-
734 ever, it may be useful in removing the need for artificial boundaries in the TC identification such
735 as the latitude band for genesis used in this study. The problem of observational bias is also an im-
736 portant aspect, in particular for the weaker storms, since forecaster interpretation and subjectivity
737 will play a role in whether a particular storm is included in the best track data, as not all storms fall
738 neatly into particular classifications. Allied to this are the different operational criteria employed
739 by the different RSMC, which contribute data to the best track archives, such as whether to include
740 tropical depressions or subtropical cyclones. This is likely the primary cause of the differences be-
741 tween the number of TCs identified in the reanalyses and IBTrACS, in particular in the SH. This
742 makes the observations less than ideal for calibrating TC identification and tracking schemes, or
743 indeed in their use in global climatological studies of TC frequencies and variability. It could be
744 concluded that, given the uncertainties in the best track data sets, they should not be considered
745 climate quality data sets. Better coordination between the RSMCs would help this situation go-
746 ing forward, although this is not necessarily part of their remit and their operational procedures
747 are tailored to their region of responsibility. The problems of objectively classifying TCs opera-
748 tionally has been recognized by the Seventh International Workshop on Tropical Cyclones who

749 have suggested that "a substantial contribution to the operational TC forecasting community could
750 be made by recommending a universal cyclone classification methodology based on the latest re-
751 search, operational forecasting capabilities, and real-time data availability". A re-evaluation of the
752 observational record over the satellite period using a combination of the satellite data and reanaly-
753 ses, using consistent identification methods for all basins, could perhaps resolve the observational
754 bias problem over historical periods covered by the satellites and provide a more complete record
755 of tropical storms for use in risk assessment and validating climate models. There has been some
756 discussion that tropical depressions and subtropical cyclones should be included in the best track
757 data for consistency (McAdie et al. 2009), since, before satellite observations became available,
758 some subtropical systems were probably classified as TSs. Tropical depressions and subtropical
759 cyclones are also associated with severe weather with TS like properties of strong winds and pre-
760 cipitation (Guishard et al. 2009; Gyakum 2011), so their inclusion can be justified in terms of their
761 impact and for a more complete record of TC activity.

762 Whilst there are deficiencies in the representation of TCs in the reanalyses, and 10 m winds in
763 particular should be used with caution, they can be complementary to the observations and provide
764 added value information on TCs such as the pre- and post-TC stages of the lifecycle. For example,
765 the tracking method used here identifies these earlier and later lifecycle stages, which can then be
766 used to study the early development of TCs and their environment and the extratropical transition
767 (Studholme et al. 2015) and how storms behave after this. The extratropical transition and its
768 aftermath are becoming increasingly important for risk analysis at high latitudes following cases
769 such as hurricane Sandy and Gonzalo and recent studies such as Haarsma et al. (2013) and is a
770 known contributor to forecast uncertainty in the extra-tropics (Anwender and Harr 2008).

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988 **LIST OF TABLES**

989 **Table 1.** Summary of the reanalysis datasets used in this study. Abbreviations, 4D-Var,
990 4D variational assimilation ; 3D-Var, 3D variational assimilation; TL255L60,
991 triangular truncation 255, linear grid, 60 vertical levels; GSI, Grid-point statisti-
992 cal interpolation; IAU, Incremental Analysis Update. 47

993 **Table 2.** The POD for the NH and SH for the direct matching method applied to the
994 reanalysis tracks (c.f. Direct Matching Results section) and the POD and FAR
995 for the NH and SH based on the objective detection method (c.f. Matching
996 Against IBTrACS subsection). 48

997 **Table 3.** Storms that match and don't match with IBTrACS in the NH by storm category,
998 for storms identified by the objective detection method applied to the reanaly-
999 sis tracks, in the first column for each reanalysis, and for the direct matching
1000 method performed in the Direct Matching Results subsection in the second col-
1001 umn with brackets for each reanalysis. Values are number per year. 49

1002 **Table 4.** Same as Table 2 but for the SH. 50

1003 **Table 5.** The POD for the NH and SH for the TC obtained from the reanalyses by the
1004 objective detection method that match with the observed Cat1-Cat5 TS only. 51

1005 TABLE 1. Summary of the reanalysis datasets used in this study. Abbreviations, 4D-Var, 4D variational
 1006 assimilation ; 3D-Var, 3D variational assimilation; TL255L60, triangular truncation 255, linear grid, 60 vertical
 1007 levels; GSI, Grid-point statistical interpolation; IAU, Incremental Analysis Update.

	<i>ERA-Interim</i>	<i>JRA25</i>	<i>JRA55</i>	<i>NCEP-CFSR</i>	<i>MERRA</i>	<i>MERRA2</i>
Assimilation	4D-Var	3D-Var	4D-Var	3D-Var	3D-VAR	3D-Var
				GSI	GSI+IAU	GSI+IAU
Model	TL255L60	T106L40	TL319L60	T382L64	$1/2^0 \times 2/3^0$ L72	Cubed Sphere
Resolution	(80km)	(120km)	(55km)	(38km)	(55km)	(50km)
Data Grid	512×256	288×145	288×145	720×361	540×361	576×361

1008 TABLE 2. The POD for the NH and SH for the direct matching method applied to the reanalysis tracks (c.f.
 1009 Direct Matching Results section) and the POD and FAR for the NH and SH based on the objective detection
 1010 method (c.f. Matching Against IBTrACS subsection).

POD						
	<i>ERA-Interim</i>	<i>JRA25</i>	<i>JRA55</i>	<i>NCEP-CFSR</i>	<i>MERRA</i>	<i>MERRA2</i>
NH Direct Match	0.95	0.95	0.95	0.95	0.95	0.95
NH Objective	0.60	0.76	0.80	0.70	0.51	0.67
SH Direct Match	0.93	0.93	0.94	0.93	0.90	0.93
SH Objective	0.76	0.84	0.87	0.83	0.61	0.79
FAR						
NH Objective	0.28	0.16	0.29	0.36	0.21	0.36
SH Objective	0.60	0.43	0.58	0.58	0.54	0.63

1011 TABLE 3. Storms that match and don't match with IBTrACS in the NH by storm category, for storms identified
1012 by the objective detection method applied to the reanalysis tracks, in the first column for each reanalysis, and
1013 for the direct matching method performed in the Direct Matching Results subsection in the second column with
1014 brackets for each reanalysis. Values are number per year.

Category		<i>ERA-I</i>	<i>JRA25</i>	<i>JRA55</i>	<i>NCEP – CFSR</i>	<i>MERRA</i>	<i>MERRA2</i>
TD	Match	2.91 (7.94)	3.26 (7.94)	5.24 (8.03)	3.50 (8.00)	2.29 (7.85)	3.48 (7.67)
	No Match	5.56 (0.53)	5.21 (0.53)	3.24 (0.44)	4.97 (0.47)	6.18 (0.62)	4.91 (0.73)
TS	Match	11.85 (22.38)	18.62 (22.53)	18.32 (22.53)	14.76 (22.32)	9.85 (22.44)	14.24 (22.45)
	No Match	11.85 (1.32)	5.09 (1.18)	5.38 (1.18)	8.94 (1.38)	13.85 (1.26)	9.73 (1.52)
CAT1	Match	8.74 (12.23)	11.18 (12.23)	11.17 (12.24)	10.09 (12.12)	7.74 (12.21)	9.76 (12.33)
	No Match	3.44 (0.00)	1.00 (0.00)	1.00 (0.00)	2.09 (0.06)	4.44 (0.00)	2.55 (0.00)
CAT2	Match	5.29 (6.35)	6.15 (6.38)	6.00 (6.35)	5.82 (6.38)	4.76 (6.35)	5.64 (6.39)
	No Match	1.06 (0.00)	0.21 (0.00)	0.35 (0.00)	0.53 (0.00)	1.59 (0.00)	0.73 (0.00)
CAT3	Match	6.15 (7.00)	6.91 (7.06)	6.82 (7.03)	6.71 (7.06)	5.82 (7.03)	6.42 (7.06)
	No Match	0.88 (0.03)	0.12 (0.00)	0.21 (0.00)	0.32 (0.00)	1.21 (0.00)	0.64 (0.00)
CAT4	Match	5.97 (6.79)	6.76 (6.79)	6.71 (6.74)	6.47 (6.79)	5.76 (6.76)	6.48 (6.76)
	No Match	0.82 (0.00)	0.03 (0.00)	0.09 (0.06)	0.32 (0.00)	1.03 (0.03)	0.33 (0.06)
CAT5	Match	1.09 (1.12)	1.12 (1.12)	1.12 (1.12)	1.12 (1.12)	1.03 (1.12)	1.09 (1.09)
	No Match	0.03 (0.00)	0.00 (0.00)	0.00 (0.00)	0.00 (0.00)	0.09 (0.00)	0.00 (0.00)

TABLE 4. Same as Table 2 but for the SH.

Category		<i>ERA1</i>	<i>JRA25</i>	<i>JRA55</i>	<i>NCEP – CFSR</i>	<i>MERRA</i>	<i>MERRA2</i>
TD	Match	0.42 (0.58)	0.48 (0.61)	0.52 (0.58)	0.48 (0.61)	0.21 (0.49)	0.44 (0.48)
	No Match	0.18 (0.03)	0.12 (0.00)	0.09 (0.03)	0.12 (0.00)	0.39 (0.12)	0.13 (0.06)
TS	Match	7.15 (9.09)	8.03 (9.06)	8.55 (9.09)	7.76 (9.18)	5.67 (9.00)	7.59 (8.91)
	No Match	2.42 (0.48)	1.55 (0.51)	1.03 (0.48)	1.82 (0.39)	3.91 (0.58)	2.00 (0.39)
CAT1	Match	4.55 (5.36)	5.09 (5.36)	5.12 (5.39)	4.94 (5.39)	3.79 (5.33)	4.75 (5.21)
	No Match	0.88 (0.06)	0.33 (0.06)	0.30 (0.03)	0.48 (0.03)	1.64 (0.09)	0.69 (0.06)
CAT2	Match	2.21 (2.64)	2.58 (2.61)	2.55 (2.64)	2.52 (2.64)	2.03 (2.61)	2.38 (2.61)
	No Match	0.61 (0.18)	0.24 (0.21)	0.27 (0.18)	0.30 (0.18)	0.79 (0.21)	0.53 (0.21)
CAT3	Match	2.55 (2.73)	2.61 (2.70)	2.70 (2.73)	2.64 (2.70)	2.12 (2.70)	2.63 (2.73)
	No Match	0.18 (0.00)	0.12 (0.03)	0.03 (0.00)	0.09 (0.03)	0.61 (0.03)	0.19 (0.00)
CAT4	Match	2.69 (2.76)	2.73 (2.73)	2.64 (2.76)	2.76 (2.76)	2.33 (2.76)	2.78 (2.76)
	No Match	0.06 (0.00)	0.03 (0.03)	0.12 (0.00)	0.00 (0.00)	0.42 (0.00)	0.06 (0.00)
CAT5	Match	0.58 (0.58)	0.51 (0.52)	0.55 (0.58)	0.55 (0.55)	0.55 (0.58)	0.59 (0.58)
	No Match	0.00 (0.00)	0.06 (0.06)	0.03 (0.00)	0.03 (0.03)	0.03 (0.00)	0.00 (0.00)

1015 TABLE 5. The POD for the NH and SH for the TC obtained from the reanalyses by the objective detection
 1016 method that match with the observed Cat1-Cat5 TS only.

	POD					
	<i>ERA1</i>	<i>JRA25</i>	<i>JRA55</i>	<i>NCEP-CFSR</i>	<i>MERRA</i>	<i>MERRA2</i>
NH Objective	0.81	0.96	0.95	0.90	0.75	0.87
SH Objective	0.88	0.91	0.95	0.94	0.75	0.92

1017 **LIST OF FIGURES**

1018 **Fig. 1.** Distribution of mean separation distances (geodesic degrees, $1^0 \simeq 111km$) between the re-
1019 analysis tracks and those of IBTrACS for tracks that match using the direct matching method
1020 (c.f. Direct Matching Results subsection) (a) NH, (b) SH, distribution of lifetimes (days) for
1021 the matched tracks (c) NH, (d) SH and the distribution of latitudes at which the matched
1022 tracks attain the peak intensity based on the 10m winds (e) NH, (f) SH. 53

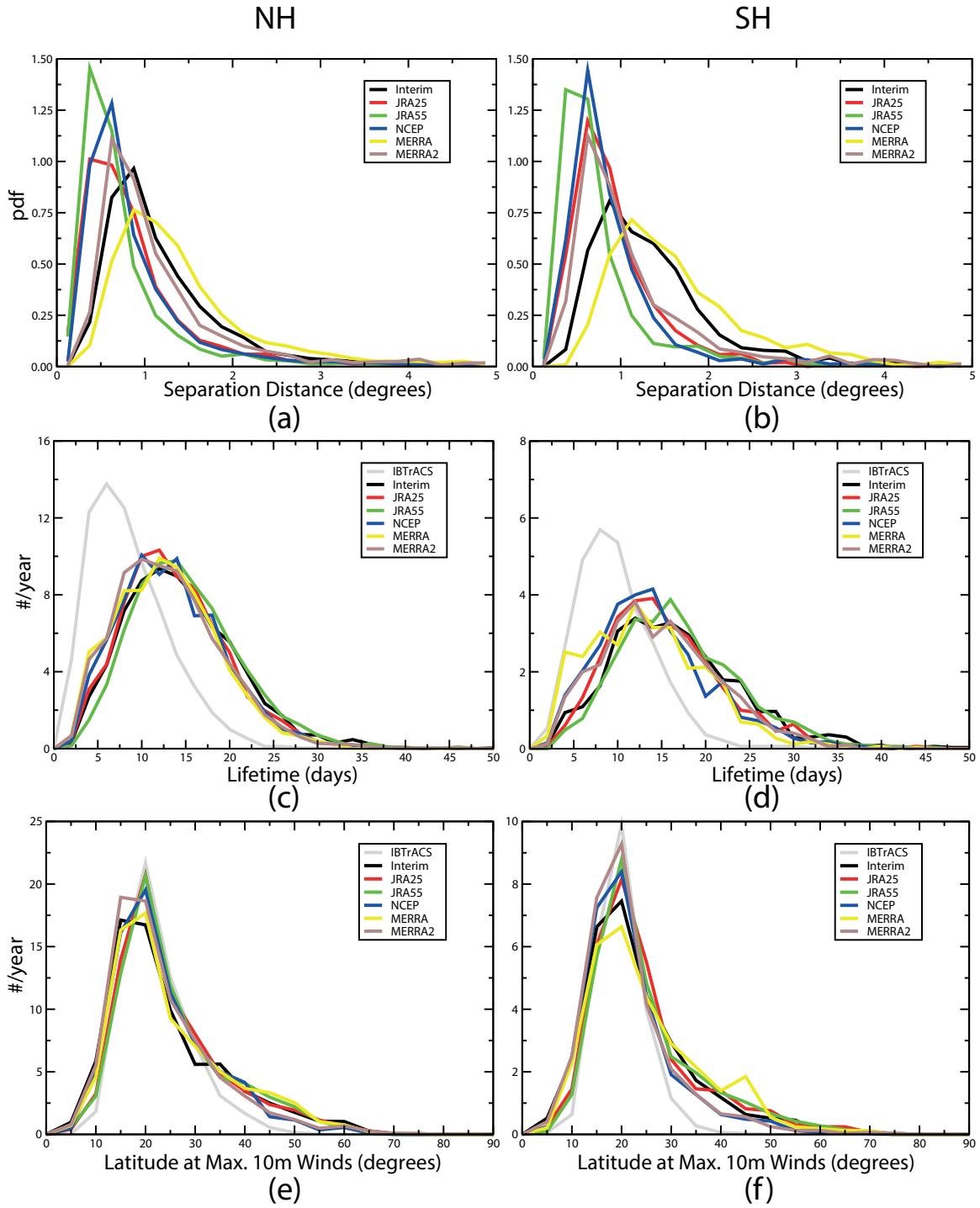
1023 **Fig. 2.** Distributions for the peak attained intensities of matched reanalysis and IBTrACS tracks
1024 obtained using the direct matching method (c.f. Direct Matching Results subsection) based
1025 on the MSLP, 10m winds and 925hPa winds (not for IBTrACS), (a) NH MSLP (hPa), (b)
1026 SH MSLP (hPa), (c) NH 10m wind speed ($m s^{-1}$), (d) SH 10m wind speed ($m s^{-1}$), (e) NH
1027 925hPa wind speed ($m s^{-1}$), (f) SH 925hPa wind speed ($m s^{-1}$). 54

1028 **Fig. 3.** Wind-pressure relationships for IBTrACS and each reanalysis and distributions for the radius
1029 of maximum winds for the reanalyses based on the direct matching method (c.f. Direct
1030 Matching Results subsection). (a) NH 10m wind speed versus MSLP, (b) SH 10m wind
1031 speed versus MSLP, (c) NH radius of maximum winds and (d) SH radius of maximum winds. 55

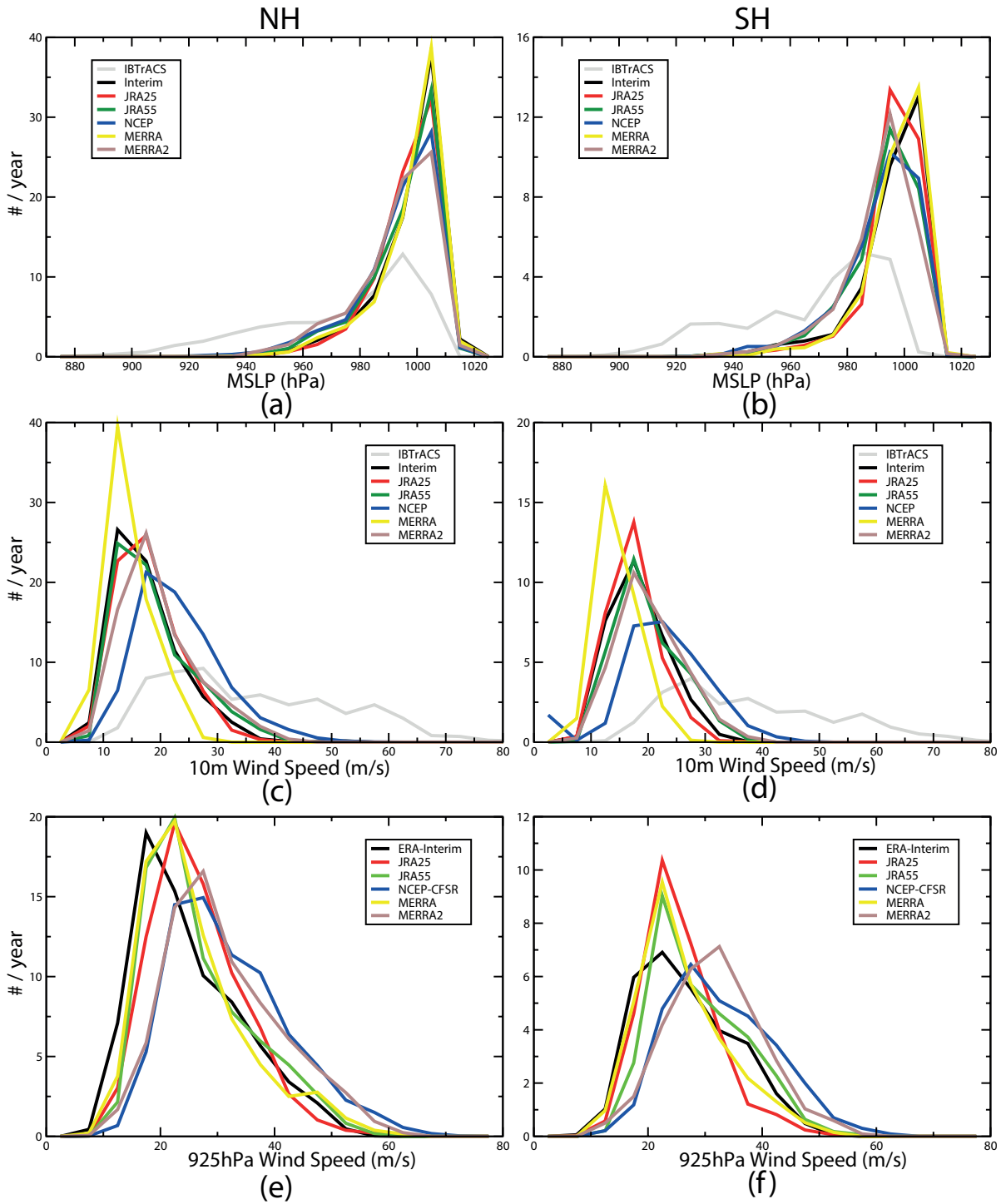
1032 **Fig. 4.** The seven basins used in this study, based on the IBTrACS definition. NI: North Indian,
1033 WP: West Pacific, EP: East Pacific, NA: North Atlantic, SI: South Indian, SP: South Pacific,
1034 SA: South Atlantic. 56

1035 **Fig. 5.** The average number of TCs per year for each of the seven basins for IBTrACS and identi-
1036 fied in the reanalyses based on the objective detection method (c.f. Objective Identification
1037 subsection). Vertical lines indicate the standard deviation. NI: North Indian, WP: West Pa-
1038 cific, EP: East Pacific, NA: North Atlantic, SI: South Indian, SP: South Pacific, SA: South
1039 Atlantic. 57

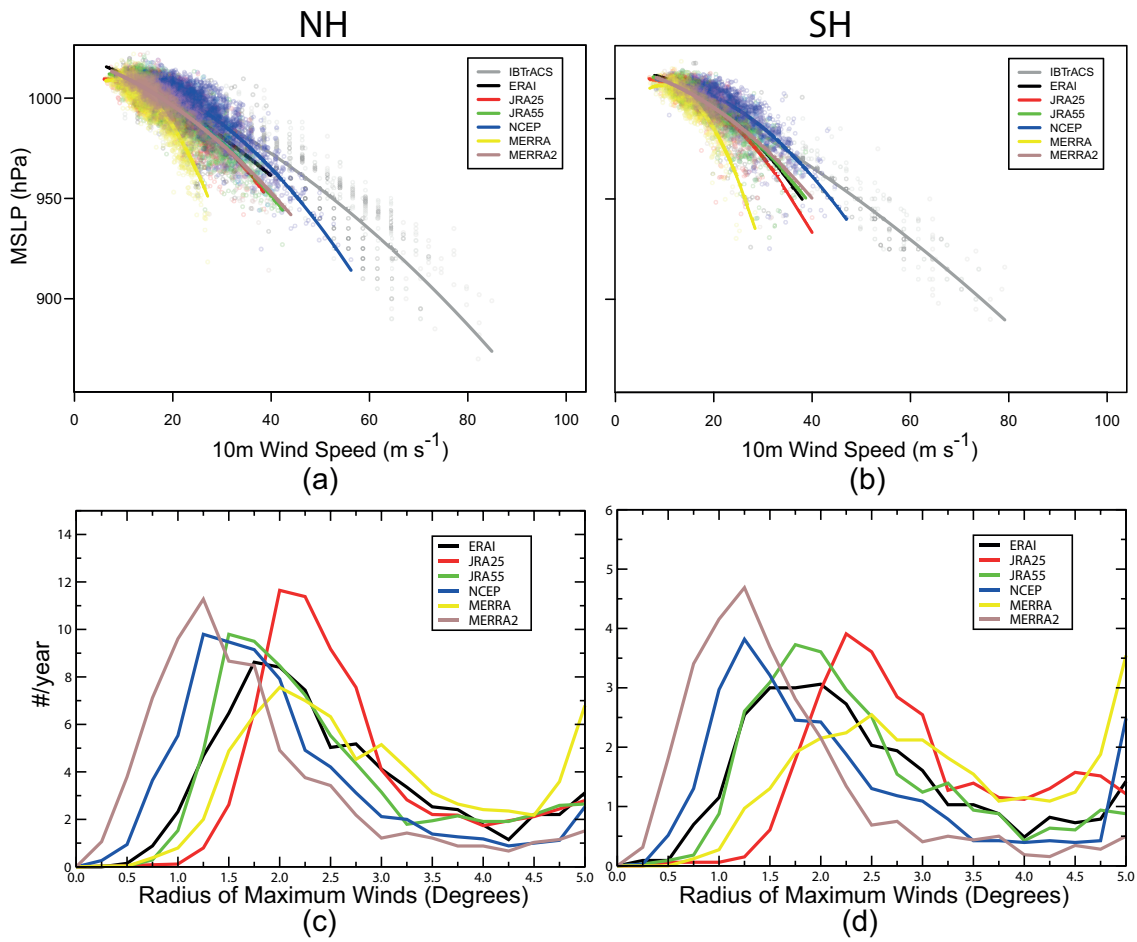
1040 **Fig. 6.** (a) Latitude at which genesis occurs in the SH for the objectively identified TCs in the
1041 reanalyses that do not match with IBTrACS (number per year), (b) examples of two tracks
1042 identified in the ERAI reanalysis with no matching track in IBTrACS, coloured dots indicate
1043 10m wind speeds ($m s^{-1}$), (c) MTSAT infrared satellite image of Storm 1 in (b) on 1800
1044 UTC 1 Jan 2011 (d) GOES West infrared satellite image of Storm 2 in (b) on 1200 UTC 24
1045 Dec 2011. 58



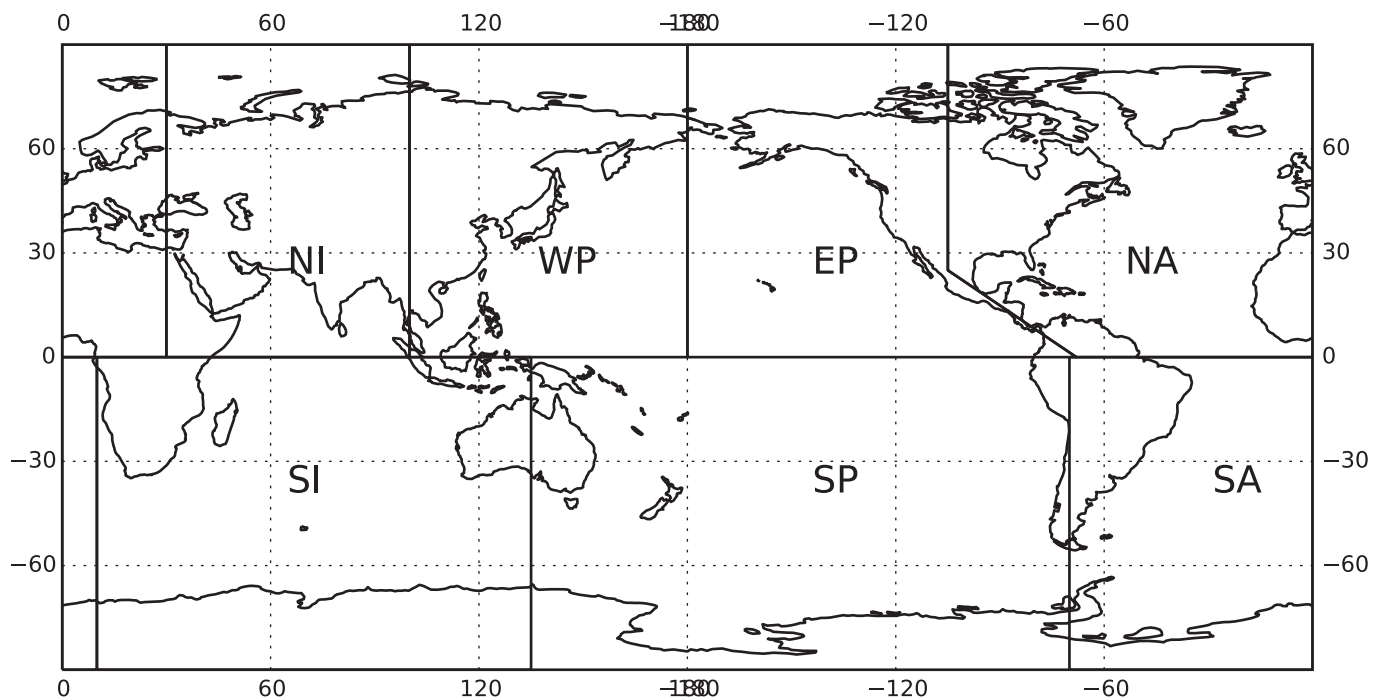
1046 FIG. 1. Distribution of mean separation distances (geodesic degrees, $1^{\circ} \simeq 111\text{km}$) between the reanalysis
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 1048 Results subsection) (a) NH, (b) SH, distribution of lifetimes (days) for the matched tracks (c) NH, (d) SH and
 1049 the distribution of latitudes at which the matched tracks attain the peak intensity based on the 10m winds (e)
 1050 NH, (f) SH.



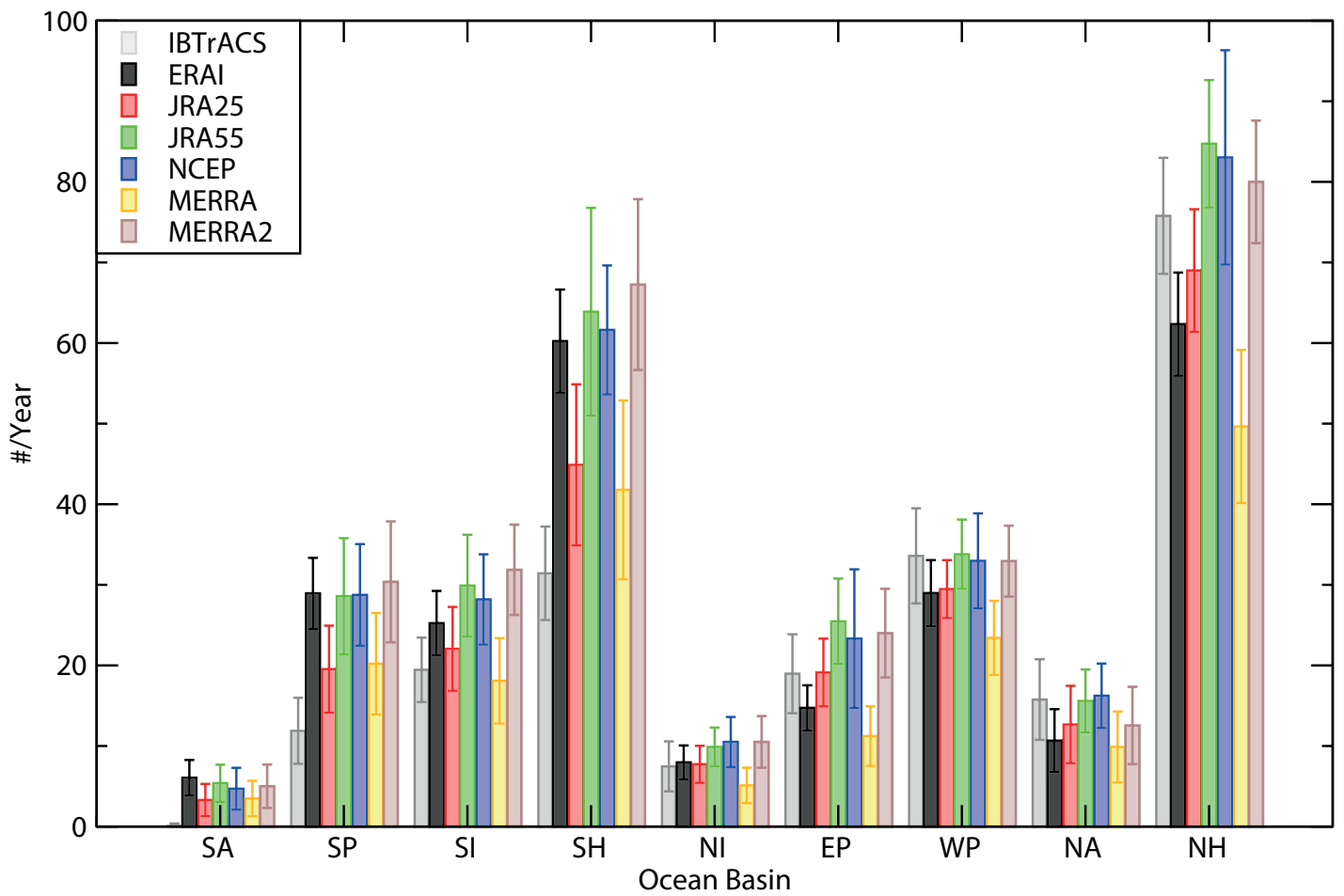
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 1054 (d) SH 10m wind speed (m s^{-1}), (e) NH 925hPa wind speed (m s^{-1}), (f) SH 925hPa wind speed (m s^{-1}).



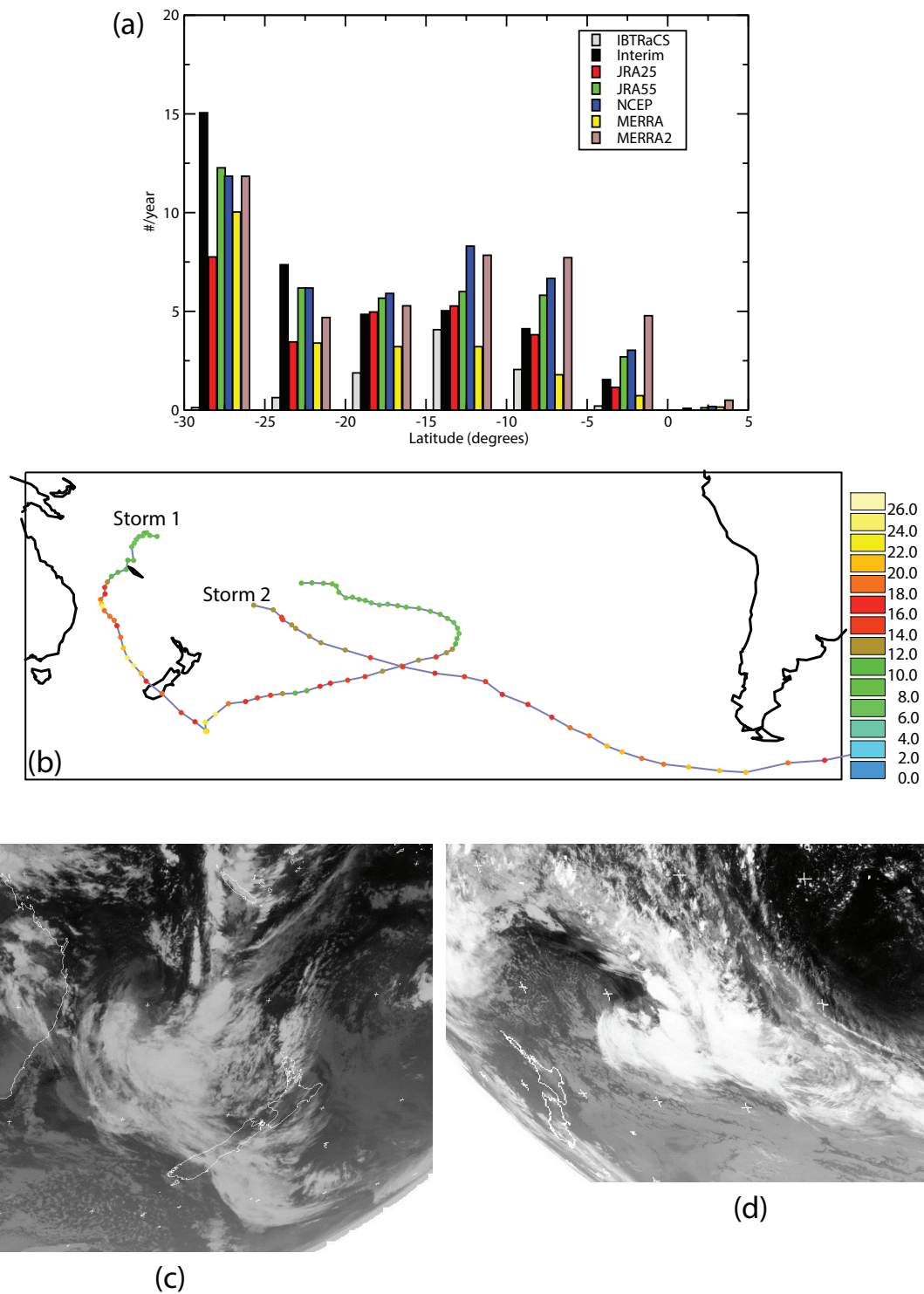
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 1056 imum winds for the reanalyses based on the direct matching method (c.f. Direct Matching Results subsection).
 1057 (a) NH 10m wind speed versus MSLP, (b) SH 10m wind speed versus MSLP, (c) NH radius of maximum winds
 1058 and (d) SH radius of maximum winds.



1059 FIG. 4. The seven basins used in this study, based on the IBTrACS definition. NI: North Indian, WP: West
 1060 Pacific, EP: East Pacific, NA: North Atlantic, SI: South Indian, SP: South Pacific, SA: South Atlantic.



1061 FIG. 5. The average number of TCs per year for each of the seven basins for IBTrACS and identified in
 1062 the reanalyses based on the objective detection method (c.f. Objective Identification subsection). Vertical lines
 1063 indicate the standard deviation. NI: North Indian, WP: West Pacific, EP: East Pacific, NA: North Atlantic, SI:
 1064 South Indian, SP: South Pacific, SA: South Atlantic.



1065 FIG. 6. (a) Latitude at which genesis occurs in the SH for the objectively identified TCs in the reanalyses that
 1066 do not match with IBTrACS (number per year), (b) examples of two tracks identified in the ERAI reanalysis with
 1067 no matching track in IBTrACS, coloured dots indicate 10m wind speeds (m s^{-1}), (c) MTSAT infrared satellite
 1068 image of Storm 1 in (b) on 1800 UTC 1 Jan 2011 (d) GOES West infrared satellite image of Storm 2 in (b) on
 1069 1200 UTC 24 Dec 2011.