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Impacts of orography on large-scale atmospheric circulation

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Some of the largest and most persistent circulation errors in global numerical weather prediction and climate models are attributable to the inadequate representation of the impacts of orography on the atmospheric flow. Existing parametrization approaches attempting to account for unresolved orographic processes, such as turbulent form drag, low-level flow blocking or mountain waves, have been successful to some extent. They capture the basic impacts of the unresolved orography on atmospheric circulation in a qualitatively correct way and have led to significant progress in both numerical weather prediction and climate modelling. These approaches, however, have apparent limitations and inadequacies due to poor observational evidence, insufficient fundamental knowledge and an ambiguous separation between resolved and unresolved orographic scales and between different orographic processes. Numerical weather prediction and climate modelling has advanced to a stage where these inadequacies have become critical and hamper progress by limiting predictive skill on a wide range of spatial and temporal scales. More physically based approaches are needed to quantify the relative importance of apparently disparate orographic processes and to account for their combined effects in a rational and accurate way in numerical models. We argue that, thanks to recent advances, significant progress can be made by combining theoretical approaches with observations, inverse modelling techniques and high-resolution and idealized numerical simulations.

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CHALLENGES IN MODELLING OROGRAPHIC IMPACTS ON ATMOSPHERIC CIRCULATION

Orography influences the atmospheric circulation through a wide range of processes and on a variety of spatial and temporal scales^{1,2} (Box 1, Fig. 1). Although orography can now be very accurately mapped using satellites at horizontal resolutions approaching tens of metres,^{3,4} the representation of its impacts on circulation in numerical weather prediction (NWP) and climate models remains notoriously difficult and several challenges must be overcome.

Unlike clouds or turbulence, which are completely unresolved even at global NWP resolutions of tens of km, the orographic spectrum is partially resolved by the model dynamics even at climate resolutions of hundreds of km. While certain orographic processes such as orographic gravity waves and low-level flow blocking become resolved at km-scale horizontal resolutions, turbulent form drag associated with small-scale hills remains unresolved even at O(~100 m) resolutions. This means that orography is within a 'grey zone' at resolutions currently used for global and even regional NWP and climate modelling.

Parametrizations accounting for unresolved orographic processes, therefore, need to accurately capture the exchange between resolved and unresolved orographic drag as horizontal resolutions are varied from km to hundreds of km. Figure 2 illustrates that this exchange is not necessarily correctly captured in two very different models used for operational global weather forecasts.^{5,6} Particularly in the Northern Hemisphere mid-latitudes, the change in the parametrized orographic torque is larger than

the change in the resolved orographic torque when the horizontal grid-spacing is refined from 150 to 16 km. As a result, the total orographic torque is not invariant with resolution, as would be expected based on angular momentum budget considerations (Box 2). This suggests that at climate model resolutions (~O(100 km)) the parametrized orographic torque is not representative of the mountain torque that is explicitly resolved at higher resolutions (~O(10 km)).

The formulations of orographic drag parameterizations also need to encompass processes acting across multiple scales and flow regimes (Box 1). By necessity, these parameterizations have so far mostly relied on simplified theoretical or empirical arguments targeted towards specific orographic processes in isolation. For gravity wave drag, the parameterizations are typically based on analytical solutions for simple flows over isolated mountains (dating back nearly 80 years⁷) using linear steady-state approximations to the equations of motion.^{8–10} For low-level flow blocking, they rely on simple empirical concepts for bluff-body dynamics.¹¹ Numerical simulations¹² and laboratory tank experiments¹³ of flow over idealized mountains have helped to refine these approaches, but evidence of how these concepts are applicable to more realistic and complex topography has only recently begun to emerge.^{14,15}

It is also not clear how to properly account for the combined effects of the different orographic drag processes, which have thus far been dealt with in rather ad hoc ways. This is due both to missing fundamental knowledge of how these processes interact and the fact that orographic drag processes, and other drag

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Box 1: Orographic effects on atmospheric circulation

The interaction between orography and the atmosphere is a fascinating fluid dynamics problem. Topographic features such as mountains, hills or valleys exert a force on the atmosphere and induce a rich variety of complex disturbances.⁶¹ These disturbances exert drag (in the sense of friction or resistance) on the flow and affect its evolution. The circulation response to orography depends both on the obstacle size and shape and the atmospheric conditions. For example, flow splitting occurs around small-scale orography in stably stratified conditions when the air flow has insufficient kinetic energy to overcome the potential energy barrier of the mountain. The divergence of the flow around the mountain creates elongated and persistent wakes downstream that, when visualized through their impact on the clouds (Fig. 1a), closely resemble water flowing over rocks in a brook. Air flow over a mountain barrier in a predominantly stably stratified environment can generate internal gravity waves that can propagate horizontally and vertically (lee-waves or vertically propagating mountain waves).⁶² These waves may affect the flow either in the vicinity of the mountain (Fig. 1b) or, in the case of vertically propagating waves, remotely, by depositing momentum away from the source region when they break. Wave breaking can also induce clear air turbulence at high altitudes with damaging consequences for aircraft.⁶³ At lower altitudes, wave breaking in the troposphere can result in an acceleration of the flow across the leeward side of the mountain, producing severe downslope windstorms.⁶⁴ (Fig. 1c). At the very small scales, terrain undulations or hills that are unable to generate gravity waves (below roughly 5 km horizontal scale) still exert orographic drag (referred to as turbulent orographic form drag) resulting in a deceleration of the flow within the boundary layer. While some of the aforementioned effects are rather small scale, their cumulative impact can project onto the large-scale circulation. Breaking of orographically generated gravity waves can, for example, influence the variability of certain aspects of the large-scale circulation, such as the breakdown of the stratospheric polar vortex.⁶⁵ Significant orographic barriers also exert a direct forcing on the planetary scales, generating Rossby waves which are affected by the rotation of the Earth and act to shape the zonally asymmetric circulation in both hemispheres.² These quasi-stationary waves and the associated flow regimes impact many aspects of regional climate. The mechanism of this large-scale wave generation can, and has been, understood through simple models. For example, the Charney–Eliassen⁶⁶ model, which approximates the atmosphere as barotropic, steady and in geostrophic and hydrostatic balance, closely reproduces the observed stationary wave pattern over the Northern Hemisphere during winter when forced with smoothed orography (Fig. 1d). This implies that even simple arguments can be used to understand the impacts of topographically forced waves on the large-scale temporal and spatial variability of the atmosphere. Models of this type provide not only a simplified understanding of the processes but also a means of determining which approximations are valid at various scales.

processes, are not well constrained by observations. This leaves their representation in models, in terms of both magnitude and effects on circulation, exposed to compensating errors (Box 2).

WHY DOES IT MATTER?

Despite these challenges, existing orographic drag parametrizations capture the basic effects of unresolved orographic processes^{16–18} and have been instrumental in advancing weather forecasting and increasing the realism of the representation of the circulation in climate models over the past decades. The first orographic gravity wave drag parametrizations, introduced in operational NWP systems and climate models in the early 1980s,^{19,20} helped alleviate the excessively strong westerly flow in the upper troposphere, improved the representation of surface pressure in the northern hemisphere (NH) winter and reduced the stratospheric cold pole bias by >10 K. A parametrization for the low-level blocking of the flow due to unresolved mountains¹¹ was also widely adopted and led to significant improvements in the predictive skill of medium-range weather forecasts. For example, in the Canadian global NWP system it led to a far bigger improvement in skill during the NH winter than any other individual change made in the past 15 years.²¹ Accounting for the turbulent orographic form drag from small-scale (<5 km horizontal) hills^{17,22} also greatly contributed to increases in predictive skill of weather forecasts. Figure 4 illustrates that the predictive skill of 10-day forecasts dramatically decreases by the equivalent of 1 day during winter in the NH if this parametrization is removed from a global NWP system. This is equivalent to the weather forecast skill

gained over the past decade through continuous improvements to the global NWP systems.²³

Given this circulation sensitivity to parametrized orographic drag, uncertainties in its representation (Fig. 3) were unsurprisingly found to impact the skill of weather forecasts on daily to seasonal timescales.²⁴ Since the parameterized orographic drag contributes substantially to the momentum budget within climate models,²⁵ its representation and uncertainties related to it have also been found to be important for climate model fidelity^{26,27} and for regional aspects of climate change.^{25,28}

The circulation sensitivity to parametrized orographic drag and the large uncertainty related to these processes highlighted in the past few years motivates the need to make advances in their representation in Earth system models. Reducing these uncertainties would allow a leap forward in predictive skill across the entire range of temporal and spatial scales. While some aspects of orographic drag parametrizations may become obsolete with resolution upgrades in NWP in the next few decades, most aspects will remain relevant for the foreseeable future. For example, form drag and flow separation across small hills only become resolved at O(100 m) and will hence still need to be parametrized in models used for predictions from hours to decades. Low-level flow blocking or gravity wave effects become resolved at km-scale, but restrictions on resolution in ensemble forecast systems and climate projections will mean that these effects will need to be parametrized in models used for these applications for many years to come.

WHAT IS NEEDED TO MAKE PROGRESS?

We argue that, thanks to recent advances in numerical models, computing techniques, observational analysis and retrieval and fluid dynamics theory, there is now scope to significantly improve our understanding of orographic drag processes and their representation in models by pursuing the three avenues detailed below.

Observational constraints and inverse modelling

It is currently impossible to directly observe the three-dimensional distribution of momentum flux induced by orographic features at global scale or even over extended mountain ranges. However, recent developments open exciting opportunities for constraining at least one of the orographic processes: mountain waves. Recent field campaigns, such as the DEEPWAVE²⁹ field experiment in New Zealand, provide in situ aircraft measurements of wave properties (including momentum fluxes). Satellite observations offer a wealth of information for constraining the representation of gravity waves,³⁰ and significant recent advances in methods of deriving gravity wave characteristics (wavelength, propagation, amplitude) from such observations^{31,32} offer new possibilities for using this information to evaluate the representation of orographic gravity waves in models. Although difficulties still exist in comparing these observations with models, progress has also been made in diagnostic techniques allowing such comparisons.^{33,34} Spaceborne or ground-based lidar observations are also increasingly being used to explore gravity wave characteristics and their representation in models.^{35,36}

Using analyses or reanalyses (our best guess of the atmospheric state) for examining systematic short-range forecast errors has proven fruitful for attributing errors due to orographic drag,^{15,37–39} since the errors remain fixed over the orography itself at short lead times. While these techniques have been and are routinely utilized to evaluate NWP models, there is also value in applying them towards understanding model errors in lower-resolution climate models⁴⁰ and such studies could be done more routinely.

The development of sophisticated mathematical and statistical techniques in the past years, especially in the context of data

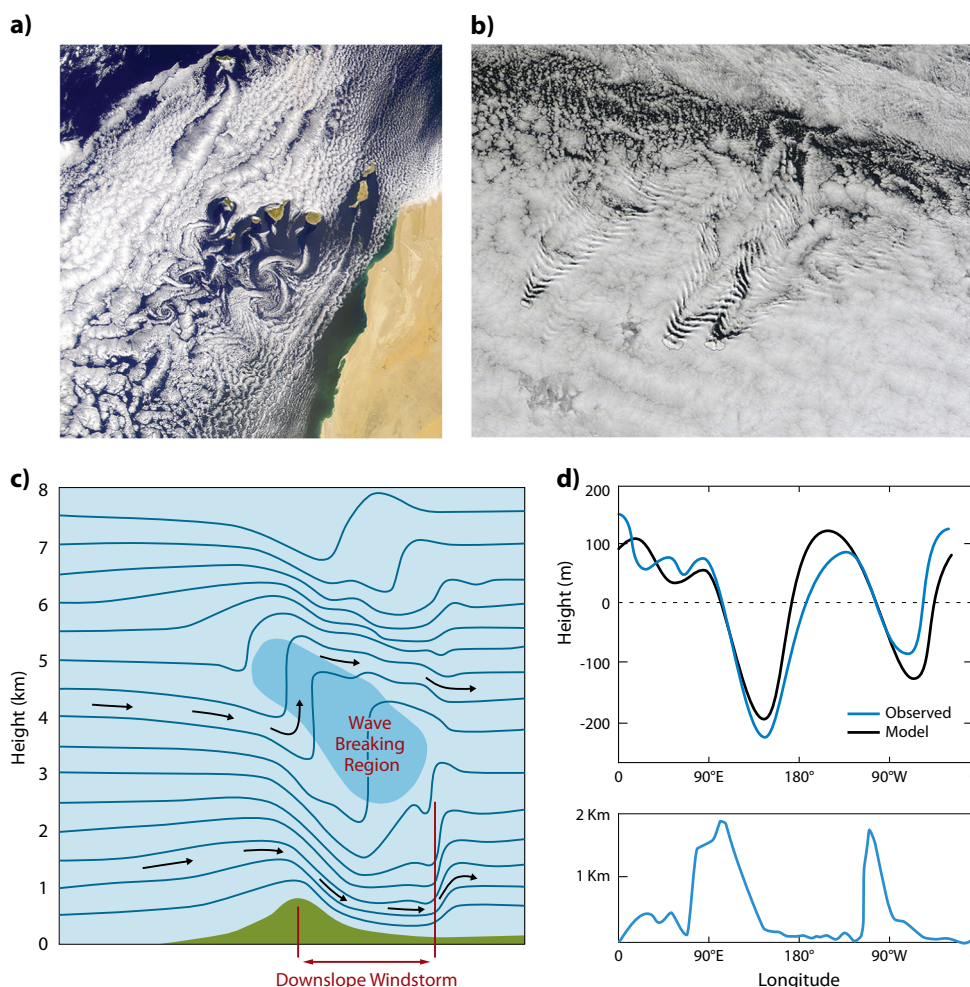


Fig. 1 Orographic effects on atmospheric circulation. **a** Island wake effects revealed by cloud-free regions in the lee of the Canary Islands, and von Karman vortices highlighted by the cloud spirals further downstream (Photo from the SeaWiFS Project, NASA/Goddard Space Flight Center, and ORBIMAGE, on 24 April 2000 used with permission from Visible Earth NASA); **b** Classical 'v' shape distinctive banded lee-wave pattern is visible in the clouds downstream of the Crozet Island over the southern Indian Ocean. The lifting and cooling of air on the crests of these waves lead to cloud formation, while sinking and heating leave the troughs cloud-free. (Photo from the Moderate Resolution Imaging Spectroradiometer (MODIS) on NASA's Terra satellite on 9 April 2014 used with permission from Visible Earth NASA); **c** Downslope windstorms can lead to severe accelerations of the near-surface winds on the leeward side of the mountain, when the vertical propagation of waves is prevented by an elevated inversion-like region created by the breaking of large-amplitude waves close to the mountain top. Adapted from Whiteman.⁷⁶ **d** Longitudinal variation of the disturbance of the geopotential height in the Charney–Eliassen model compared to the observed 500-hPa geopotential height perturbations at 45°N in January (top) and smoothed profile of topography at 45°N used in the Charney–Eliassen model (bottom). (Redrawn after Holton,⁷⁷ with permission from Elsevier)

assimilation and ensemble forecasting, also offers new ways of constraining uncertain parameters within, and errors associated with, orographic drag parametrizations. For example, inverse modelling techniques employing data assimilation,^{41,42} perturbed parameter ensembles⁴³ and statistical emulators,^{44,45} can be used to recover information on uncertain model parameters by fitting the model forecast to observed data, such as winds, temperature or pressure. Some of these techniques have been applied to orographic drag processes in simplified models.⁴² Applying them to NWP and climate models seems very promising. To optimize a parametrized process through data assimilation, for example, it is necessary to have one or more parameters that are not well constrained. At the same time, it is also necessary to have an observation that is informative about that process or parameter. Surface stress is a good candidate because surface pressure responds to surface stress,²⁰ and surface pressure is a well-observed quantity over land. The match of surface pressure to observations could be optimized by adjusting surface stress at each model grid point. This can be achieved, for example, through

an 'augmented control variable' approach, which is now available in complex global data assimilation systems.⁴⁶

High-resolution simulations

With increasing computational power and improvements in numerical methods, comprehensive high-resolution simulations with spatial resolutions that can realistically⁴⁷ capture the fluid dynamics over complex mountain terrain at scales relevant for improving orographic drag parametrizations are becoming increasingly available. Such simulations can be considered as high-resolution virtual laboratories for exploring orographic drag processes. The added benefit of using such simulations over realistic complex topography initialised from our best guess of the observed atmospheric state, as opposed to idealised ones, is that they allow us to quantify (in real terms) the contribution of the parametrized orography to the model error.¹⁵

Studies with comprehensive high-resolution models have already been successfully used to demonstrate the improved representation

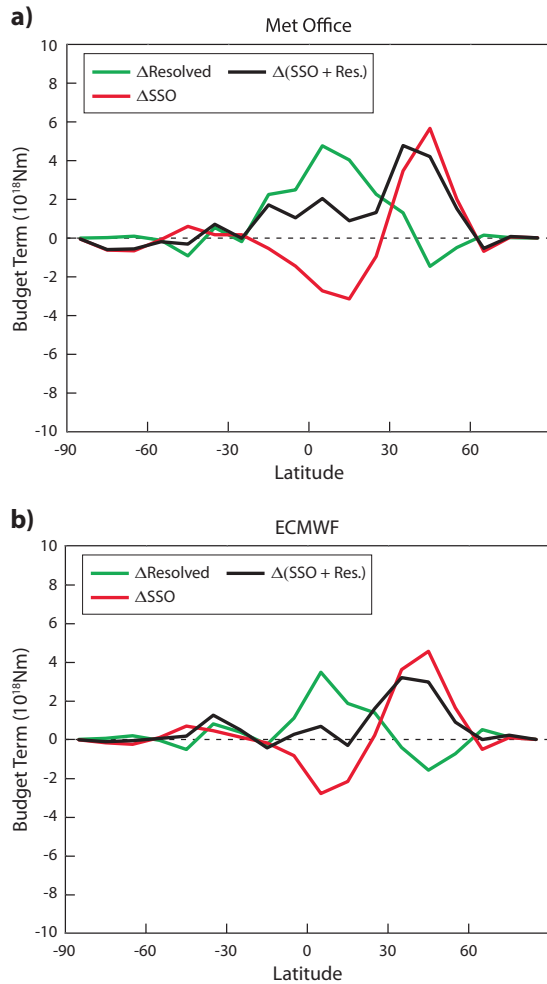


Fig. 2 Difference in the resolved (green), parametrized (red) and total (resolved+parametrized, black) orographic torques between horizontal resolutions of 16 and 150 km. The torques are integrated over 10° latitude bands and averaged over a set of 24 h forecasts performed during December 2015 with models used for operational weather prediction at **a** the UK Met Office and **b** the European Centre for Medium-Range Weather Forecasts (ECMWF). The sign of the differences is such that the resolved westward torques between 30 and 60°N (which act as a drag on the westerly flow) are larger at 16 km compared to 150 km resolution, while the parametrized torques are smaller. The reduction in parametrized orographic torque greatly exceeds the increase in resolved orographic torque, leading to a smaller total orographic torque at higher resolution. The differences in the total (resolved+parametrized) orographic torques are, depending on latitude, between 20% and 40% of the total orographic torque in the UK Met Office model at 150 km resolution. See ref. ⁶ for further details of the methodology

of flow around orography at increased resolutions on mesoscales.⁴⁸ Km-scale simulations at which low-level blocking and orographic gravity waves become resolved have also been used to demonstrate the potential of existing parametrizations to capture the time mean and the variability of the resolved orographic drag and to explore their limitations for small islands with real terrain.^{14,18} We are now entering an era where such simulations become feasible over extensive and complex mountain ranges such as the Himalayas,¹⁵ the Rockies and even globally.⁴⁹

As an example, Fig. 5 compares the monthly averaged zonal momentum flux by resolved gravity waves in a high-resolution (2 km) regional simulation over the Rockies with the parametrized momentum flux in an equivalent low-resolution (32 km)

Box 2: The atmospheric momentum budget

The zonally averaged vertically integrated absolute angular momentum budget equation provides a useful framework for understanding how surface stress (often also referred to as surface drag) relates to the large-scale circulation^{67,68}:

$$\frac{\partial \langle \rho m \rangle}{\partial t} = - \left[\rho_0 \frac{\partial h}{\partial \lambda} \right] - \frac{1}{r \cos \phi} \frac{\partial}{\partial \phi} \langle \rho m v \cos \phi \rangle - [\tau_x r \cos \phi] \quad (1)$$

where $m = (u + \Omega r \cos \phi) r \cos \phi$ is the absolute atmospheric angular momentum, ρ is the atmospheric density, v is the meridional wind, u is the zonal wind, r is the radius of the Earth, ϕ is latitude, λ is longitude, ρ_0 is the pressure at the surface, h is the height of the surface, and τ_x is the zonal mean surface stress from unresolved (or parametrized) drag processes acting at scales smaller than those resolved by the model dynamics. The operators $\langle \dots \rangle$ and $[\dots]$ denote the vertical and zonal integral, respectively. This formulation reflects the fact that local drag forces are actually stress divergences, i.e. momentum exchanges, hence the vertical integral of the momentum equation leaves only the exchanges between the atmosphere and the surface, represented in the surface stress. This surface stress is then the net effect of all drag processes induced by interaction with the surface, wherever they occur within the atmosphere. The angular momentum equation illustrates that the rate of change of absolute angular momentum of the atmosphere (LHS of Eq. (1)) is balanced by, from left to right, the resolved mountain torque, meridional convergence of angular momentum in the free atmosphere, and surface torques from parametrized processes such as orographic drag associated with unresolved orographic features (e.g. turbulent orographic form drag, blocking of the flow at low levels and orographic gravity waves) and other planetary boundary layer processes. In the steady-state limit, the dominant balance is between the convergence of the angular momentum flux in the free atmosphere and the total surface torque due to resolved mountains and unresolved processes.

The pressure forces (first term on the RHS of Eq. (1)) from large-scale mountain chains are notoriously difficult and costly to measure, with such measurements being limited to only a few transects across particular mountain ranges.⁶⁹ The convergence of the angular momentum flux, on the other hand, is well observed on large scales, as is independently verified by observations of the length of day.⁷⁰ The balance implied by Eq. (1) means that this also provides a constraint on the total zonally integrated surface torque.⁷¹ However, the horizontal distribution of the total surface torque, and its partition between the different processes (resolved and unresolved), is not well constrained. As a result, different models partition the total surface torque in different ways. This partitioning is dependent on: the scales resolved by the model dynamics; details of how the orographic fields are derived; the formulation and combination of the parametrization schemes used; and the tuning of uncertain parameters within these schemes. These parameter values often depend on subjective choices aimed at improving specific NWP skill scores⁷² or calibrating the model climate.⁷³ State-of-the-art global NWP and climate models, therefore, differ in their resolved orography,⁷⁴ their total parametrized surface stress and their partitioning between the various processes⁷⁵ (Fig. 3). Often, parametrized boundary layer turbulent stress and orographic stress are traded against each other. The UK Met Office Unified Model has, for example, much more low-level blocking and gravity wave drag and considerably less turbulent friction and turbulent orographic form drag compared with the Integrated Forecasting System of the European Centre for Medium-Range Weather Forecasts, but the total surface stress only differs by ~10–15% in the NH mid-latitudes. This is because the models are tuned to respect the observational constraint (the angular momentum flux convergence), thus giving rise to the issue of compensating errors.

simulation. The resolved momentum flux has been derived from the 2-km model-simulated velocity fields, which explicitly resolve the orographic gravity waves. In this case, it appears that the parametrized momentum flux in the low-resolution model provides a reasonable representation of the well-resolved ‘truth’.

Spectral filtering methods⁵⁰ used to isolate the gravity wave velocity perturbations in order to compute the momentum fluxes, as are employed in Fig. 5, can be used to determine the spectral contributions to momentum fluxes in high-resolution simulations. Since much of the gravity wave parametrization theory is based on linear wave solutions, this allows for a direct comparison of the scale dependence of the orographic drag deduced from the models with the analytical one, as had been done previously with observed momentum fluxes.⁵¹ Pursuing such an approach in the future is highly relevant to the parametrization problem, since it allows both to determine the dominant scales contributing towards the resolved momentum fluxes (and therefore, which scales the parametrizations need to account for) and to evaluate the validity of the theory.

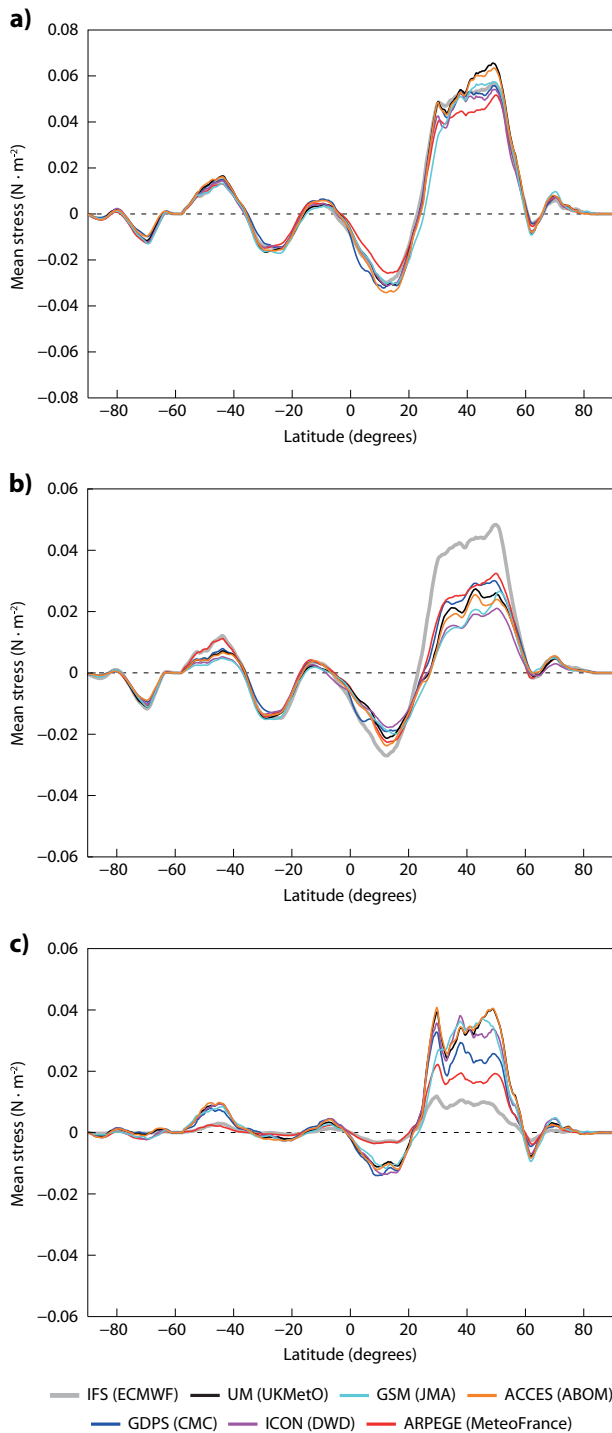


Fig. 3 The atmospheric momentum budget. Zonal averages (land points only) of **a** total parameterized surface stress, and contributions from **b** the turbulent friction and turbulent orographic form drag, and **c** low-level blocking and orographic gravity wave drag, from the results submitted by different groups to the Working Group for Numerical Experimentation of the World Meteorological Organization Drag project, for short range forecasts for January 2012

Future analysis of km-scale experiments should be targeted towards understanding: how orographic drag is distributed in the horizontal and vertical domain; how this distribution depends on the flow characteristics, namely the thermal stratification, the wind

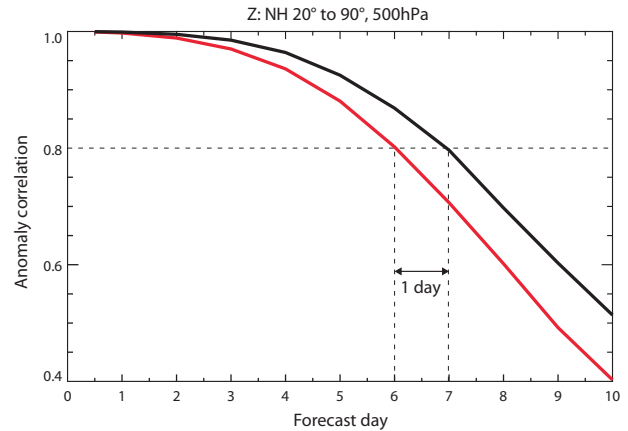


Fig. 4 Almost 1 day of skill is lost when the turbulent orographic form drag parametrization¹⁷ is switched off in global 10-day weather forecasts performed with the Integrated Forecasting System of the European Centre for Medium-Range Weather Forecasts: in the control experiment (black), forecasts 7 days ahead have an accuracy of 80% in terms of anomaly correlation of the geopotential height at 500 hPa, while in the experiment where the form drag scheme was switched off (red) this level of accuracy is lost after 6 days. In other words, in the new experiment the degree of accuracy of the forecasts only matches the degree of accuracy of numerical weather prediction systems of 10 years ago.²³ The simulations have been performed at a horizontal resolution of 32 km, for the period December–February 2016–2017

speed and wind direction over mountain ranges of different sizes (e.g. Pyrenees, Alps, Himalayas, Rockies, etc.) as well as isolated mountains (e.g. Kilimanjaro, Fuji); and which processes contribute to the total orographic drag. Particular emphasis should be given to processes that are currently not accounted for in parametrizations, such as lee-wave drag and transient wave–mean flow interaction, so that the consequences of their omission may be evaluated.

Simulations at even higher resolution ($O(100\text{ m})$), which can start to resolve turbulent form drag associated with orographic features with scales $<5\text{ km}$, are also becoming available at continental scale.⁵² These could be used in the future to help constrain the partitioning between form drag and the other orographic drag processes.

Better theoretical understanding

The focus on orographic drag processes in recent years has highlighted the gaps in our theoretical understanding of these processes. One of the key gaps lies in the partitioning between different drag processes, illustrated by the inter-model spread in surface stress shown in Fig. 3. This discrepancy may be due to missing processes not accounted for and/or simplifications made in the parametrizations. Theories for more complex parametrizations, for example, accounting for drag effects due to vertical windshear,⁵³ downslope windstorms⁵⁴ and transient gravity wave–mean flow interaction⁵⁵ do exist. However, simple approximations are maintained in global models due to the considerable leap that needs to be made from idealized theory to practical implementation of a parameterization scheme suitable for weather and climate models. There are also several processes that are not well understood theoretically and are not incorporated into existing parametrizations, such as mean flow time dependence⁵⁶ and the interactions of multiple scales of topography in a stratified flow.⁵⁷

A better theoretical understanding of the large-scale circulation response to different forms of orographic drag, such as the non-

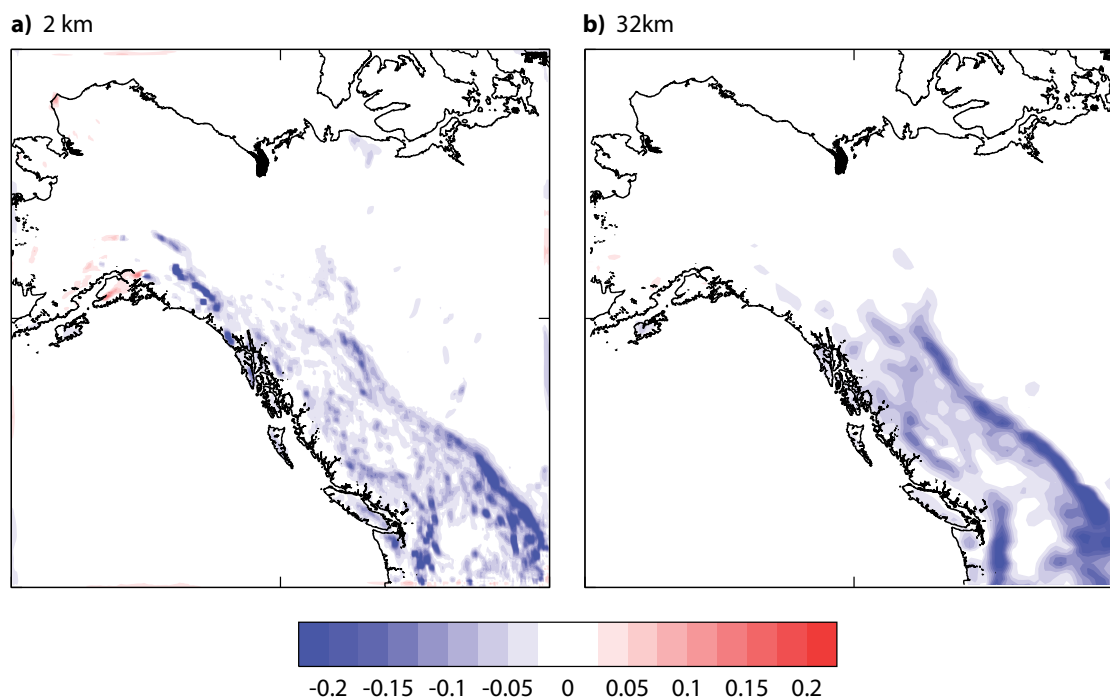


Fig. 5 The mean zonal component of gravity wave momentum flux (N m^{-2}) at 7 km above sea level as predicted by 1-month-long simulations (December 2015) of flow across the Rockies with the Met Office Unified Model. The right-hand panel shows the parametrized flux produced by a simulation with a coarse (32 km) grid. The left-hand panel shows the simulated flux resolved on a much finer (2 km) grid, after it has been coarse-grained onto the 32 km grid. Comparison of the two momentum flux fields allows one to determine the realism of the parametrization scheme

local response to a torque,^{58,59} is also required. If these responses would be understood, one could use the indirect or far-field response to a certain process to obtain a ‘fingerprint’ of that process and to evaluate its accuracy. An example is the way in which the amplitude and period of the Quasi-Biennial Oscillation is used as a constraint on the non-orographic gravity wave drag.⁶⁰

Progress at both the process level and in terms of circulation impacts can be achieved by combining a hierarchical modelling approach with our existing theoretical knowledge. In addition to the high-resolution simulations with realistic terrain, simulations that start from well-known idealized mountain shapes in idealized flow, for which we have analytical solutions, and gradually increase in complexity towards real-world mountain ranges in real flow can showcase the limitations of established concepts and provide a quantitative basis for new understanding. Idealized general circulation experiments, in which drag is prescribed with different geographical or vertical distributions, could also be used to develop observable large-scale fingerprints of different processes contributing to drag as a means of constraining the partitioning of drag. These will help to bridge the gap between the simplifications made in current parametrizations, the available theory and more advanced parametrizations.

A PROMISING FUTURE FOR OROGRAPHIC DRAG

Key challenges in representing orographic drag processes include the ‘grey-zone’ problem, which encompasses the unclear separation between resolved and unresolved orographic scales, the overlap between different processes and the difficulties in directly observing orographic drag. Overcoming these issues will be instrumental in the development of a future generation of increasingly accurate weather and climate models. Rapid progress is within reach by making combined use of theoretical approaches, emerging observational constraints and inverse

modelling and high-resolution simulations along the avenues exposed within this perspective.

DATA AVAILABILITY

The data used in Figs 2–5 are available at the following: <https://doi.org/10.5281/zenodo.2552060>. Original unprocessed data from the numerical experimentation done to produce Figs 2, 4 and 5 are available upon request from the first author.

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AUTHOR CONTRIBUTIONS

I.S., A.v.N., T.G.S. and S.V. conceived and wrote a first draft of the manuscript, based on discussions at the ECMWF drag workshop in September 2016 that was attended by all authors. A.v.N. provided Fig. 2, A.Z. provided Fig. 3, I.S. provided Fig. 4, S.V. provided Fig. 5. N.M. wrote the abstract. All other authors provided comments, suggested edits and contributed additional expertise in proposing the three avenues for making progress. All authors provided comments and suggested edits for the initially submitted draft and the changes made during the review process and approved the final draft.

ADDITIONAL INFORMATION

Competing interests: The authors declare no competing interests.

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