

# *The EUSTACE project: delivering global, daily information on surface air temperature*

Article

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119

120 Capsule

121

122 The main goals and activities of the EUSTACE project are discussed along with some key

123 results, including a global, multi-decadal daily air temperature record from satellite and *in*

124 *situ* measurements.

125 Abstract

126

127 Day-to-day variations in surface air temperature affect society in many ways, but daily  
128 surface air temperature measurements are not available everywhere. Therefore, a global  
129 daily picture cannot be achieved with measurements made *in situ* alone and needs to  
130 incorporate estimates from satellite retrievals.

131 This article presents the science developed in the EU Horizon 2020-funded EUSTACE project  
132 (2015-2019, <https://www.eustaceproject.org>) to produce global and European, multi-  
133 decadal ensembles of daily analyses of surface air temperature complementary to those  
134 from dynamical reanalyses, integrating different ground-based and satellite-borne data  
135 types. Relationships between surface air temperature measurements and satellite-based  
136 estimates of surface skin temperature over all surfaces of Earth (land, ocean, ice and lakes)  
137 are quantified. Information contained in the satellite retrievals then helps to estimate air  
138 temperature and create global fields in the past, using statistical models of how surface air  
139 temperature varies in a connected way from place to place; this needs efficient statistical  
140 analysis methods to cope with the considerable data volumes. Daily fields are presented as  
141 ensembles to enable propagation of uncertainties through applications. Estimated  
142 temperatures and their uncertainties are evaluated against independent measurements and  
143 other surface temperature data sets.

144 Achievements in the EUSTACE project have also included fundamental preparatory work  
145 useful to others, for example: gathering user requirements; identifying inhomogeneities in  
146 daily surface air temperature measurement series from weather stations; carefully  
147 quantifying uncertainties in satellite skin and air temperature estimates; exploring the

- 148 interaction between air temperature and lakes; developing statistical models relevant to
- 149 non-Gaussian variables; and methods for efficient computation.

150 Body text

151 EU Surface Temperature for All Corners of Earth (EUSTACE,  
152 <https://www.eustaceproject.org>) is a 4-yr research project funded by the European Union  
153 Horizon 2020 research and innovation programme (EU H2020; Grant Agreement 640171;  
154 see Appendix A for a list of the Consortium's institutions) that started on 1 January 2015.  
155 EUSTACE has used temperature estimates from satellites to boost the amount of  
156 information available beyond that provided by weather stations and ships to help to  
157 construct a prototype global, multi-decadal daily air temperature record presented on a  
158 0.25° latitude by 0.25° longitude grid.

159

160 Near-surface air temperature (typically measured at a height of about 2 m above ground  
161 level at meteorological stations) is a fundamental quantity for many of the activities  
162 undertaken in climate science and in many of the societal concerns that climate services aim  
163 to support; it is something that we all experience directly in our day-to-day lives. Near-  
164 surface air temperature has been measured almost continuously in some places and across  
165 the global oceans by ships for well over a century. Designated as an Essential Climate  
166 Variable (ECV), these records allow for the construction of a useful climate data record  
167 (CDR) in those places for the period covered. Globally, however, there a number of locations  
168 where either access to the measurements is not possible, or no air temperature records  
169 exist. As well as long records of direct measurements of near-surface air temperature, we  
170 have information from satellite retrievals (i.e. remotely-sensed, indirect estimates) of  
171 temperature. However, satellite retrievals tend not to pertain to the air temperature that  
172 we experience directly, but either to an average temperature of a higher layer in the

173 atmosphere or to the skin temperature of the surface of the Earth. These quantities are  
174 related to near-surface air temperature, more or less tightly depending on the type of  
175 surface and the surface-lower-atmosphere interactions. Therefore, it is possible to use  
176 satellite-derived temperatures together with near-surface air temperature measurements  
177 to create a more complete climate data record of air temperature. Thus, EUSTACE created a  
178 prototype global climate data record of near-surface air temperature for every day since  
179 January 1850 using both direct measurements of air temperature and estimates of it based  
180 on satellite skin temperature retrievals.

181

182 Near-surface air temperature products provide valuable information for a range of activities,  
183 from the monitoring of current conditions (e.g. Sánchez-Lugo et al. 2019) to the assessment  
184 of past variability (e.g. Osborn et al. 2017) to their use in seasonal-to-decadal forecasting  
185 (e.g. Kushnir et al. 2019), climate model evaluation (e.g. Walters et al. 2019), detection and  
186 attribution of climate change (e.g. Jones and Kennedy 2017), Intergovernmental Panel on  
187 Climate Change Assessments (e.g. Hartmann et al. 2013), agricultural modelling (e.g.  
188 Weedon et al. 2011), health modelling (e.g. Xu et al. 2019) and other downstream uses.  
189 Such a daily surface air temperature product could form part of the future operational  
190 monitoring system for surface air temperature over the polar regions, over Africa and South  
191 America. EUSTACE has already enabled monitoring of lake surface water temperature to be  
192 included in the annual State of the Global Climate reports (for the years 2015, 2016, 2017  
193 and 2018; Woolway et al., 2016, 2017a and 2018; Carrea et al., 2019). EUSTACE products are  
194 complementary to products from dynamical reanalyses (e.g. Buizza et al. (2018)) with much  
195 of the work dedicated to the preparation of input surface temperature observations, for

196 which EUSTACE has performed thorough uncertainty analyses, which were previously  
197 lacking.

198

199 Dynamical reanalyses combine historical and recent observations with numerical weather  
200 prediction models to produce dynamically-consistent reconstructions of past weather and  
201 climate. These reanalyses require observational data with well-characterised uncertainties.  
202 The new, validated, estimates of uncertainty in satellite surface skin temperature  
203 observations developed by EUSTACE are of benefit to them. EUSTACE products also provide  
204 an alternative source of near-surface air temperature data that is independent from  
205 numerical weather prediction models and extends further back in time than most dynamical  
206 reanalyses.

207

208 Results from scientific projects are often not produced in a format that can be used easily by  
209 others; in general, processing or translation is needed. Two-way interaction with potential  
210 users from the start of a project helps to increase the relevance and usability of products to  
211 various potential user groups. EUSTACE collected information on user requirements in several  
212 ways, via: user consultation workshops; questionnaires and interviews; a literature review on  
213 user requirements (Bessembinder et al. 2016; Bessembinder 2017, including the results from  
214 a large number of national and EU projects); testing of example mock-up datasets; and  
215 describing specific use cases with “trail blazer” users.

216

217 These activities resulted in greater insight into how climate data are used, data format  
218 preferences, and which variables are needed (i.e. not just daily mean temperature, but also  
219 minimum and maximum temperature), amongst other things. We used many of the user  
220 requirements collected to design the EUSTACE data file structure and the user guides; for  
221 example, a quick start guide is provided as part of the product user guide, together with  
222 example use cases.

223

224 While many of the ideas used within EUSTACE have been trialled elsewhere for individual  
225 regions (e.g. Cristóbal et al. (2008)), or for different time scales (e.g. Kilibarda et al. (2014)),  
226 EUSTACE has brought them together for the first time to create global, multi-decadal daily  
227 products. EUSTACE has performed an integrating function, bringing together products and  
228 expertise from a wide range of European, national and international initiatives. EUSTACE has  
229 also followed much of the road map of “recommended steps towards meeting societal  
230 needs for surface temperature understanding and information” set out previously in the  
231 EarthTemp Network Community Paper (Merchant et al. 2013). In particular, EUSTACE has  
232 made progress in seven out of the ten broad aims identified therein:

- 233 • develop more integrated, collaborative approaches to observing and understanding  
234 Earth’s various surface temperatures;
- 235 • build understanding of the relationships between different surface temperatures,  
236 where presently inadequate;
- 237 • make surface temperature datasets easier to obtain and exploit for a wider  
238 constituency of users;



- 239 • consistently provide realistic uncertainty information with surface temperature  
240 datasets;
- 241 • communicate differences and complementarities of different types of surface  
242 temperature datasets in readily understood terms;
- 243 • rescue, curate and make available valuable surface temperature data that are  
244 presently inaccessible; and
- 245 • build capacities to accelerate progress in the accuracy and usability of surface  
246 temperature datasets.

247

248 Computer code has been developed both to estimate air temperature from satellite data  
249 and to create daily maps of mean air temperature; this code has been publicly released  
250 (Rayner 2019). Information contained in the satellite retrievals helps to create more-  
251 complete fields in the past, via statistical models of how surface air temperature varies in a  
252 connected way from place to place. As the data volumes involved are considerable, the  
253 EUSTACE partnership included statisticians and computer scientists, enabling the  
254 development of efficient analysis methods. As a result, EUSTACE has been able to  
255 demonstrate that these methods can be built into a fully functional processing system, with  
256 research-level maturity (EUMETSAT, 2014) which exploits the features of modern high  
257 performance computing resources to deliver the more-complete datasets described below.  
258 This system could be used in future to update some of the EUSTACE data sets described  
259 here to enable their use in climate monitoring.

260

261 The datasets that are currently commonly used to monitor surface temperatures globally  
262 are constructed as a combination of air temperature observations over land and sea surface  
263 temperature observations over ocean. The current versions of the most widely used global  
264 near-surface temperature datasets, HadCRUT4 (Morice et al., 2012), NOAAGlobalTemp  
265 (Smith et al., 2008; Vose et al., 2012) and GISTEMP (Hansen, 2010), extend from the mid-  
266 19th century to present and are derived from *in situ* observations only; temperature  
267 retrievals from satellites are not used in their construction. These global temperature  
268 datasets are presented at monthly resolution because summaries of monthly average  
269 temperatures are more commonly available for individual meteorological stations and cover  
270 a greater region of the Earth than daily or sub-daily summaries in the 19th century and early  
271 20th century. The density distribution of available *in situ* temperature observations limits  
272 the spatial resolution of these products. For example, HadCRUT4 is provided as monthly  
273 fields on an equi-angle latitude-longitude grid at 5° resolution.

274

275 Surface air temperature datasets covering land regions, but not ocean or sea ice, are  
276 available at higher spatial and temporal resolutions. For example, Rhode (2013a; 2013b) use  
277 a larger number of meteorological stations than do HadCRUT4, NOAAGlobalTemp or  
278 GISTEMP, together with a statistical interpolation algorithm, to produce a monthly surface  
279 air temperature dataset at higher spatial resolution; an experimental daily analysis has also  
280 been produced. Other high-resolution datasets of air temperatures over land are available  
281 and are commonly used in climate modelling (Harris et al., 2013) and hydrological modelling  
282 (Weedon et al., 2011). Higher temporal resolution air temperatures derived from land  
283 meteorological station observations are also available, including the daily GHCN-D databank

284 (Menne et al., 2012), and the sub-daily HadISD databank (Dunn et al., 2016). Gridded  
285 temperature fields based on GHCN-D are available in the HadGHCN-D dataset (Caesar, et al.,  
286 2006) covering a time period from 1950 to present. HadISD is presented as time series for  
287 individual meteorological stations only. However, none of these latter datasets are based on  
288 homogenised data (see below).

289

290 The existing coarse-resolution global temperature datasets are widely used in global and  
291 regional climate assessments; however their utility is limited in some applications that  
292 require information at high temporal and/or spatial resolutions, such as the assessment of  
293 temperature extremes, national climate assessments, regional impact studies and validation  
294 of climate simulations from high-resolution climate models. These global temperature  
295 datasets are also often expressed in terms of temperature anomalies (temperatures relative  
296 to average conditions over some reference period), rather than in terms of absolute  
297 temperature information, which is commonly needed in these applications. EUSTACE  
298 provides products that can be used for the study of absolute temperatures, as well as  
299 providing information relevant to temperature anomalies.

300

301 Figure 1 provides an overview of the EUSTACE process and shows how different activities  
302 linked together to transform the source datasets (Appendix B) into the series of EUSTACE  
303 products (Appendix C). Source data sets were chosen to maximise our opportunity to  
304 quantify the components of uncertainty (in the case of satellite data) and the amount of  
305 historical daily information (in the case of weather station data). Wrapped around these  
306 scientific developments were interactions throughout the project with potential users.

307 Evaluation against independent reference measurements (Veal, 2019a) and comparison  
308 with other related products (Veal, 2019b) put EUSTACE work into context.

309

310 Through this development process, EUSTACE has contributed to advancing and enabling  
311 climate science in five main areas:

- 312 1) Detecting and correcting for non-climatic discontinuities in weather station series: to  
313 provide an accurate picture of variations in air temperature, measurements at  
314 weather stations have been checked for any jumps in the series and then corrected  
315 (Squintu et al., 2019a and b). Such discontinuities might have arisen from changes in  
316 the surroundings of the weather station, the instruments used, the location of the  
317 station, or the measurement procedure (Brugnara et al., 2019).
- 318 2) Estimating consistent skin temperature uncertainties: EUSTACE used satellite data on  
319 the surface skin temperature of the land, ocean and ice, obtained from European  
320 reprocessing projects with diverse approaches to estimating uncertainty. Therefore,  
321 we derived consistent uncertainty estimates for these data over all surfaces in order  
322 to use them together effectively (Ghent et al. 2019; Nielsen-Englyst et al. 2019a).
- 323 3) Estimating air temperature from satellite data: while in some locations air  
324 temperature records can span periods of a century or more, in many areas there is a  
325 lack of information. EUSTACE has helped to provide daily air temperature  
326 information by using temperature estimates from satellite measurements to boost  
327 the amount of information beyond that already available from weather station  
328 records and ships (Nielsen-Englyst et al. 2019; Høyer et al. 2018; Kennedy and Kent,  
329 2019).

330 4) Understanding the role of lakes: a number of EUSTACE studies explored various  
331 aspects of the relationship between lake surface water temperature and air  
332 temperature, demonstrating the place of lakes in the global climate system, their  
333 response to climate change and the importance of using spatially-resolved data to  
334 explore aspects of the response of lakes to climate change (Woolway and Merchant,  
335 2017; 2018; Woolway et al. 2017b, c, d; 2018b).

336 5) Estimating complete fields: EUSTACE used cutting-edge statistical methods to exploit  
337 the links between air temperature in different places and through time to estimate  
338 daily air temperatures in places and at times when neither direct measurements, nor  
339 estimates from satellite were available

340

341 Hereafter, we will briefly discuss these activities, together with the independent validation  
342 of EUSTACE products.

343

#### 344 **Detecting and correcting for non-climatic discontinuities in weather station series**

345

346 Most instrumental temperature series suffer from non-climatic artefacts (i.e. discontinuities  
347 or “breaks”; e.g., due to the relocation of weather stations, changes in the instrument  
348 shelter, changes in observation practices) which often result in sudden changes in the time  
349 series (e.g. Peterson et al., 1998; Brandsma and Können, 2006). Changes like this are not  
350 often adequately documented, so we need to use an automated method to detect them  
351 that we can apply to a global dataset. Correcting for these changes is termed

352 “homogenisation”. Until recently, homogenisation efforts have mostly addressed the  
353 monthly or annual time scales and have only adjusted shifts in the mean value. This is not  
354 sufficient when dealing with daily data as inhomogeneities can affect not just the mean, but  
355 the entire distribution of variables (Trewin, 2013). The effects of, for example, shelter  
356 changes on temperature depend non-linearly on the ambient weather conditions such as  
357 sunshine and wind.

358

359 Homogenisation of daily and sub-daily data has received more attention in recent years (e.g.  
360 Aguilar et al. 2008), but efforts are still rare compared to work on monthly data (Venema et  
361 al. 2012). Existing methods correcting daily or sub-daily temperature data can be grouped  
362 into three basic categories:

363 1) Corrections of the mean: Methods that start from monthly mean break sizes (i.e.  
364 sizes of non-climatic discontinuities), which are then distributed to individual days.  
365 Daily corrections are computed by fitting a spline or piecewise linear function  
366 between monthly mean corrections (e.g. Vincent et al. 2002). This is the easiest  
367 approach, but comes with a risk that the tails of the distribution would not be  
368 properly corrected.

369 2) Corrections of higher order moments of the distribution: Methods that directly  
370 adjust the distribution of daily temperature based on a daily reference series (e.g.  
371 Trewin, 2013). This is better suited for extremes, but it requires longer and better  
372 correlated reference series than method 1).

373 3) Methods that incorporate basic physics such as the effects of radiation and  
374 ventilation on the temperature shield (e.g., Auchmann and Brönnimann 2012). This  
375 requires detailed metadata that are not usually available for large datasets.

376 Until quite recently, no global dataset of homogenised daily land surface air temperature  
377 was available. Corresponding homogenisation work was restricted to a few regions such as  
378 Canada (Vincent et al. 2002), the Mediterranean region (e.g., Brunet et al. 2006, Kuglitsch et  
379 al. 2009), Australia (Trewin, 2013) and China (Xu et al. 2013).

380

381 Most break-detection methods require highly correlated reference series. However, a non-  
382 climatic network-wide break point (e.g., the simultaneous introduction of new instruments)  
383 can be difficult to detect if reference series are from the same network. For global studies,  
384 only unhomogenised daily temperature data have been available through GHCN-Daily and  
385 other sources, which are not suitable in all locations for analysing trends in extremes, for  
386 example. Berkeley Earth have produced an experimental gridded daily temperature product  
387 over land (see a description of their method in Rohde et al. (2013a; b)), but their  
388 homogenised daily station series are not available and the analysis was constructed without  
389 directly homogenising daily data. Rather, Rohde et al. (2013 a; b) constructed fields of daily  
390 anomalies (from their monthly mean values) and added them to the existing homogenised  
391 monthly dataset.

392

393 EUSTACE has combined multiple break-detection algorithms (those of Caussinus and Mestre  
394 (2004), Toreti et al. (2012), and Wang (2008)). We applied them either to annual and semi-

395 annual averages of differences between each station and neighboring reference series (our  
396 relative tests; all methods used), or to the averages of the target station alone (our absolute  
397 test; Wang (2008) only used), in the absence of neighboring stations or if available reference  
398 series are not suitable (Brugnara et al. (2019) provides details). Using multiple methods of  
399 detecting discontinuities provides an ability to assess the robustness of the results. Figure 2  
400 illustrates the coverage of the EUSTACE station dataset and indicates the type of break  
401 detection method applied to each station (relative or absolute) and also where application  
402 of the break detection methods has not been possible because of insufficient record length  
403 (i.e., less than 10 years). A simple likelihood index is formed from a 50-member break  
404 detection ensemble and users of the EUSTACE global station dataset can select a likelihood  
405 threshold appropriate to their needs, such that the detection power is maximised whilst  
406 minimising the false alarm rate. This is the first global daily station dataset with estimated  
407 locations of non-climatic discontinuities and their likelihood, together with valuable  
408 metadata, e.g. on resolution of measurements.

409

410 In addition to break detection, the EUSTACE global station dataset has undergone other  
411 quality checks both on the air temperature measurements themselves and on reported  
412 station altitudes (Brugnara et al. 2019). Appendix C provides a link to the resulting dataset of  
413 daily mean, maximum and minimum temperature.

414

415 For European weather station series, EUSTACE has made adjustments, where possible, to  
416 reduce the impact of non-climatic discontinuities. Briefly, we used an iterated quantile-  
417 matching approach (an example of method type 2 above) to adjust the distributions of the



418 measurements, not just their means, by comparing to the measurement distributions at  
419 nearby reference stations (Squintu et al. (2019a; b) give details). The homogenisation brings  
420 the distributions before and after each station change much closer together, adjusting for  
421 the non-climatic effects of such discontinuities.

422

423 Applying the quantile matching to the whole European station dataset has an impact on the  
424 apparent trends in temperature over Europe (see Squintu et al., 2019a). Sometimes, the  
425 EUSTACE corrections increase the trend and sometimes they decrease it. Where stations  
426 previously showed negative trends since 1951, they show positive trends in most cases after  
427 homogenisation; in all cases making them more consistent with their neighbouring stations.

428

429 This is the first time that a pan-European station dataset of daily data has been  
430 homogenised to reduce the impact of non-climatic discontinuities. The homogenised  
431 European station dataset is provided separately from the global station dataset and  
432 comprises part of the European Climate Assessment and Dataset (ECA&D) product. A  
433 gridded 100-member ensemble dataset available either on a  $0.1^\circ$  latitude by  $0.1^\circ$  longitude  
434 grid or a  $0.25^\circ$  latitude by  $0.25^\circ$  longitude grid, based on the homogenised station records  
435 has also been developed as a contribution to the next version of the E-OBS dataset (Cornes  
436 et al., 2018). A two-step method (documented in Cornes et al., 2018) was used to create the  
437 ensemble: (i) the daily values were fitted with a Generalised Additive Model, to capture  
438 large-scale spatial trends and (ii) the residuals from this were then interpolated using  
439 stochastic Gaussian Random Field simulation. Appendix C provides a link to the CEDA  
440 catalogue record for these datasets of daily mean, maximum and minimum temperature.

441

442 **Estimating consistent skin temperature uncertainties**

443

444 EUSTACE uses surface temperature retrievals over land, ocean and ice based on information  
445 gathered by infra-red satellite sensors. One of our key aims is to estimate the uncertainty in  
446 our air temperature products, so first we addressed the inconsistency in the availability of  
447 uncertainty estimates for skin temperature retrievals over different surfaces. Here skin  
448 temperature is the temperature at a few microns below the top-most surface of the land,  
449 ocean or ice.

450

451 Uncertainty in surface skin temperature retrieved from satellites arises from various sources  
452 (Merchant et al., 2015):

453 1) Radiometric noise in the measurements made by the satellite sensor. This is usually  
454 the simplest component of uncertainty, and a standard “uncertainty propagation”  
455 can be applied to derive the surface skin temperature uncertainty associated with  
456 any surface skin temperature retrieval, given information about the radiometric  
457 noise. There is usually no or negligible correlation of error from this source between  
458 different surface skin temperature retrievals.

459 2) Limitations of the retrieval process would introduce uncertainty into the surface skin  
460 temperature even if the actual radiometric measurements made had zero error. For  
461 physically-derived retrievals, this component can be isolated and estimated if  
462 representative simulations of the retrieval process are available ; this is not the case

463 where purely empirical relationships are used. An important aspect of this  
464 component of uncertainty is that the errors are likely to be correlated in space and  
465 time, and therefore may not “average out” in a simple way when transforming data  
466 from finer to coarser spatio-temporal scales.

467 3) Effects that are more systematic, principally: sensor calibration (which may drift over  
468 time) and radiative transfer simulation (including the effects of imperfect instrument  
469 characterisation and incorrect surface emissivity assumptions, although sub-pixel  
470 emissivity variability over land is usually considered random despite having local,  
471 coherent structure. See Ghent et al. 2019 for further discussion of uncertainties  
472 arising from misspecification of emissivity).

473

474 In addition to the above, error is introduced into surface skin temperature estimates  
475 because of imperfect cloud detection (when infrared sensors are used, as in EUSTACE; see  
476 Bulgin et al. 2018), unrecognised atmospheric aerosol, sensor anomalies, signal  
477 contamination, geo-location error, corrupted data streams, etc. Errors arising from these  
478 contributing sources are often far from Gaussian in their distributions, with complex effects  
479 on surface skin temperature uncertainty. These uncertainties have not been quantified in  
480 EUSTACE.

481

482 For all surfaces, EUSTACE estimated uncertainties partitioned according to the correlation  
483 structure of the different contributing error sources, following the method developed by  
484 Merchant et al. (2014) and expanded in Merchant et al (2015). Uncertainties are split into  
485 those arising from uncorrelated random effects, from effects which are locally correlated

486 (these arise from atmospheric effects and/or from uncertainties in the specification of  
487 emissivity) and from effects which are correlated over large space and time scales. The  
488 derivation of uncertainties in land surface temperature is documented in Ghent et al. (2019)  
489 and in Nielsen-Englyst et al. (2019a) for ice surface temperature. Uncertainties in sea  
490 surface temperature are as calculated by Merchant et al. (2014).

491

492 Links to EUSTACE products containing these consistently-estimated uncertainties are given  
493 in Appendix C.

494

#### 495 **Estimating air temperature from satellite skin temperature**

496

497 Before we can use the satellite data to estimate air temperature, we have to understand the  
498 relationship between surface air temperature and surface skin temperature and how it  
499 varies throughout the day, by surface type and through the seasons. The challenges are  
500 different in each domain, so EUSTACE explored the relationship separately over land, ocean  
501 and ice. Based on our understanding of the factors influencing the relationship in each case,  
502 we developed multiple linear regression relationships. As well as *in situ* measurements and  
503 satellite skin temperature estimates, these use extra information to help to categorise the  
504 way the skin/air temperature relationship behaves, such as vegetation, latitude and snow  
505 cover. Inclusion of altitude was found to provide no additional skill due to a lack of high  
506 altitude weather stations, although it does affect the relationship. Wind speed has a clear  
507 influence on the relationship (Good 2016), but use of wind speed information (from a

508 dynamical reanalysis) in the regression provided no additional skill. The changing vegetation  
509 fraction information used also acts as a proxy for some other relevant surface effects, such  
510 as urbanisation, but there was no explicit attempt here to model the impact of urbanisation.  
511 The uncertainty arising from excluded effects is also not dealt with explicitly in the error  
512 model. We withheld a pre-defined set of *in situ* measurements from the regression to use in  
513 validation of the results. We then used the regression relationships to estimate air  
514 temperature when and wherever a satellite skin temperature retrieval is available, i.e. in  
515 clear-sky conditions over the period of record.

516

517 The relationship between skin and air temperature is not straightforward; Good (2016)  
518 explores this over land. Simultaneously-measured air and skin temperature vary relative to  
519 each other over the course of a day. Depending on conditions, the skin temperature can  
520 become much warmer than the air temperature when the sky is clear, but when cloud is  
521 present, the skin temperature quickly decreases to a value close to the air temperature. The  
522 daily maxima and minima in the skin and air temperatures usually occur at different times of  
523 day and the amplitudes of their diurnal cycles are often quite different. These differences  
524 also vary with season and with location. Nielsen-Englyst et al. (2019b) found a very different  
525 relationship over ice-covered surfaces in Greenland with the closest coupling between skin  
526 and air temperature occurring at noon in the summer under clear skies, when the sun  
527 warms the surface. At other times, particularly in darkness, the surface is often colder than  
528 the air above it through radiative cooling and the formation of a surface inversion layer.  
529 Under overcast skies, the surface can become warmer than the overlying air during more of  
530 the day. Spatial mismatches between satellite retrievals and *in situ* measurements mean

531 that care needs to be taken on the resolution of satellite data used to develop the  
532 relationships. Consequently, we train our regression over land on skin temperature at 0.05°  
533 latitude by 0.05° longitude resolution, as the relationship with air temperature has been  
534 shown to peak at this resolution (Sohrabinia et al. 2014). Weather stations were  
535 preferentially selected for model training if their land cover type matched the dominant  
536 land cover type in the surrounding 5° latitude by longitude area. Retrievals from infrared  
537 sensors are only available in clear sky conditions, so we might expect that to bias our  
538 understanding of the relationship. By using *in situ* measurements from both clear and  
539 cloudy conditions, we mitigate the impact of this (see Høyer et al. 2015; Nielsen-Englyst et  
540 al., 2019a; Kennedy and Kent, 2019 for details on the relationships between skin and air  
541 temperature across different surfaces).

542

543 Once a regression relationship has been derived, that relationship is used to estimate air  
544 temperature where we have skin temperature retrievals. We perform this procedure  
545 separately over land, ocean and ice and build up a global picture of air temperature based  
546 on the available satellite measurements (see an example in Figure 3). Global regression  
547 coefficients are used over land. Here, the estimation is most challenging, largely due to a  
548 lack of representative station measurements, in high altitude regions (for both daily  
549 minimum and maximum temperature) and at high latitudes and/or with high snow cover  
550 (for daily maximum).

551

552 Since we previously estimated our skin temperature retrieval uncertainties arising from  
553 components with different correlation structures, when we propagate those through the

554 regression-based air temperature estimation together with the uncertainties inherent in the  
555 estimation, we can also derive components of uncertainty in the air temperature estimates  
556 arising from random, locally-correlated and systematic effects. This means that the  
557 uncertainties in our air temperature estimates are also estimated consistently across the  
558 different surfaces and can be propagated appropriately through an application.

559

560 EUSTACE air temperature estimates from satellite are provided on a 0.25° latitude by 0.25°  
561 longitude grid in separate files for each surface (land, ocean and ice). Daily mean  
562 temperatures are provided over ocean and ice and daily maximum and minimum is  
563 provided over land. Appendix C provides access information.

564

### 565 **Understanding the role of lakes**

566

567 EUSTACE has undertaken work using both lake surface water temperature from satellites  
568 and from *in situ* measurements gathered by the project to better understand the  
569 relationship between lake surface water temperature and near surface air temperature.

570

571 Lakes can show an amplified response of summer surface water temperature to near surface  
572 air temperature variability over the lake. This amplification of response is variable, but greater  
573 for cold lakes (e.g., those situated at high latitude and high elevation) and for deep lakes  
574 (Woolway and Merchant, 2017). Over-lake atmospheric boundary-layer stability is found to  
575 be more frequently unstable, with over-lake air temperature lower than lake surface water

576 temperature, at lower latitudes (Woolway et al., 2017b). In summer, the frequency of  
577 unstable conditions decreases with increasing lake area, as a result of an increase in wind  
578 speed with lake size, affecting heat and carbon fluxes between the atmosphere and the lake.  
579 A study of Central European lakes shows variable warming rates across the year, but these  
580 lakes have warmed most in spring with significant trends seen over the last few decades  
581 (Woolway et al., 2017c). Abrupt changes seen in these lakes in the 1980s are consistent with  
582 abrupt changes in air temperature at the same time. Warming trends seen across nineteen  
583 large Northern Hemisphere lakes (Woolway and Merchant, 2018) vary significantly across  
584 lakes as well as between them. Deeper areas of large lakes exhibit longer correlation time  
585 scales of lake surface water temperature anomalies and a shorter stratified warming season.  
586 Deep areas of large lakes consequently display higher rates of increase of summer lake surface  
587 water temperature.

588

589 Wind speed has a substantial impact on stratification of lakes, which can have a greater  
590 influence than air temperature (Woolway et al. 2017d), and is a controlling factor on lake-air  
591 turbulent heat fluxes. Variations in turbulent heat fluxes over lakes have a marked seasonal  
592 cycle in some cases, with heat loss higher over large lakes and at low latitudes (Woolway et  
593 al., 2018b). The relative contribution of latent and sensible heat fluxes to the total heat flux  
594 differs between lakes and with latitude.

595

596 The relationship between lake surface water temperature and near surface air temperature  
597 is a two-way interaction. Air temperature influences lake temperature (via its role in  
598 turbulent fluxes) and the presence of a lake has an impact on the air temperature in its



599 vicinity; an impact that metaphorically has some “memory” of earlier air temperature  
600 anomalies by virtue of the thermal inertia of the lake. The lake influence can be substantial,  
601 and in some instances be in excess of 2°C. In some regions, in particular where lakes are  
602 abundant (e.g., Northern Europe), their influence on the surrounding climate needs to be  
603 considered. For EUSTACE, the key question is how the lake modifies the dynamics over time  
604 of the daily minimum, maximum, and mean air temperature in its vicinity. EUSTACE has  
605 estimated the region of influence of lakes globally, provided in the Supplemental material to  
606 facilitate the inclusion of this effect in future air temperature analyses.

607

#### 608 **Estimating more-complete fields**

609

610 Having used surface skin temperature retrievals over all surfaces of Earth to estimate near  
611 surface air temperature, we have global, but not globally-complete, fields covering the last  
612 few decades. Gaps remain due to the impact of clouds on the satellite estimates, for  
613 example. We also have over a century and a half of spatially-incomplete data from ships and  
614 weather stations. Night-only ship data were used, to avoid daytime biases, and adjusted to  
615 represent air temperature at 2 m following Kent et al., 2013. To try to complete the picture,  
616 we needed to use statistical modelling to capture information on how temperature covaries  
617 between locations. This information is contained in both the satellite estimates from the  
618 recent past and the weather station and ship measurements (Woodruff et al. 2011). The  
619 statistical modelling helps us understand unobserved regions on any given day.

620

621 The state-of-the-art in the spatial statistics research community was previously far ahead of  
622 the methods that had been introduced to the Earth sciences, both in terms of generality and  
623 computational efficiency. In particular, methods capable of propagating uncertainty from  
624 multiple input data sources and realistic modelling of uncertainty due to spatial variability  
625 had seen only very limited use in the Earth sciences.

626

627 Current methods for spatial interpolation in Earth sciences that also include statistical  
628 uncertainty estimates fall mainly into two categories: low-dimensional function  
629 representations (e.g. Banerjee et al., 2008, Wikle, 2010), and local covariance-based kriging  
630 methods (e.g. Furrer et al., 2006). Given a realistic computational effort, none of these  
631 approaches provide full quantification of uncertainties on long and short spatial and  
632 temporal scales simultaneously; low-dimensional basis methods cannot capture small-scale  
633 variability and dealing with statistical non-stationarity is challenging for covariance-based  
634 methods. New techniques for statistical spatio-temporal models have been developed  
635 recently by combining numerical methods for stochastic partial differential equations  
636 (SPDEs) with efficient Bayesian computations for Markov random fields. When combined  
637 with methods for fast computations for hierarchical statistical models (e.g., Rue et al., 2013)  
638 they can handle multiple scales as well as non-stationarity (Lindgren et al., 2011, Bolin and  
639 Lindgren, 2011), for a cost similar to that of low-dimensional models. Previously, these  
640 methods have successfully been used in ecology, epidemiology, and geology, but not until  
641 now for datasets of the size and resolution of global historical daily temperature datasets.  
642 EUSTACE development has made extensive use of these methods to create a global daily  
643 mean air temperature analysis on a  $0.25^\circ$  latitude by  $0.25^\circ$  longitude grid.

644

645 We model daily mean air temperature measurements, first, as an average of each day's  
646 maximum and minimum temperature and, second, as a combination of the true  
647 temperature plus bias terms (including accounting for locally-correlated biases in the air  
648 temperature estimates from satellite) and other errors affecting each measurement type.  
649 We then assume that the true daily mean air temperature can be modelled as a linear  
650 combination of three different components: a moving long-term average climatology; a  
651 large-scale component representing inter-annual variability and a daily, weather-related  
652 component. Each component is modelled as a linear combination of Gaussian variables and  
653 is solved conditioned on the other components, starting with the climatology. The solution  
654 is improved iteratively starting with the climatology, followed by the large-scale and then  
655 the local component, moving from the broadest and slowest scales, to the shortest and  
656 fastest. The process is then repeated. The estimation of the climatology component benefits  
657 directly from the inclusion of satellite-derived data. The time-variation of the large-scale  
658 component is informed largely by the long-term *in situ* measurements from ships and  
659 weather stations. The correlations captured by the local component benefit from both the  
660 satellite-derived and *in situ* data. Different types of errors in the input measurements are  
661 associated with the individual component to which they are most relevant. For example,  
662 station biases arising from non-climatic discontinuities are associated with and estimated as  
663 part of the large-scale component, because breaks in the station series are identified at an  
664 annual resolution. To make the computation tractable, we use a combination of local linear  
665 basis functions. These basis functions combine to describe variation in space (for the daily  
666 component) and, in some cases, also in time (for the large-scale component). The basis  
667 functions are defined on a nested triangular mesh which also helps to speed up the

668 computation. This Bayesian method allows us to represent uncertainty in the process by  
669 drawing samples from the posterior distributions of the model components. Figure 4  
670 illustrates the additional information this generates and the uncertainty in different  
671 components of the process for 1 January 2006.

672

673 We generate ten samples of possible representations of mean near surface air temperature  
674 for each day from 1 January 1850. The usefulness of the complete field is determined  
675 strongly by the availability of measurements to constrain the analysis. Therefore, where we  
676 have estimated values which add no additional information (as defined by climatology or  
677 large-scale uncertainties greater than a threshold), we mask these out of the analysis (white  
678 areas in top right panel of Figure 4). In addition, in a few limited areas the statistical model  
679 produced extreme climatological values; these were also masked. Consequently, the  
680 analysis is not globally-complete.

681

682 The purpose of EUSTACE is to provide information on daily near surface air temperature to  
683 enable assessments of vulnerability to its daily variations, rather than for monitoring of  
684 large-scale changes on longer timescales. Nonetheless, it is important to know how the  
685 global analysis compares to data sets developed for large-scale monitoring. The upper  
686 panels of Figure 5 shows regional annual average near surface air temperature anomaly in  
687 the EUSTACE global analysis v1.0 since 1850 for Europe and North America, together with  
688 the same quantity in: a blend of CRUTEM4 (Jones et al., 2012) and HadNMAT2 (Kent et al.,  
689 2013); NOAA GlobalTemp (Smith et al., 2008; Vose et al., 2012); GISTEMP (Hansen, 2010);  
690 and Berkley Earth (Rohde et al., 2013a and b). From 1895 onwards, the data sets agree

691 closely. Prior to 1895, there are very few daily station measurements in the EUSTACE global  
692 station data set, so the EUSTACE analysis v1.0 relies on night marine air temperature to infer  
693 values over Europe. This causes a discrepancy in the EUSTACE analysis when compared to  
694 the global surface temperature monitoring data sets, which are themselves instead based  
695 on monthly weather station values. Monthly average data are more plentiful for the late  
696 nineteenth century, having been digitised separately from daily values. Over North America,  
697 the agreement is good back to 1870.

698

699 More pertinent to the aims of EUSTACE is the ability of the global analysis v1.0 to represent  
700 the evolution of daily near surface air temperature at a particular location. Having withheld  
701 a large number of station records from the development of the analysis, we can examine  
702 how the analysis compares to these records over the course of example years. The lower  
703 panels of Figure 5 show this for Cimbaj, Uzbekistan in 1975 and for Fort Nelson, Canada in  
704 2003. The station records for these locations were not included in the analysis so provide an  
705 independent comparator. The uncertainty in the analysis is larger for Cimbaj than for Fort  
706 Nelson (shown by the envelope around the EUSTACE analysis v1.0 time series). Nonetheless,  
707 in both locations, the analysis compares well on a day-to-day basis with the record of daily  
708 mean near surface air temperature from GHCN-D v3.26. In particular, we see that the gaps  
709 in the Fort Nelson record for 2003 are completed by the EUSTACE analysis method, which  
710 uses information from other weather station records and air temperature estimated from  
711 satellite to infer the missing values.

712

713 The EUSTACE prototype global daily air temperature ensemble is openly available via the  
714 CEDA archive (see Appendix C).

715

## 716 **Validation**

717

718 The EUSTACE daily air temperature estimates (both the air temperatures estimated from  
719 satellite and the global analysis) were matched with withheld validation measurements  
720 from land stations, ice stations, moored buoys, ships and ice buoys. These data were  
721 excluded from both the derivation of regression relationships between skin temperature  
722 retrievals from satellite and air temperature and from the production of the global daily  
723 analysis fields. Veal et al. (2019a) presents the full evaluation, but Figure 6 summarises the  
724 results for the EUSTACE global analysis.

725

726 Over ocean, the EUSTACE global analysis v1.0 performs well over the period 1850-2015,  
727 with a global median discrepancy (robust standard deviation, RSD) of +0.00 K (1.76 K)  
728 against withheld ship measurements (Woodruff et al., 2011) adjusted to a height of 2 m.  
729 The highest discrepancies (analysis minus validation data) are found in the Southern Ocean,  
730 although matchups are sparse here. The global analysis also performs well in most land  
731 regions with a global median discrepancy (RSD) against weather station measurements of -  
732 0.23 K (1.76 K), however seasonal median discrepancies over central Asia are high, 6-10 K in  
733 winter at some stations (these most erroneous data have been masked out of the final  
734 product). Over permanent ice domains, the global analysis performs less well, especially

735 over sea-ice: regional median discrepancies (RSDs) against ice buoy data are +1.19 K (4.60 K)  
736 in the Arctic and +4.76 K (6.81 K) in the Antarctic; note that these latter two statistics are  
737 affected by the sparsity of *in situ* measurements against which to compare the EUSTACE  
738 analysis in these regions, but are dominated by a drift over the Poles in the analysis which  
739 has largely been masked out of the final product. The regional median discrepancies (RSDs)  
740 over land-ice (including the Antarctic ice-shelf) against weather station data are lower:  
741 +0.37 K (4.04 K) in the Arctic and +0.47 K (2.68 K) in the Antarctic.

742

743 In addition, estimates of uncertainty are also evaluated using the withheld data. The  
744 uncertainty estimates are assessed by first binning the matchup discrepancies by the value  
745 of the uncertainty on the EUSTACE temperature estimate. Matchup statistics (median and  
746 RSD of the matchup discrepancies) are calculated for each bin. The matchup discrepancy has  
747 contributions from the uncertainty in the *in situ* reference data as well as the uncertainty on  
748 the EUSTACE temperature estimate. There is also a contribution from matching two  
749 different spatial scales, i.e. a point *in situ* value with the EUSTACE 0.25° grid box estimate.  
750 The expected match up variance can be modelled as the sum of the squares of these  
751 contributions. The actual and modelled matchup discrepancy variances are plotted in Figure  
752 7. Assuming our estimates of the uncertainty in the reference data and the matchup process  
753 are good then, if the EUSTACE uncertainty estimates are also good, for each bin the  
754 matchup RSD (blue bar) should match the modelled value (dashed line). If the blue bars are  
755 higher than the dashed line then the matchup discrepancy RSD exceeds the modelled value,  
756 indicating that the EUSTACE uncertainty estimate is too low. The uncertainty estimates for  
757 the EUSTACE global analysis v1.0 show little agreement with expectation over ocean

758 (overestimated and showing little variation with actual discrepancy), but good agreement  
759 over land. Since the EUSTACE analysis validates extremely well in comparison to withheld  
760 data over the ocean, this mitigates the impact of the less-effective uncertainty estimates  
761 here. Analysis uncertainties are underestimated over ice regions, particularly in the  
762 Northern Hemisphere and over Southern Hemisphere land ice; here, this arises from  
763 assumptions in the analysis method about the correlation structure of errors in the over-  
764 sampled air temperature estimates from satellite.

765

766 The EUSTACE matchup data base is available for non-commercial use (see Appendix C for  
767 details).

768

#### 769 **Priorities for future work**

770

771 EUSTACE relies on good retrievals of surface skin temperature from infrared satellite  
772 instruments. Adequate removal of values contaminated by cloud between the surface and  
773 the sensor is crucial for accurate skin temperature retrieval, but also for correct estimation  
774 of uncertainties and for accurate estimation of air temperature from skin temperature. The  
775 skin temperature datasets currently used in EUSTACE are sporadically contaminated by  
776 uncleared clouds. Use of improved satellite retrievals will improve the EUSTACE products.

777

778 As a proof-of-concept, EUSTACE has demonstrated that inclusion of air temperatures  
779 estimated from satellite enables the more-stable estimation of the climatological



780 component of the global analysis (where biases in air temperature estimates from satellite  
781 are not large or there are sufficient *in situ* measurements to inform their correction), as  
782 compared to use of *in situ* measurements alone. Use of longer satellite datasets would  
783 improve the amount of information available to the analysis and improve results further.  
784 Since the inputs to the EUSTACE analysis were fixed, more satellite data have become  
785 available (i.e. version 2 of the Arctic and Antarctic Ice Surface Temperatures from thermal  
786 infrared satellite sensors (AASTI) dataset over ice, Globtemperature land surface skin  
787 temperature from a further Moderate Resolution Imaging Spectroradiometer sensor, and  
788 stable sea surface temperatures from the Advanced Very High Resolution Radiometer series  
789 in the ESA SST CCI v2.1 dataset).

790

791 With more satellite skin temperature information would come the possibility of developing  
792 and applying regionally-varying regression relationships over land. EUSTACE air temperature  
793 estimates from satellite over land currently employ a global relationship determined by  
794 latitude, snow cover and fractional vegetation cover; this results in some (sometimes large)  
795 regionally-varying biases in the resultant air temperature estimates, which are reduced in  
796 the global analysis through the additional statistical modelling undertaken there and the  
797 inclusion of measurements made *in situ*.

798

799 Interactions with users have demonstrated that information on daily maximum and  
800 minimum temperatures are needed in addition to the daily mean. Although EUSTACE  
801 undertook modelling work to enable the production of a global analysis of maximum and  
802 minimum via the mean and the diurnal temperature range, it proved impossible to pull it

803 through into production within the timeframe of the project. Methods developed  
804 demonstrate promise and have applicability beyond surface temperature diurnal  
805 temperature range to other non-Gaussian variables. These prototyped methods would also  
806 enable full propagation of components of uncertainty with different correlation length  
807 scales through to the final analysis; the current EUSTACE global analysis simplifies the  
808 assumptions made to enable the calculations, but consequently results in underestimated  
809 uncertainties, especially over polar regions where satellite data are plentiful.

810

811 Pull-through of the lake influence mask (see Supplemental material) as a covariate (as  
812 distance from coast or altitude are currently specified) in the EUSTACE global analysis has  
813 the potential to improve the air temperature fields local to large lakes (with an influence on  
814 the scale of the EUSTACE grid box or larger, i.e.  $0.25^\circ$  in latitude and longitude).

815

816 The availability of daily measurements made *in situ* could be increased substantially by  
817 continuing the current international data rescue and digitisation efforts (see Brönnimann et  
818 al. (2018), for example) and by making these and other daily measurements openly  
819 available. Each new set of digitised data has the potential to improve a global analysis of air  
820 temperature by better constraining the statistical modelling, particularly when targeted to  
821 regions currently under-represented in the EUSTACE global station dataset (see Figure 2) or  
822 in under-sampled areas of the ocean, such as the Southern Ocean (Brönnimann et al.  
823 (2018)).

824

825 In the course of our work, we have identified the following needs to extend the current  
826 observing system: more simultaneous Voluntary Observing Ship measurements of sea-  
827 surface and near-surface air temperature (because the network is declining and provides  
828 the only means of measuring near-surface air temperature over ocean globally) and more  
829 weather station measurements of near-surface air temperature in certain surface regimes  
830 (e.g. desert, deep forest, ice, high elevation, high latitude ) to both better define the  
831 relationship between skin and near-surface air temperature there and provide more data  
832 for validation.

833

#### 834 **Summary and conclusions**

835

836 The potential for future improvements outlined above notwithstanding, EUSTACE has  
837 produced a number of novel outcomes:

- 838 • a global daily station dataset with estimated locations of non-climatic discontinuities  
839 and their likelihood;
- 840 • a pan-European station dataset homogenised to reduce the impact of non-climatic  
841 discontinuities and gridded ensemble analyses for Europe;
- 842 • consistently-estimated components of uncertainty in satellite skin temperature  
843 retrievals over different surfaces of Earth;
- 844 • air temperature estimates from satellite for each surface (land, ocean and ice) with  
845 propagated uncertainty components;

- 846       • a deeper understanding of the role of lakes in responding to and influencing  
847       surrounding surface air temperature;
- 848       • a global, multi-decadal daily analysis of surface air temperature incorporating both  
849       measurements made *in situ* and estimated from satellite data; and
- 850       • validation of products using withheld reference data.

851

852 These data have been made publicly available, where not restricted by source data licenses,  
853 both for direct use and to form the basis of future onward developments (see Appendix C  
854 for details).

855 APPENDIX A

856 The EUSTACE team

857

858 The EUSTACE consortium included 9 organisations:

- 859 1) Met Office (United Kingdom)
- 860 2) The University of Reading (United Kingdom)
- 861 3) Science and Technology Facilities Council (United Kingdom)
- 862 4) University of Leicester (United Kingdom)
- 863 5) Koninklijk Nederlands Meteorologisch Instituut-KNMI (Netherlands)
- 864 6) University of Bern (Switzerland)
- 865 7) University of Bath (United Kingdom)
- 866 8) Danmarks Meteorologiske Institut (Denmark)
- 867 9) University of Edinburgh (United Kingdom)

868

869 An External Expert Advisory Board comprised: Prof. Peter Thorne (University of Ireland,  
870 Maynooth); Dr. Elizabeth Kent (National Oceanography Centre, Southampton); and Prof.  
871 Doug Nychka (National Centers for Atmospheric Research and Colorado School of Mines).

872

873 APPENDIX B

874 EUSTACE input data

875

876 The EUSTACE data products are based on a number of input data sources, summarised in  
877 Tables A1-A3.

878

879 Table A1 here

880

881 Table A2 here

882

883 Table A3 here

884

885 APPENDIX C

886 EUSTACE products

887

888 The EUSTACE data products have been catalogued in the Centre for Environmental Data  
889 Analysis (CEDA) archive, with individual download pages pointing to the data. Two  
890 products, the European homogenised data and the gridded European dataset, which also  
891 form part of the European Climate Assessment & Dataset (ECA&D) are made available  
892 separately via ECA&D.

893

894 The EUSTACE data products and their availability and licenses are summarised in the table  
895 below.

896

897 Table A4 here

898

899 Data are made available on an open license (Open Government Licence  
900 <http://www.nationalarchives.gov.uk/doc/open-government-licence/version/3/>) where  
901 possible. For the station datasets and the matchup data base, this was not possible due to  
902 the licensing conditions of the input datasets, which meant they could only be made  
903 available for non-commercial use. These have been made available under a non-  
904 commercial license (Non-Commercial Government  
905 <http://www.nationalarchives.gov.uk/doc/non-commercial-government-licence/version/2/>).

906

907 In addition, EUSTACE has produced:

- 908 • User requirements reports;
- 909 • Product user guides, including detailed guidance on uncertainties and information
- 910 content in the products; and
- 911 • Peer-reviewed journal articles.

912

913 Links to all of these can be found on the EUSTACE website  
914 (<https://www.eustaceproject.org>).

915

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921 over ocean. We also thank Esther Conway for her work on the data management plan and  
922 Roy Mandemakers for provision of a step-by-step-guide for using the Climate4Impact portal  
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925 the manuscript.

926

927



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929

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1240 Tables.

1241

1242 Table A1. Satellite data on which EUSTACE products are based and period of data used.

1243

Satellite instrument	Satellite programme	Variables used	Data producers
Along Track Scanning Radiometer (ATSR) series, 1991-2012	ESA	Sea surface temperature at 0.2m depth on 0.25° latitude by 0.25° longitude grid	ESA CCI SST, experimental v1.2 (A) ATSR Level 3C data product. See Appendix C for data access.
Advanced Very High Resolution Radiometer (AVHRR) series, 2000-2009	NOAA	Ice surface skin temperature on instrument swath	AASTI v1.0 dataset generated by Met Norway and DMI within the NORMAPP and the NAACLIM projects. See Appendix C for data access.

<p>Moderate Resolution Imaging Spectroradiometer (MODIS) Aqua + Terra, 2000-2016</p>	<p>NASA</p>	<p>Land surface skin temperature on instrument swath</p>	<p>USGS/NASA (via ESA GlobTemperature). MODIS Collection 6 radiances downloaded from the NASA Level-1 and Atmosphere Archive &amp; Distribution System Distributed Active Archive Center [<a href="https://ladsweb.modaps.eosdis.nasa.gov/">https://ladsweb.modaps.eosdis.nasa.gov/</a>]. See Appendix C for data access.</p>
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1245

1246 Table A2. Weather station air temperature measurements on which EUSTACE products are  
 1247 based and period of data used.

1248

Dataset	Link	Reference
Global Historical Climatology Network – Daily (GHCN-D), version 3.22, 1850-2015	<a href="http://doi.org/10.7289/V5D21VHZ">http://doi.org/10.7289/V5D21VHZ</a>	Menne et al., 2012
International Surface Temperature Initiative (ISTI), v1.00 stage 2, 1850-2015	<a href="http://www.surface temperatures.org/databank">http://www.surface temperatures.org/databank</a>	Rennie et al., 2014
European Climate Assessment & Dataset (ECA&D), 1950-2015	<a href="https://www.ecad.eu/">https://www.ecad.eu/</a>	Klein-Tank et al., 2002
Data rescued by ERA-CLIM project, various		Stickler et al., 2014
DECADE project, 1931 onwards	<a href="http://www.geography.unibe.ch/research/climatology_group/research_projects/decade/index_eng.html">http://www.geography.unibe.ch/research/climatology_group/research_projects/decade/index_eng.html</a>	Hunziker et al., 2017
Southern Alps homogenized, 1871-2015		Brugnara et al 2016
Data from the national weather service of Argentina	Servicio Meteorologico Nacional Argentina	

1249

1250 Table A3. Marine *in situ* measurements on which EUSTACE products are based and period of  
1251 data used.

1252

Dataset	Link	Reference
HadNMAT2 observations, derived from ICOADS release 2.5.1, 1850-2010	<a href="http://www.metoffice.gov.uk/hadobs/hadnmat2/">http://www.metoffice.gov .uk/hadobs/hadnmat2/</a>	Kent et al., 2013

1253

1254 Table A4. EUSTACE products and their access and licensing information

1255

Short name	Descriptive name	Dataset link	License
Satellite skin temperatures			
Global satellite land surface temperature, v2.1	EUSTACE / GlobTemperature: Global clear-sky land surface temperature from MODIS Aqua on the satellite swath with estimates of uncertainty	<a href="http://catalogue.ceda.ac.uk/ui/d/Of1a958a130547febd40057f5ec1c837">http://catalogue.ceda.ac.uk/ui d/Of1a958a130547febd40057f5 ec1c837</a>	Open

	components, v2.1, 2002-2016		
	EUSTACE / GlobTemperature: Global clear-sky land surface temperature from MODIS Terra on the satellite swath with estimates of uncertainty components, v2.1, 2000-2016	<a href="http://catalogue.ceda.ac.uk/uuid/655866af94cd4fa6af67809657b275c3">http://catalogue.ceda.ac.uk/uuid/655866af94cd4fa6af67809657b275c3</a>	Open
Global satellite ice surface temperature, v1.1	EUSTACE / AASTI: Global clear-sky ice surface temperature from the AVHRR series on the satellite swath with estimates of uncertainty components, v1.1, 2000-2009	<a href="https://catalogue.ceda.ac.uk/uuid/60b820fa10804fca9c3f1ddfa5ef42a1">https://catalogue.ceda.ac.uk/uuid/60b820fa10804fca9c3f1ddfa5ef42a1</a>	Open
Global satellite sea surface	EUSTACE / CCI: Global clear-sky sea surface temperature from the (A)ATSR series at 0.25	<a href="https://catalogue.ceda.ac.uk/uuid/b8285969426a4e00b74814342">https://catalogue.ceda.ac.uk/uuid/b8285969426a4e00b74814342</a>	Open

temperature, v1.2	degrees with estimates of uncertainty components, v1.2, 1991- 2012		
Surface air temperatures from <i>in situ</i> measurements			
European station measure- ments	EUSTACE/ECA&D: European land station daily air temperature measurements, homogenised	<a href="https://catalogue.ceda.ac.uk/uuid/81784e3642bd465aa69c7fd40ffe1b1b">https://catalogue.ceda.ac.uk/uuid/81784e3642bd465aa69c7fd40ffe1b1b</a>	Non- commercial use only
Global Station Measure- ments	EUSTACE: Global land station daily air temperature measurements with non- climatic discontinuities identified, for 1850-2015	<a href="http://catalogue.ceda.ac.uk/uuid/7925ded722d743fa8259a93acc7073f2">http://catalogue.ceda.ac.uk/uuid/7925ded722d743fa8259a93acc7073f2</a>	Non commercial use only
Validation match up database, v1.0	EUSTACE: coincident daily air temperature estimates and reference measurements, for validation, 1850-2015, v1.0	<a href="https://catalogue.ceda.ac.uk/uuid/4b34a2c6890f4e518cacc88911193354">https://catalogue.ceda.ac.uk/uuid/4b34a2c6890f4e518cacc88911193354</a>	Non- commercial use only

E-OBS	EUSTACE / E-OBS:  Gridded European  surface air temperature  based on homogenised  land station records  since 1950	<a href="https://catalogue.ceda.ac.uk/uuid/b2670fb9d6e14733b303865c85c65d">https://catalogue.ceda.ac.uk/uuid/b2670fb9d6e14733b303865c85c65d</a>	Non  commercial  use only
Surface air temperature estimates from statistical analysis			
Air  temperature  estimates  from  satellite, v1.0	EUSTACE: Globally  gridded clear-sky daily  air temperature  estimates from satellites  with uncertainty  estimates for land, ocean  and ice, 1995-2016	<a href="https://catalogue.ceda.ac.uk/uuid/f883e197594f4fbaae6edebafb3fddb3">https://catalogue.ceda.ac.uk/uuid/f883e197594f4fbaae6edebafb3fddb3</a>	Open
Global air  temperature  estimates,  v1.0	EUSTACE: Global daily air  temperature combining  surface and satellite  data, with uncertainty  estimates, for 1850-  2015, v1.0	<a href="https://catalogue.ceda.ac.uk/uuid/468abcf18372425791a31d15a41348d9">https://catalogue.ceda.ac.uk/uuid/468abcf18372425791a31d15a41348d9</a>	Open

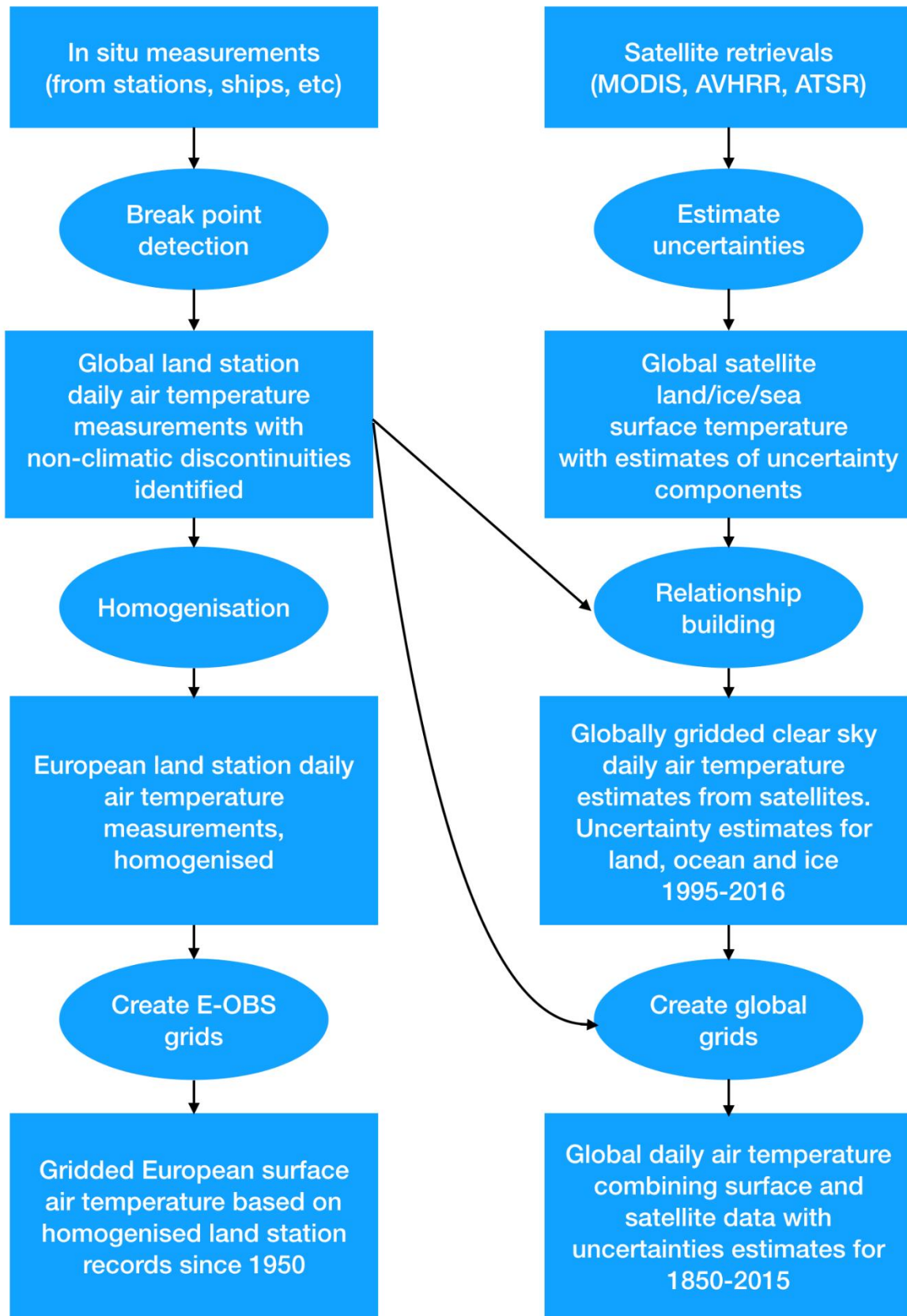
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1258 Figures

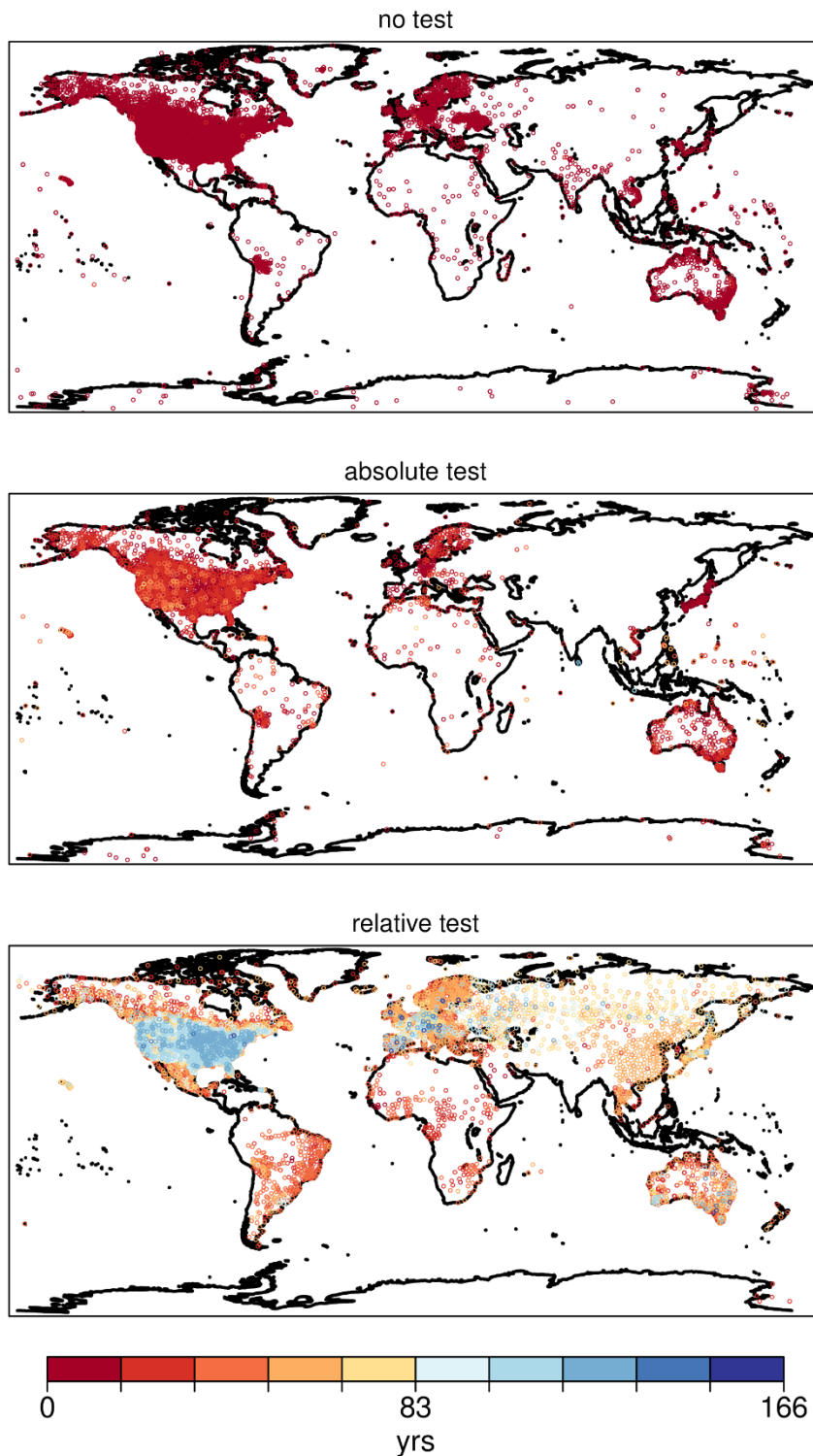
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1261 Figure 1. Schematic of work undertaken in the EUSTACE project. Top-most boxes denote  
1262 input data. Ovals denote new development. Other boxes denote EUSTACE products (see  
1263 also Appendix C). Connections between different components are indicated by arrows.

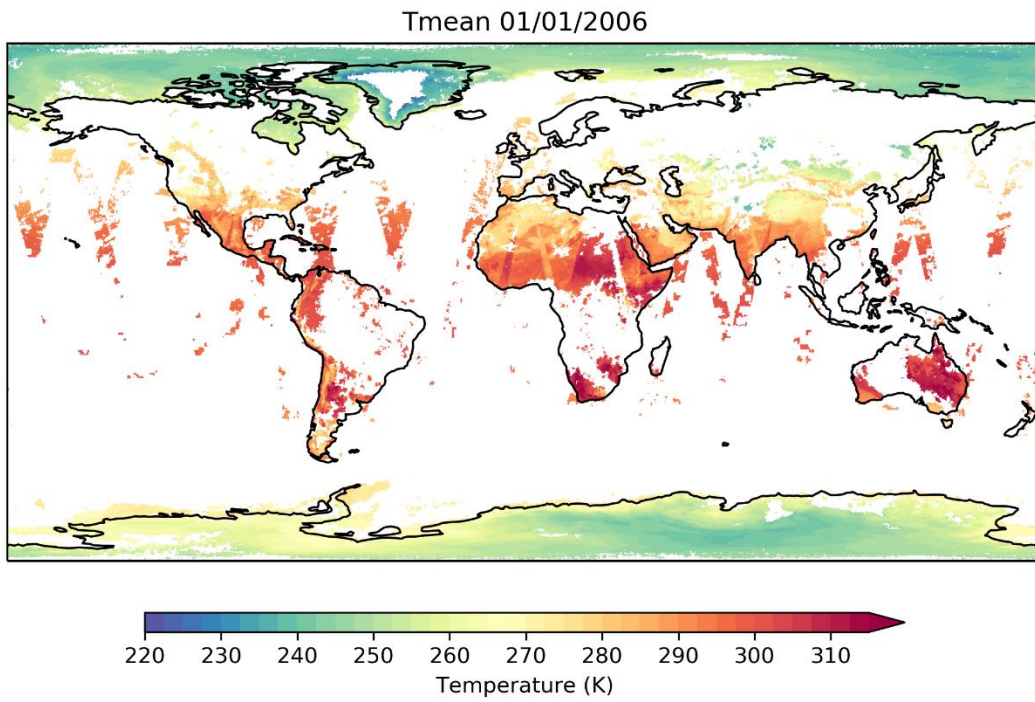
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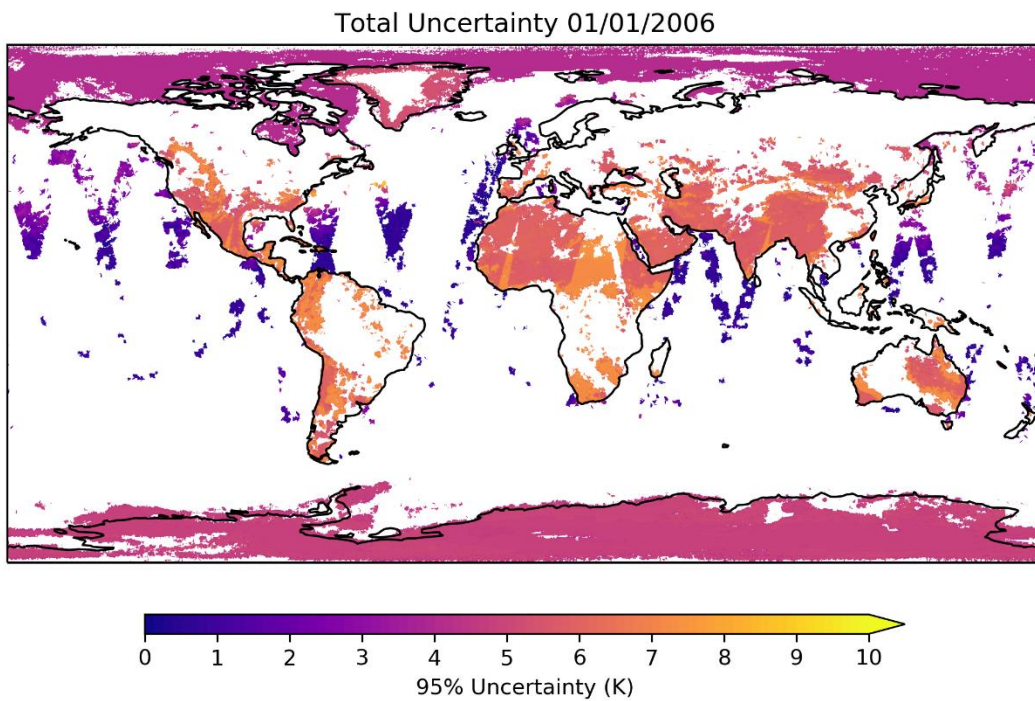
1265

1266 Figure 2. Map of weather stations included in the EUSTACE global station air temperature  
 1267 data set and break-detection tests applied (see text). Color of symbols represents length of  
 1268 daily surface air temperature record available. Top: no test applied. These stations are those

1269 which have records shorter than 10 years. Middle: only absolute test applied. Bottom:  
1270 relative test applied.

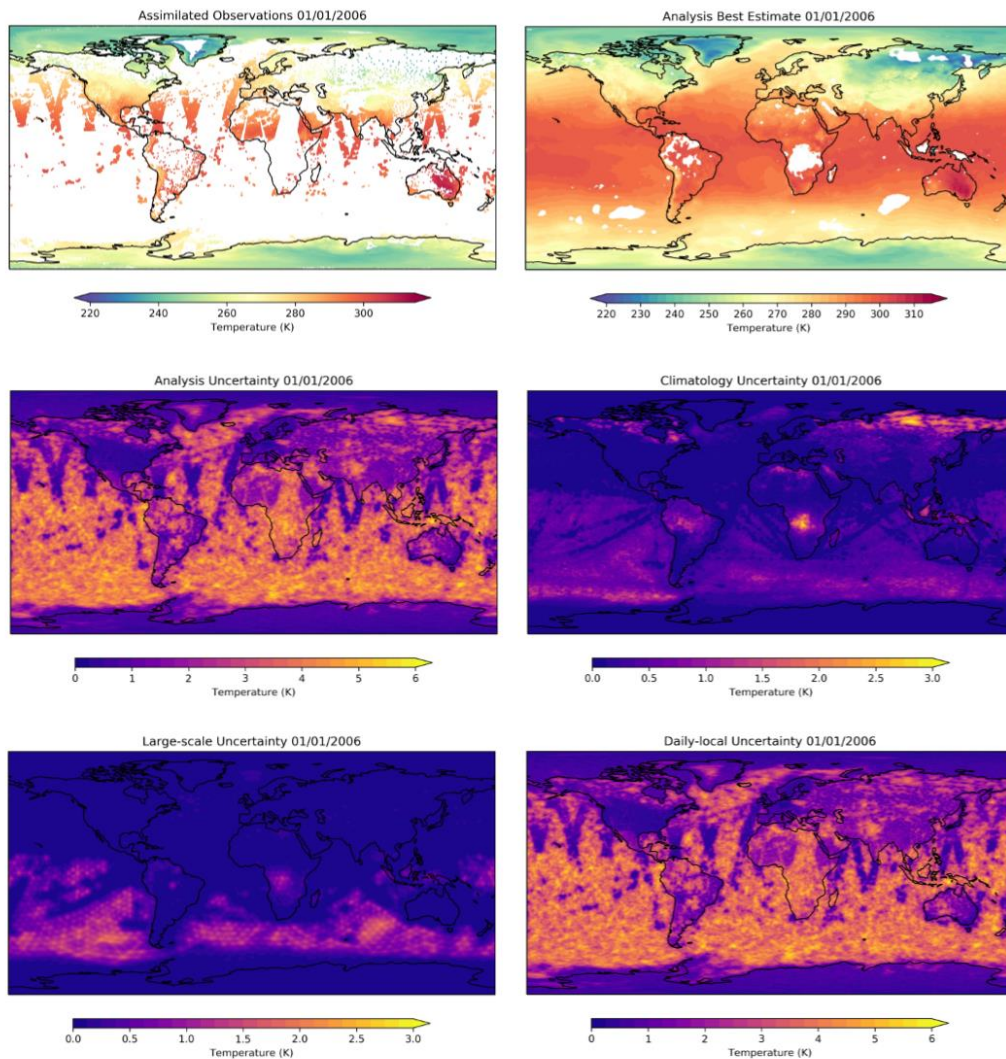


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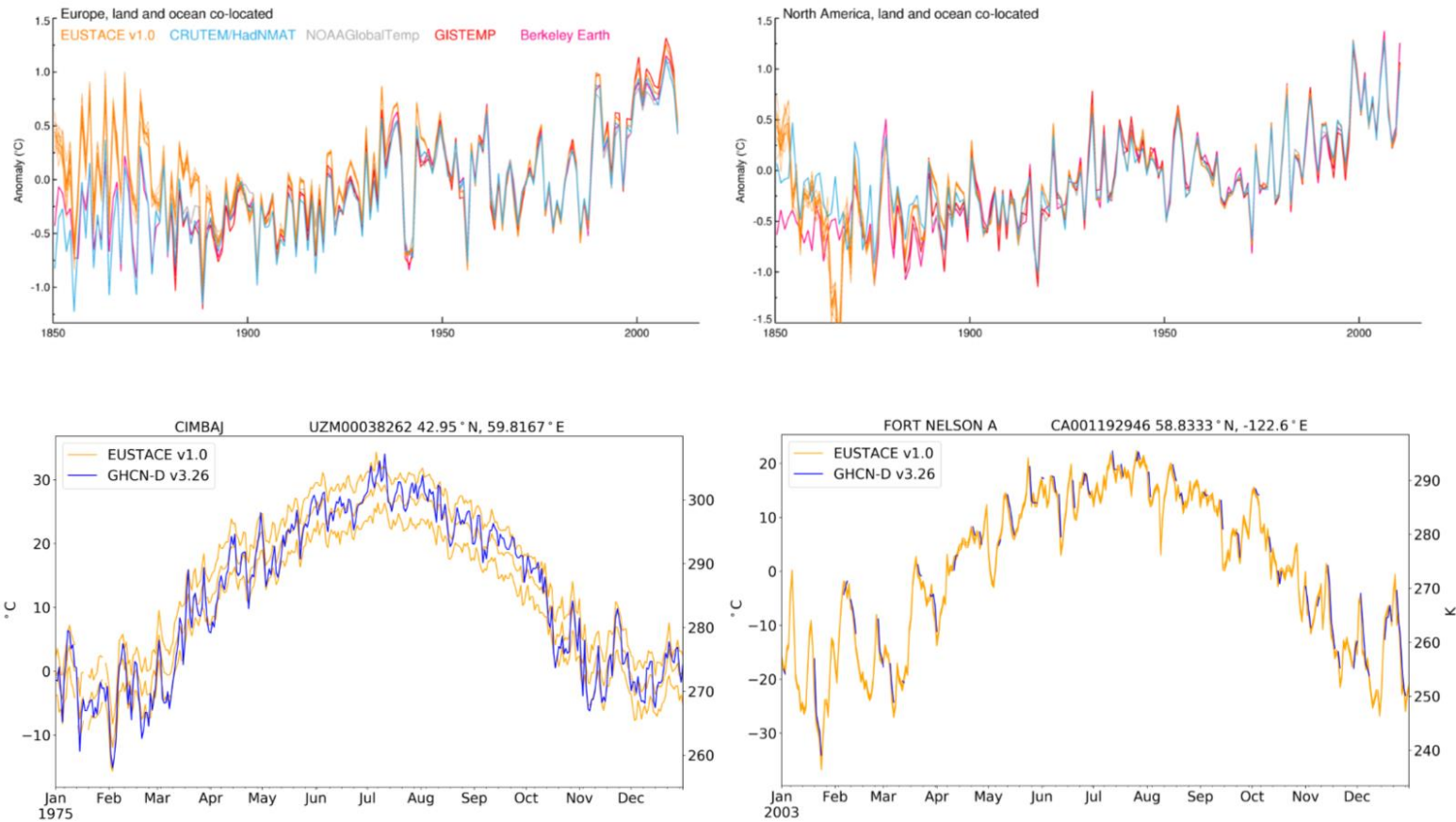
1272

1273 Figure 3. EUSTACE air temperature estimates from satellite. (Top) daily mean air  
1274 temperatures (K) estimated for 01 01 2006. (Bottom) combined uncertainty (K).



1276

1277 Figure 4. Air temperature (K) for 01 01 2006. Top left: input observations of air temperature  
 1278 (K). Top right: best guess combined *in situ* and satellite measurements from EUSTACE  
 1279 statistical infilling (K). Areas with climatology or large-scale component uncertainty above a  
 1280 threshold are masked. Middle left: total uncertainty (K) in the infilled analysis. Middle right:  
 1281 uncertainty (K) in the climatology component. Bottom left: uncertainty in the large-scale  
 1282 component (K). Bottom right: uncertainty in the local component (K).

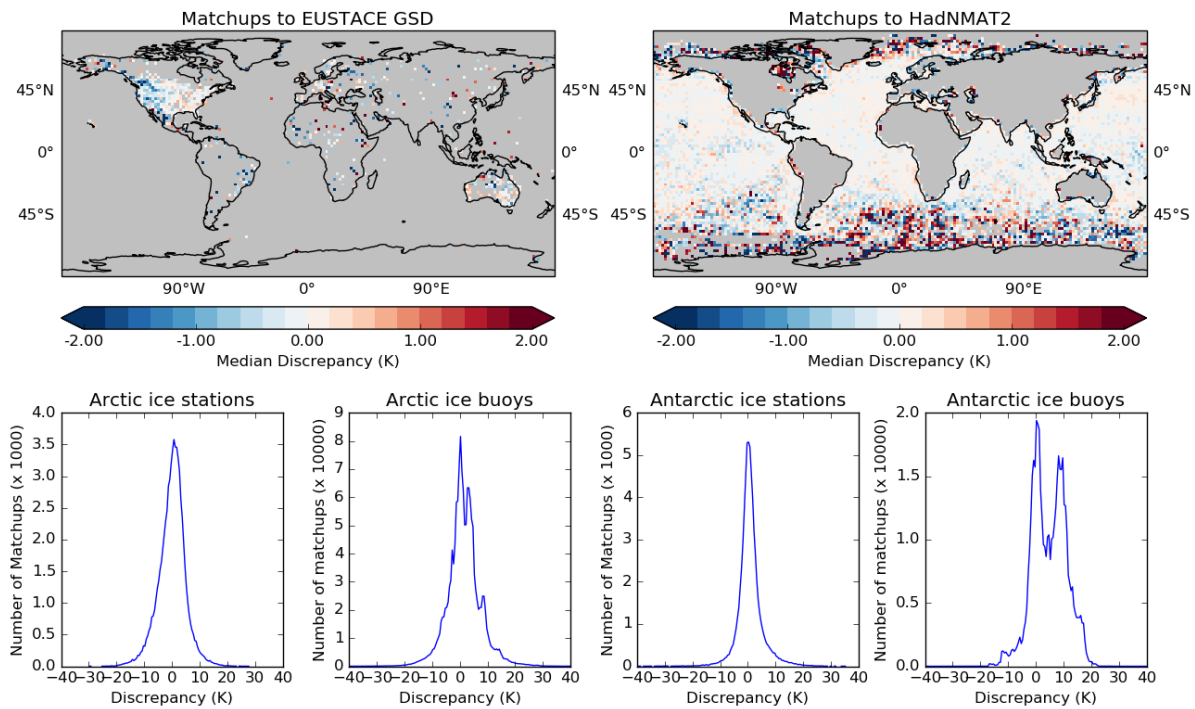


1283

1284 Figure 5. (Top) Annual regional average near surface air temperature anomaly (relative to 1961-1990) in a number of global surface  
 1285 temperature data sets, 1850-2015 (left: Europe; right: North America). Orange: EUSTACE global analysis v1.0; cyan: a blend of CRUTEM4 and  
 1286 HadNMAT2; grey: NOAAGlobalTemp; red: GISTEMP; pink: Berkley Earth. (Bottom) Daily near surface air temperature (K and °C) over the

1287 course of a year (left: Cimbaj, Uzbekistan in 1975; right: Fort Nelson, Canada in 2003). Orange: EUSTACE global analysis v1.0 (ensemble mean  
1288 and range); royal blue: GHCN-D v3.26 station measurements.

1289

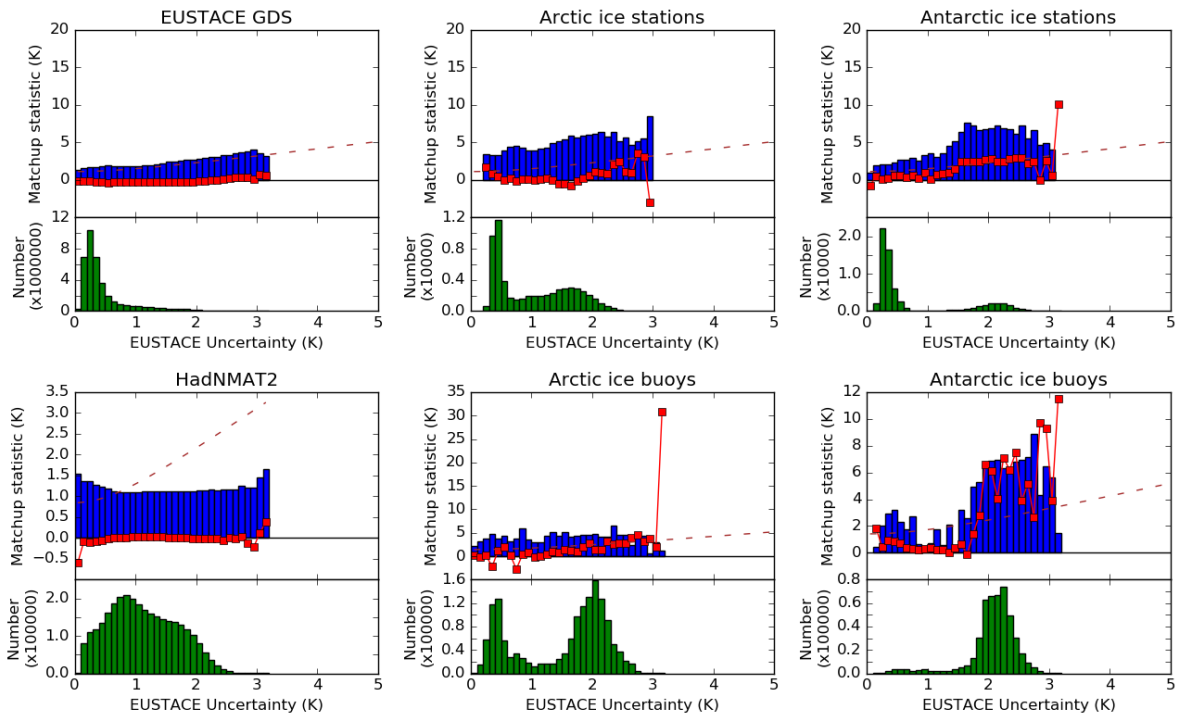


1290

1291 Figure 6. Validation of the EUSTACE global analysis v1.0, 1850-2015 against independent  
1292 reference data. (Top left) median discrepancy (K) over land, compared to withheld station  
1293 measurements. (Top right) median discrepancy (K) over ocean, compared to withheld ship  
1294 measurements corrected to 2m. (Bottom row, left to right) discrepancy (K) between  
1295 EUSTACE analysis and withheld reference data over ice-covered regions: Arctic land; Arctic  
1296 sea ice; Antarctic land and Antarctic sea ice.

1297





1298

1299 Figure 7. Validation of the uncertainty estimates for the EUSTACE global analysis v1.0, 1850-  
 1300 2015, against independent reference data. Top left: land; top middle: Arctic land ice; top  
 1301 right: Antarctic land ice; bottom left: ocean; bottom middle: Arctic sea ice; bottom right:  
 1302 Antarctic sea ice. Dashed line: modelled discrepancy; combined EUSTACE uncertainty and  
 1303 uncertainty in the validation data (K). Blue bars: robust standard deviation of discrepancies  
 1304 between the analysis and the validation data (K). Red line: median discrepancy (K). Green  
 1305 bars: number of matchups.